

***working draft for community input***

# **Ice Drilling Program**

# ***draft* Long Range Science Plan**

# **2023-2033**

Prepared by the U.S. Ice Drilling Program in collaboration with its  
Science Advisory Board and with input from the research community

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## 1 **Executive Summary**

2 *The Executive Summary will be written when the rest of the report is finalized.*

## 3

## 4

## 5 **Ice Coring and Drilling Science Goals**

### 6

### 7 **I. Past Climate Change**

8

9 Earth's climate system involves local, regional, hemispheric, and global phenomena. It is impossible to

10 understand global climate without understanding both individual components of the system and the

11 system as a whole, as evidenced by data from a large number of locations and over a range of time scales.

12 Issues articulated by many U.S. scientists (e.g., ICWG, 2003) were central to the themes in the

13 International Partnership in Ice Core Sciences (IPICS) white papers (Brook and Wolff, 2006); hence a

14 number of the categories below reflect those themes.

15

16 **1. Industrial and Instrumental Period:** Spatially distributed evidence from ice cores spanning the

17 industrial and instrumental period (last 100-200 years) is still needed to establish human impacts on the

18 climate, cryosphere and atmosphere, study modern surface processes, and calibrate models and remote

19 sensing data with in situ data. As shallow ice cores (generally <200 m), these records are relatively easy

20 to recover and consequently more records can be collected to evaluate spatial patterns of change.

21

22 Over the past 200 years, human activities have had significant impact on atmospheric composition and

23 climate, yet the impacts in polar and remote high-latitude and high-altitude regions are not fully

24 understood. Shallow ice coring programs have been, and will continue to be done through individual or

25 small-group projects at targeted sites (e.g., ice coring in mid-latitude temperate glaciers or in selected

26 areas of the Arctic and Antarctic such as Summit and Disko Bay Greenland) and internationally

27 coordinated scientific traverses (e.g., International Trans-Antarctic Science Expedition, Norwegian-U.S.

28 Scientific Traverse of East Antarctica). These shallow coring arrays must also be updated in coming years

29 since many ice cores, including the International Trans-Antarctic Scientific Expedition cores, were

30 established in the late 1990s to early 2000s and are thus missing the last 20 years of information.

31 Returning to sites and taking firn cores to depths of 10-15 would provide such an update. While shallow

32 coring has been done in a number of locations, more cores are needed in order to understand whether

33 observed patterns are regional, hemispheric, or global. Through a combination of over-snow science

34 traverses and coordinated individual site efforts, an extensive array of relatively easy-to-recover ice core

35 records, driven by individual and group proposals, is a mainstay of the ice coring community with the

36 following objectives:

- 37 • Determine accumulation rate and temperature changes on the Greenland and Antarctic ice sheets
- 38 and in alpine regions where instrumental records are rare.



81 represents a critical pre-industrial baseline against which to compare 20<sup>th</sup> century changes in climate, the  
82 cryosphere, and atmospheric composition and chemistry. Existing quantitative reconstructions of climate  
83 spanning the past two millennia continue to be debated, in part due to a lack of annually-resolved records  
84 prior to 1600 B.P. in many areas, and due to the highly regional nature of many climate processes. A  
85 coordinated international effort to recover a spatial array of annually resolved and calibrated 2,000-year  
86 ice core records has several primary objectives:

- 87 ● Determining regional and high-resolution temporal patterns of temperature, precipitation, sea ice  
88 extent, and atmospheric composition and chemistry to better understand climate forcing and  
89 particularly climate feedbacks that will also operate in the near future.
- 90 ● Evaluating 20<sup>th</sup> century warming, precipitation, atmospheric circulation, sea ice, and atmospheric  
91 composition and chemistry changes in the context of the past 2,000 years.
- 92 ● Establishing the extent and regional expression of the so-called Little Ice Age and Medieval Climate  
93 Anomaly phenomena, and constraining their relationships with regional climate patterns like the  
94 North Atlantic Oscillation (NAO), Arctic Oscillation (AO), El Niño Southern Oscillation (ENSO), and  
95 Monsoons.
- 96 ● Calibrating local, regional, and global climate models against a recent but sufficiently long pre-  
97 anthropogenic period.
- 98 ● Determining the sensitivity of alpine glaciers and ice sheet margins to the relatively warm Medieval  
99 Climate Anomaly and relatively cold Little Ice Age, with implications for the impact of future warming  
100 on water resource availability and sea level rise.
- 101 ● Quantifying spatial and temporal patterns of climate-forcing mechanisms that are regionally variable  
102 (e.g., greenhouse and reactive gases, sulfate, terrestrial dust, biological material, and carbon aerosols  
103 [black and biogenic carbon]), and the record of solar variability.
- 104 ● Assessing the relative roles of anthropogenic and natural forcing on climate evolution prior to and  
105 into the industrial period.
- 106 ● Quantifying anthropogenic emission sources and levels prior to the start of the industrial revolution  
107 (1760s), such as from early metal smelting, biomass burning, or other human activities.

108  
109  
110  
111 (image here)

112 Scientific drilling on the Mt. Hunter Plateau of Denali provides a 2,000-year record of precipitation and atmospheric circulation  
113 in Central Alaska. Drilling at this site was accomplished by wind and solar energy without the need for gas-fueled generators.  
114 Photo credit: Top) *Seth Campbell, CRREL*; Bottom) *Dom Winski, Dartmouth*

115  
116 New coring associated with this effort will include Arctic, Antarctic, and mid-latitude sites. Recent and  
117 desired future U.S. or U.S./International efforts include Central Alaska Range; British Columbia (Mt.  
118 Waddington); Detroit Plateau on the Antarctic Peninsula; multiple locations on the coastal WAIS ice  
119 shelves and ice domes; the Aurora Basin in Antarctica; Hercules Dome (the 2,000-year record would be  
120 part of a deeper core); Greenland coastal ice caps, high accumulation rate sites in Greenland, and  
121 Northwest (Qaanaaq) and South Dome sites (the 2000-year record would be part of a deeper core). This  
122 list is not exclusive, but illustrates the diversity of discussions within the research community.

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**3 Large-Scale Global Climate Change:** Large changes in global climate driven by external forcing have involved significant interactions among ice sheets, carbon cycle, vegetation, dust, ocean and atmospheric circulation resulting in rapid changes in regional climate. Understanding Earth system dynamics especially in times of rapid transitions is critical to making improvements in current Earth system models for assessment of future climate. Incomplete evidence about processes and dynamics of abrupt millennial scale change and regional impacts requires evidence from ice cores to develop more complete knowledge of the underlying mechanisms. Ice cores are uniquely placed to provide the contrasting polar elements of climate in high resolution and are the only source of past atmosphere allowing measurements of greenhouse and other trace gases. Scientific challenges include the following (mainly from the International Partnerships in Ice Core Sciences (IPICS) terminations and seesaws initiative):

- Further develop ice core proxies for different aspects of the Earth system, for example to reconstruct conditions at the ocean surface (e.g., sea ice, marine biological productivity, ocean evaporation conditions), and in the boreal continental biosphere (e.g. forest fires, land ice extent).
- Continue to improve understanding of climate forcings on millennial timescales (greenhouse gases, solar irradiance, aerosols)
- Improve the absolute and relative chronologies of individual ice cores, in the ice solid and gas phases, and construct consistent multi-ice-core chronologies.
- Develop methods to synchronise ice core records to those from other palaeoclimate archives, in particular for previous glacial cycles where radiocarbon is unavailable.
- Identify further sites to complement the spatial picture and implement plans to fill the “gaps”: examples include further cores at coastal domes around Antarctica.
- Apply newer methods to improve the resolution of data from some existing sites. Quantify and understand the spatial and temporal evolution of rapid climate changes, and assess how this varies with background climate (orbital forcing, greenhouse gas concentration, land ice masses).
- Construct, using ice cores carefully synchronised to other records, the sequence of events (including forcings and responses) through several glacial-interglacial transitions at highest resolution possible.
- Use these reconstructions with Earth system modelling to provide a stringent test of mechanisms. This will require an increase in modeling capability to assess changes at sufficient resolution through multiple terminations

Under the auspices of IPICS, the international scientific community aspires to a network of ice cores. Specific U.S. contributions to this network include the completed WAIS Divide core and the South Pole ice (SPICE) core, and an upcoming core at Hercules Dome. In Greenland, two potential sites in the Northwest (Qaanaaq) and at South Dome would also contribute to IPICS goals. The projects may vary in scope and logistical needs, but many are envisioned to be drilling campaigns conducted in two or three seasons with minimal logistics. Site-specific records of climate and environmental change are the primary objective; it will not be necessary to undertake the full suite of measurements possible in an ice core, although clearly such measurements provide data for a variety of future projects. The Foro 1800 drill was used to successfully drill the South Pole Ice Core to a depth of 1,751 m (age ~ 49,000 years), and the Foro 3000 drill will be used at Hercules Dome. Individual and small group projects targeting specific aspects of

165 climate and atmospheric variability on millennial to glacial timescales can be conducted at low-  
166 accumulation sites such as Dome C, Antarctica or at ice margin sites with agile drill systems and moderate  
167 logistics requirements.  
168

169 **4. High-resolution Records of the Last Interglacial; A Warm-Earth Analog:** The last interglacial (LIG)  
170 period (~130k to 110k years ago) was warmer than present due to differences in Earth’s orbital  
171 configuration, and can provide clues about how the Earth will respond as human activities continue to  
172 force global warming. Critical science priorities for ice cores spanning the Last Interglacial include:

- 173 ● Determining whether the West Antarctic Ice Sheet experienced partial or total collapse during the  
174 LIG, and determining the extent of the Greenland Ice Sheet during this warmer time. These objectives  
175 are critical for constraining sea-level rise estimates in a warmer world.
- 176 ● Quantifying the temperature, precipitation, atmospheric circulation, and sea-ice extent of Greenland  
177 and Antarctica during the LIG.
- 178 ● Establishing whether rapid climate change events occurred during the warmer world of the LIG.
- 179 ● Determining whether the lack of an abrupt climate change during the deglacial warming (i.e. Bolling  
180 Allerod warming and Antarctic Cold Reversal) contributes to a climate “overshoot” and a warmer  
181 interglacial.
- 182 ● Comparing the evolution of the LIG with our present interglacial period, the Holocene.
- 183 ● Investigating glacial inception at the end of the LIG.

184 Existing ice core records of the last interglacial are primarily from low accumulation sites in East Antarctica  
185 are insensitive to changes in the marine segments of the ice sheets. The detailed behavior of polar climate,  
186 greenhouse gases, ice sheet size, and other earth system attributes recorded by ice cores are not well  
187 known for this period, and require high-accumulation conditions. Results from the North Greenland  
188 Eemian (NEEM) ice core in Greenland, and similar results from other Greenland ice cores, have shown  
189 that the Eemian record located there is at least partially recoverable, but not in stratigraphic order. Large  
190 volumes of ice from the last interglacial have been shown to outcrop at the surface of Taylor Glacier,  
191 Antarctica; however a complete and undisturbed stratigraphic sequence of the warming from the climate  
192 period Marine Isotope Stage 6 has yet to be recovered.

193  
194 (image here)  
195 The bubbles visible in this piece of ice from an Antarctic ice core contain carbon dioxide and other gases that were trapped in the  
196 ice when formed many thousands of years ago. Ice cores provide the only natural archive of ancient air. Photo credit: *Oregon*  
197 *State University*.  
198

199 The search for sites with unfolded ice will continue in both polar regions along with efforts to interpret  
200 folded ice; likely targets are relatively high accumulation sites in Antarctica, such as Hercules Dome, where  
201 last interglacial ice is likely to be preserved, and possible new sites in Greenland, including near the Camp  
202 Century site in Northwest Greenland.

203  
204 The U.S. community, represented by the IDP Ice Core Working Group (ICWG), has prioritized Hercules  
205 Dome as the next deep ice core site in Antarctica, due to its likely preservation of ice from the last  
206 interglacial period (Jacobel et al., 2005) and its sensitivity to a potential collapse of the WAIS (Steig et al.,

207 2015), as well as its potential to provide bubble-free ice (below the problematic bubble-to-clathrate  
208 transition zone, e.g. Neff 2014) for gas studies during the last glacial-interglacial transition.  
209

210 **5. Evidence from the ice sheet prior to 800,000 years B.P.:** Each time ice cores have extended further  
211 back in time they have revealed new facets of climate dynamics. The record from the EPICA core at Dome  
212 C extends back to just over 800,000 years, and shows that different styles of glacial-interglacial cycles  
213 occur even under superficially similar external forcing. The Dome C site was selected to recover old, but  
214 not the oldest ice. Antarctic ice sheet inception is thought to have occurred 35 million years ago, and  
215 although basal processes may have removed or altered the very oldest ice in many places, it is likely that  
216 ice older than 800,000 years is preserved in East Antarctica.  
217

218 The primary reason to seek this older ice is to further understand one of the major puzzles of climate  
219 system history: Why did the climate system change from a dominantly 41k- to a 100k- year glacial cycle  
220 about one million years ago? Numerous research objectives related to this transition, and the earlier time  
221 period, could be addressed with ice core records extending back ~ 1.5 million years, including:

- 222 ● Evaluating the CO<sub>2</sub>-climate relationship prior to 800 ka, to determine whether the change to 100-kyr  
223 cycles and/or the long-term cooling trend from 1.5 – 0.8 Ma was related to changes in greenhouse  
224 gas concentrations.
- 225 ● Clarifying whether 23k-year climate cycles are present in ice core records prior to the transition to  
226 100-kyr cycles around 1 Ma. The 23k-year cycles are not present in marine proxy records of this age,  
227 but are present in both marine records and ice cores after the transition.
- 228 ● Investigating the high-resolution nature of glacial transitions during the 41k-year world.
- 229 ● Determining if rapid climate change events like Dansgaard-Oeschger events were present during the  
230 41k-year world.  
231

232 (image here)

233 100,000-year 'sawtooth' variability in Antarctic climate over the last 800,000 years is mirrored by generally, similar variability in  
234 atmospheric carbon dioxide (as well as methane and nitrous oxide, not shown) and global ice volume inferred from deep ocean  
235 oxygen isotope records from marine calcium carbonate. Whether Antarctic climate followed the ice volume record prior to this  
236 time, when ice volume records are dominated by a 40,000-year period, is not known, neither are the mean levels of greenhouse  
237 gases and the temporal variability of those levels. *Figure from Severinghaus et al. (2013).*  
238

239 There are two complementary, but very different, ways of accessing ice older than 800,000 years. The  
240 first is drilling at very low accumulation rate sites in interior East Antarctica. This has the advantage of  
241 recovering a continuous record, which, in the younger part, can be compared to other ice cores (an  
242 important consideration for drilling at very low accumulation sites where record integrity may be an  
243 issue).  
244

245 The second method is to make use of “blue ice” sites such as Taylor Glacier (Aciego et al., 2007), Mt.  
246 Moulton (Dunbar et al., 2008) and Allan Hills (Spaulding et al., 2013; Higgins et al., 2015) where old ice  
247 may outcrop at the surface via slow ablation or be present in the shallow subsurface. Continuous records  
248 require careful site selection, however discoveries are possible from sites with easier access, through  
249 smaller and less expensive projects. A site in the Allan Hills (Kehrl et al, 2018) has been shown by ice



250 penetrating radar to likely have a continuous record to ~250 ka with several hundred more meters of ice  
251 below. Different drilling requirements are needed for the two approaches. Development of blue ice  
252 sampling techniques should continue, given the potential for large volume sampling, very old ice (see  
253 below) and the possibility that continuous ice core records will not be discovered. Consideration of sites  
254 where only old ice might be preserved (for example areas where there is no accumulation today but has  
255 been in the past) should also continue.

256  
257 The IPICS “Oldest Ice” workshop resulted in a paper (Fischer et al., 2013) describing the state of knowledge  
258 of possible oldest ice sites; although it is possible to use modeling to identify possible locations (Lieferringe  
259 and Pattyn, 2013) it was the general conclusion that more reconnaissance was needed before choosing a  
260 site. The European project “Beyond EPICA” has selected a site near Dome C for deep drilling. The site is  
261 shallower than the existing Dome C ice core site to limit the chance of basal melting. Locations near Dome  
262 Fuji are also under consideration for a Japanese led deep core. Choosing a location with confidence is still  
263 difficult; mainly due to poorly-known geothermal heat flux. Determination of the spatial variability of  
264 geothermal heat flux is critical to the identification of potential drilling sites for oldest ice. Large regions  
265 which may have the optimal site conditions for Oldest Ice, remain unexplored. The Rapid Access Ice Drill,  
266 currently in a testing phase, should be able to quickly create access holes for spatially-distributed  
267 measurements of geothermal heat flux to facilitate site selection. New and ongoing radar, laser altimetry,  
268 gravity and magnetic data from ICECAP and Antarctica’s Gamburtsev Province (AGAP) airborne surveys  
269 are helping identify potential sites, but additional observations and model calculations are needed, such  
270 as those being conducted for the COLDEX project. In Greenland, locations on the west side of the east  
271 mountain range where the first ice sheet originated might result in ice more than one million years old.  
272 Since the stratigraphy is likely to be disturbed in that area, methods for dating ice that is not in order  
273 stratigraphically should be further developed before drilling for ice older than 800,000 years in Greenland.

274  
275 Rapid sampling of or access to the near basal region of the East Antarctic ice sheet is needed for site  
276 selection for the oldest ice project because temperature and heat flow measurements are needed to  
277 constrain models of ice sheet dynamics that are needed to predict potential locations of old ice. The Rapid  
278 Access Ice Drill (RAID) would be useful for this purpose. In addition, a more agile drill that could create  
279 holes as deep as 1,000 m would accelerate discovery. There are additional complementary international  
280 efforts to explore for oldest ice sites; these include oldest ice site selection programs that involves rapid  
281 access with new tools under development, including SUBGLACIOR (Alemany and others, 2014), and  
282 ICEDIVER (Winebrenner, 2021).

283  
284 **6. Pre-Quaternary atmosphere:** The possibility that very old ice (>1.5 million years) is preserved in special  
285 environments (for example, in debris-laden glaciers) in Antarctica (e.g., Yau et al., 2015) is exciting because  
286 it would provide a window into the composition of the atmosphere and climate during times when global  
287 environmental conditions were very different from today. Such sites will likely range from blue ice  
288 locations, where drilling issues are essentially identical to those mentioned above, to debris-laden glaciers  
289 or similar environments, which will require specialized drilling equipment; for example the Agile Sub-Ice  
290 Geologic Drill (ASIG) has proven to be useful in some cases.

291

292 **7. Large-volume sampling of climatic intervals and tracers of high interest:** Rare isotopes, ultra-trace  
293 species, micro-particles, biological materials, and other measurements that have not yet been fully  
294 exploited in ice core research offer new opportunities for discovery if large volumes of ice are made  
295 available. Examples include  $^{14}\text{C}$  of  $\text{CH}_4$  to trace methane hydrate destabilization during past warming  
296 events (Petrenko et al., 2017), nano-diamonds,  $^3\text{He}$ , and micrometeorites as tracers of extraterrestrial  
297 impacts, and  $^{14}\text{C}$  of  $\text{CO}$  as a tracer for atmospheric oxidizing capacity or past cosmic ray flux (depending  
298 on site characteristics). In the case of traditional drill sites, multiple cores or replicate coring technology  
299 are needed to obtain larger sample sizes, and in situ melting has been suggested (but not yet successfully  
300 used) as a means of sampling large volumes of air from deep ice core sites. A large-volume ice coring effort  
301 was recently conducted at the Law Dome DE-08 site in East Antarctica, and promising future sites include  
302 Dome C in Antarctica and Das 2 in Greenland.

303  
304 Blue ice areas such as Taylor Glacier and Allan Hills currently provide the best opportunities for rapid  
305 collection of large samples of ancient ice without a heavy logistics burden (e.g., Buizert et al., 2014). Ice  
306 sections ranging in age from Early Holocene to 1 M-year have already been clearly identified at these sites  
307 (e.g., Higgins et al., 2015; Korotkikh et al., 2011) and are ready for access/sampling by future projects.  
308 Continued studies at these sites that would provide more detailed and complete age maps of the desired  
309 outcropping ice areas.

310  
311 Depending on the site and scientific target, a range of ice drilling and sampling tools may be appropriate.  
312 The Blue Ice Drill, Eclipse Drill, Foro 400 drill, 4" drill, hand augers, and chainsaws have all been successfully  
313 used. Continuing efforts to maintain and upgrade the capability to explore and retrieve the ice at these  
314 sites are needed.

315  
316 (image here)  
317 Caption: A large-volume ice core drilled on the Taylor Glacier ablation zone, Antarctica. Bubbles in the ice at the site contain  
318 evidence of ancient atmospheric composition. The Blue Ice Drill is an easily-transportable drill capable of retrieving quality firn  
319 cores of approximately 9.5 inches in diameter as well as quality solid ice cores of the same diameter up to 70 meters below the  
320 firn-ice transition. Photo credit: *Jeff Severinghaus*.

321  
322  
323 **8. Ancient microbial life:** Ice sheets provide chronological reservoirs of microbial cells entombed during  
324 atmospheric deposition and studies have shown that microbial DNA and viable organisms can be  
325 recovered from ice cores collected from both Greenland and the Antarctic as well as temperate glaciers  
326 (e.g., Christner et al., 2001, 2003; Miteva et al., 2004). In addition, the distributions of microbial cells  
327 themselves can serve as climatic records in deep ice (Santibáñez and others, 2018). Many questions  
328 remain regarding how these organisms survive in deep ice for tens to hundreds of thousands of years, the  
329 origin of these airborne microorganisms and what their diversity and biogeographic distribution reveals  
330 about climate during deposition. The ability to obtain larger volumes in conjunction with advances in  
331 molecular techniques such as metagenomic analyses (Simon et al., 2009) and methods that can amplify  
332 smaller quantities of nucleic acids will enable more detailed study of the genomic potential of resident  
333 microbes and available preserved organic carbon material (D'Andrilli et al. 2017a,b) and how they  
334 integrate with our understanding of ice core ecology. There is interest in investigating the physiology of

335 microorganisms recovered from ice cores to elucidate unique physiological properties that enable them  
336 to survive in ice for extended periods of time and that may offer important biotechnological applications  
337 (Cavicholi et al., 2002). For example, studies have shown novel, ultra small microbial isolates from deep  
338 Greenland glacier ice that may inform on how organisms survive energy deprivation for extended periods  
339 of time (Miteva, 2005). There is also interest in investigating the reservoir of ice core carbon, not only as  
340 paleoclimate stores of carbon cycling signatures, but also as reservoirs of reactive carbon that may directly  
341 impact surrounding environments when exposed to the atmosphere with melting, calving, and retreat  
342 (D'Andrilli & McConnell, 2021).

343  
344 Recent studies have characterized organic carbon materials within various Antarctic ice cores and shown  
345 changing carbon signals measurable from different climates (spanning back to 27,000 years ago (D'Andrilli  
346 et al., 2017a, b) and within the Holocene of Arctic and the Antarctic ice cores (Grannas et al., 2006; Xu et  
347 al., 2018; King et al. 2019; Vogel et al., 2019; D'Andrilli & McConnell, 2021). High temporal resolution  
348 organic carbon data from the WAIS Divide ice core emphasized the highly complementary nature of  
349 carbon surveys with routinely surveyed geochemical assays in paleoclimate atmospheric reconstructions  
350 (D'Andrilli et al., 2017a). Notably, the preservation paradigm of geochemicals in ice cores also extends to  
351 biological and other organic materials, therefore it will become increasingly important to characterize  
352 their concentrations and qualitative nature now, in the ice, before it melts to reconstruct our past, learn  
353 about our present, and help better predict their impacts in the future in a warming climate. The inclusion  
354 of organic material (OM) chemical characterizations in ice coring efforts improves geochemical and  
355 biological paleoclimate atmospheric composition interpretations at broad and fine scales, ice sheet  
356 carbon storage assessments and comparisons with other ecosystem OM, understanding of signatures  
357 arising from preservation mechanisms, and the ability to predict the fate of carbon in ice sheets in the  
358 Arctic and Antarctic.

359  
360 **9. Borehole Array for Spatial Variations in Climate:** Although borehole observations do not provide a  
361 detailed climate history, an array of boreholes linked to an ice core can provide information on the spatial  
362 variability in climate history for any of the ice cores mentioned above. Notably, the WATSON instrument  
363 has thus far shown a unique microscale approach to OM and microbial distribution within the ice sheets  
364 and in the future requires a joining of the microscale, fine scale, and broad scale ice core researchers to  
365 best understand small to large spatial variations of materials in ice. (Eshelman et al., 2019; Malaska et al.,  
366 2020). Discoveries from ice borehole research will continue as additional technology is developed. See  
367 also section IV.1 below.

368  
369 **Summary**  
370 Advances in understanding climate require arrays of ice cores with depths ranging from tens of meters to  
371 3,000 m, and the requirements for the coring or sampling vary. Agile drills currently at IDP-Wisconsin  
372 should be continually maintained in good working condition so that they can be used for new projects.  
373 Clean hand augers and agile drills are needed for biological studies in glaciers. The Foro 3000, capable of  
374 coring up to 3,000 m has been created in the same family of drills as the Foro 1650 (aka Intermediate  
375 Depth Drill for coring up to 1,650 m). The large-diameter Blue Ice Drill for blue ice areas was used  
376 successfully on Taylor Glacier, Antarctica, and elsewhere; with continued science attention to blue ice

377 areas as well as to large-volume sampling in general, an upgraded Blue Ice Drill may be useful. Estisol-140  
378 has replaced the IsoparK-HCFC-141b combo for use as drilling fluid in deep drilling since HCFC-141b is no  
379 longer produced. While Estisol-140 had some issues in the first season of use, changes to the drill and  
380 handling procedures in the second season mitigated many of the issues. Neither Isopar-K nor Estisol-140  
381 are “clean” enough for biological and OM chemistry measurements in ice cores. For those types of  
382 research, the outer layers of the ice cores must be shaved off manually or within the continuous flow  
383 analysis melting system, since they introduce large contaminants into the measurements (Christner et al.,  
384 2005; D’Andrilli et al., 2017b). Table 1 lists characteristics for drills needed for the areas of the science  
385 outlined above that are available through IDP. In addition to the drills in Table 1, several U.S. universities  
386 have university-owned drills that may be available by request directly to the university, for example the  
387 WISSARD/SALSA clean access drill, Roving Drill, and Shot-hole Drill at the University of Nebraska-Lincoln,  
388 and also the Rapid Access Ice Drill (RAID) at the University of Minnesota.  
389

390 **Table 1. Requirements of drills for studies of climate change.**

	Diam (cm)	Depth (m)	Ambient temp(C)	Clean coring ?	Transport type	Site occupancy	Int'l aspects	Drill Name
<200 years	5-7	Horiz.	-20	yes	Backpack	Days	US	-
<200 years	5	15	-30	sometimes	Backpack	Days	US	Hand auger
200 year	7-10	400	-50	no	Twin otter/ Lt traverse	Days/ weeks	US	Eclipse, 4", Foro
200 year	7-10	400	-5 warm ice	no	Twin otter/ Lt traverse	Days/ weeks	US	Thermo-mechanical
2k array	7-10	<1,500	-50	sometimes	Twin otter/ Lt traverse	Weeks/ month	US part of IPICS	700 Drill, Foro 1650
40k array	10+	1-3k	-50	no	Twin otter/ Herc	1-2 seasons	US or shared	Foro 1650, Foro 3000
Interglacial	10+	1-3k	-50	no	Herc	Multiple seasons	US only or US-led	Foro 1650, Foro 3000
>800k years	10+	3k	-50	no	Herc & traverse	Multiple seasons	IPICS	Foro 3000
Site selection oldest ice	2-4	<1,000	-50	no	Herc & traverse	2 days	IPICS	RAID
>800k years (blue ice)	25	5-20	-40	no	Twin otter	1-2 seasons	US	Blue Ice Drill
Pre-Quaternary atmosphere	7-25 rock-ice mix	200	-40	no	Helicopter	1-2 seasons	US	ASIG
Tracers; large diameter cores	10-25	200+	-40	no	Helicopter/Basler/Traverse	1-2 seasons	US, Australia, France	Blue Ice Drill; 4" Drill
Ancient microbial life/ organic material	25	200+	-40	sometimes	Helicopter twin otter, herc	1-2 seasons	US	SCHWD
Borehole Array	8	200 to 3.5k	-40	no	Twin Otter/Lt Travers	Week	US	RAID

391 Additional information on the drills is given in the IDP Long Range Drilling Technology Plan.

392

393 **II. Ice Dynamics and Glacial History**

394 Rapid changes in speed of fast-flowing tidewater glaciers, outlet glaciers, and ice streams observed over  
395 the past decade create an urgency to understand their dynamics. In West Antarctica, ongoing rapid loss  
396 of ice in the region of Thwaites Glacier and the Amundsen Sea is occurring, with possible accelerated loss  
397 due to ocean-driven melting at the grounding zone and nearby areas (e.g., Scambos et al., 2017). A  
398 complete retreat of the Thwaites Glacier basin over the next few centuries would raise global sea level by  
399 more than three meters. It is possible that processes such as hydrofracture and ice cliff failure could lead  
400 to a more rapid collapse of Thwaites Glacier within the next few decades. Reducing uncertainty in the  
401 projected contribution of Thwaites Glacier to sea level rise requires substantial and coordinated  
402 collaborations involving a multidisciplinary, international scientific community. The Ross Sea sector of the  
403 Antarctic Ice Sheet, with its spatially diverse and changing natural environment, can facilitate process  
404 studies through field sampling and data which will be used in evaluating the current state of the ice sheet,  
405 quantifying the glaciological and oceanographic processes that may play a role in rapid decay of the ice  
406 sheet, and interpreting past ice sheet changes from subglacial and ice-proximal geologic records to  
407 understand ice sheet sensitivity to climate forcing on different timescales. In general, predicting responses  
408 of glaciers and ice sheets to future possible environmental change requires models that incorporate  
409 realistic ice dynamics (Alley and Joughin, 2012). Ice loss on the Greenland Ice Sheet is also happening at  
410 a dramatic rate, and contains an additional 7.4 meters of sea level equivalent. Predicting dynamic ice loss  
411 of major ice streams, like the northeast Greenland ice stream is a major challenge to the international  
412 community. For both the Antarctic and Greenland Ice Sheets, understanding the history of past ice sheet  
413 change is key for pinpointing ice sheet sectors most sensitive to climate change. Measurements and  
414 observations of present day conditions are needed to develop and validate such models. Properties of the  
415 ice and the ice-bed interface exert strong control on the flow of glaciers and ice sheets. Instruments  
416 deployed down boreholes drilled to the bed are needed to collect basic data concerning the spatial and  
417 temporal distribution of ice properties, sediments, and subglacial hydrology.

418  
419 Another approach to understand future ice-sheet response to local and global climate is to reconstruct its  
420 history. Histories of ice dynamics (thinning and divide location) and climate (accumulation and  
421 temperature) can be inferred from observations from ice cores, basal ice samples, and boreholes near ice  
422 divides. Ice core and borehole data, including basal ice samples for gas analysis, depth-profiles of age,  
423 layer thickness, temperature, ice fabric, and bubble density all provide constraints for ice flow models.  
424 For example, the depth-age relationship contains information about past accumulation and past thinning;  
425 a thin annual layer at depth could imply either low accumulation in the past or ice sheet thinning  
426 (Waddington et al., 2005; Price et al., 2007). Radar-detected layers can also be used to infer the flow  
427 history of glaciers and ice sheets and the history contained in the layers is much richer if their age is known  
428 (Waddington et al., 2007, Dahl-Jensen et al. 2013); ice cores can be used to date intersecting radar layers.  
429 The high quality radio echo sounding data from the Center for Remote Sensing of Ice Sheets (CREGIS) and  
430 Operation IceBridge both in Antarctica and Greenland make it possible to detect internal layers reaching  
431 to the bedrock. Disturbances, folding, and larger structures are observed that strongly influence the local  
432 ice dynamics and point towards the need for more complex and anisotropic ice deformation relations.

433

434 Specific observational data needed to improve and validate models of ice sheet response to  
435 environmental change include:

436

437 **1. Basal conditions and geothermal flux:** Direct measurements of bed conditions including frozen/thawed  
438 bed, basal pore water pressure, slip, and sediment properties and deformation are needed to develop  
439 and test realistic models of the controls on the fast flow of ice streams and outlet glaciers. Determination  
440 of whether a bed is frozen or thawed requires coupled thermo-mechanical flow models. A necessary  
441 boundary condition is a realistic realization of the geothermal flux. Geothermal flux has been determined  
442 at a few locations from borehole thermometry, but we expect the geothermal flux to vary significantly  
443 over spatial scales of less than 25 km (Fahnestock et al., 2001). In Greenland borehole temperature  
444 reconstructions imply low values in south Greenland (<40 mW/m<sup>2</sup>, values of 50 mW/m<sup>2</sup> at Greenland Ice  
445 Core Project (GRIP) and Camp Century and higher values at NEEM (80 mW/m<sup>2</sup>) and North-GRIP (130  
446 mW/m<sup>2</sup>). Until recently the only measurement in West Antarctica was from Siple Dome (69 mW/m<sup>2</sup>), but  
447 recent borehole temperature measurements from the WAIS Divide borehole indicate a geothermal flux  
448 of at least 90 mW m<sup>-2</sup> and possibly much higher (Cuffey et al., 2016). Measurements of ~285 mW m<sup>-2</sup> at  
449 Subglacial Lake Whillans were made (Fisher et al, 2015). Additional measurements are needed to provide  
450 boundary conditions for ice sheet models. Based on the data to date, geothermal flux values vary  
451 considerably throughout West Antarctica and further investigation is required to provide boundary  
452 conditions for ice sheet modeling.

453

454 Measurements at the bed of glaciers and ice sheets are hampered because of difficulties accessing the  
455 bed, and keeping boreholes open long enough to deploy sensors. Rapid-access drills that are portable and  
456 capable of drilling to the bed of glaciers and ice sheets in less than one field season are needed to make  
457 basic measurements including temperature, heat flux, pressure, deformation, and slip, and to sample  
458 basal sediments and bedrock. The U.S. RAID drill is a step in this direction. Hot-water drills capable of  
459 accessing the bed through 500 m to 2,500 m of ice are urgently needed. Logging tools to detect  
460 temperature, diameter, inclination, azimuth, and pressure are needed in connection with the production  
461 of ice boreholes.

462

463 **2. Remote sensing of basal conditions:** Remote sensing such as active and passive seismic arrays and radio  
464 echo sounding complement in situ measurements of bed conditions and englacial properties. Seismic  
465 imaging requires arrays of shallow holes for emplacing sources. The capability for producing large  
466 numbers of shallow holes (25 - 100 m depth, 5 - 10 cm diameter) should be maintained within IDP. The  
467 Rapid Air Movement (RAM) drill is being refurbished and enhancements for increased portability, and  
468 increased efficiencies (drilling rate, reduced size, and power consumption) are needed to improve the  
469 agility of the RAM drill for creating shot holes.

470

471 **3. Sub-ice shelf mass balance:** Ice shelves buttress discharge from ice sheets and ice sheets grounded  
472 below sea level can become unstable after their buttressing ice shelves disintegrate. Recent work  
473 indicates that ocean temperatures control rates at which the ice shelves melt, and emerging observations  
474 (Jenkins et al., 2010; Stanton et al., 2013) and model results (Favier et al. 2014; Pattyn et al., 2013;  
475 Gagliardini et al., 2010; Pollard and DeConto, 2009) indicate that sub-shelf melting exerts strong control

476 on the mass balance of ice sheets. Although measurements near the grounding line have been made and  
477 more are being conducted, coverage is still sparse. The melting process is determined by boundary layer  
478 physics that operate on spatial scales of centimeters. Access holes large enough for deploying instruments  
479 on moorings, autonomous underwater vehicles, and remotely operated vehicles are needed to acquire  
480 short-term spatially-distributed data. Additionally, long-term observatories at targeted sites are needed  
481 to document temporal variability. All these experiments should be directly related to grounding-zone  
482 studies and linked to oceanographic campaigns beyond the ice shelves.

483  
484 (image here)

485 Heat and mass exchange in sub-ice shelf cavities impact ice flow and ice sheet mass balance. Image credit: *WISSARD project*.

486

487 **4. Grounding zone processes:** Improved understanding of processes in grounding zones is needed to  
488 assess the role of fast-flowing ice streams and outlet glaciers on the stability of ice sheets. Conceptual  
489 geological models of grounding-line environments have been inferred from stratigraphic successions.  
490 Remote sensing studies using satellite observations and geophysical surveys have been conducted at  
491 grounding lines of major ice streams, and studies of processes at modern grounding lines are underway  
492 (Anandakrishnan et al., 2007; Alley et al., 2007; Horgan et al., 2013; Christianson et al., 2013). Few  
493 direct measurements or materials have been collected at grounding lines and grounding zones of fast-  
494 flowing ice streams and outlet glaciers (Begeman et al., 2018; Venturelli et al., 2020). Small diameter  
495 access holes are needed to deploy instruments to measure spatial and temporal changes in these critical  
496 areas. Shallow coring (~300 m) at Crary Ice Rise could provide a short (1000 year) climate record  
497 while also dating the timing of grounding at this site and possibly accessing marine ice-filled relic basal  
498 crevasses.

499

500 **5. Rheological properties of ice:** Rheological properties of ice depend strongly on temperature, impurities,  
501 and texture, including grain size and fabric (Cuffey and Paterson, 2010). Improved understanding of the  
502 controls on the rheology is needed to develop realistic models of deformation of ice sheets. These models  
503 are needed to help develop depth-age relationships in ice cores, understanding flow and shear, and also  
504 to establish past, present, and future responses to possible environmental changes. Folding of deep ice  
505 and large structures forming at the base of the ice are believed to be related to the rheological structure  
506 of ice. Studies at Siple Dome (Pettit et al., 2011, Bay et al., 2001) and Dome C (Pettit et al., 2011), for  
507 example, have shown that strong vertical gradients in the effective viscosity of ice are likely present at  
508 depth in the ice sheets. These strong variations in ice rheology have the potential to lead to folding (such  
509 as at NEEM, Dahl-Jensen et al., 2013) or the formation of shear bands. Sensors that measure depth  
510 profiles of temperature, fabric, optical stratigraphy, tilt, and borehole diameter in boreholes are now  
511 available and can be calibrated against ice core properties. Rapid-access drills that can drill through ice up  
512 to 3,000 m thick are needed to deploy such sensors. In particular, the ability to drill multiple holes along  
513 a flow line can provide key spatial changes in ice properties. In addition, a system to rapidly access the ice  
514 sheet and then extract ice cores from selected depths would allow analyses of ice properties at depths of  
515 special interest; such a drill does not yet exist but should be planned.

516



517 **6. Glacial history:** Defining the extent and volume of ice sheets under paleoclimatic conditions warmer  
518 than the present (Eemian, Marine Isotope Stage-14, Pliocene, and mid-Holocene warm periods in  
519 Greenland) is an important indicator of future ice sheet vulnerability. Although a variety of indirect  
520 approaches have been used to constrain the history of ice sheets (glacial geology, paleoceanography,  
521 etc.), the most direct method is to determine the age of basal ice across an ice sheet bed. Basal ice age  
522 can be modeled with age-depth flow models, or more directly by dating trapped air in basal ice. Slow-  
523 moving ice in the vicinity of ice divides contains a record of past ice dynamics (thinning and divide  
524 location). Depth profiles of age and temperature from ice cores and boreholes can be used to extract  
525 histories of accumulation and ice dynamics (Waddington et al., 2005; Price et al., 2007). Records from  
526 coastal domes and coastal ice caps are of special interest because they can be used to infer past extents  
527 of ice sheets and the history of deglaciation (Conway et al., 1999). Intermediate depth (~1,500 m) cores  
528 to measure depth-profiles of age and temperature at targeted coastal domes are needed to help constrain  
529 the deglaciation of ice sheets. Coring on ice domes near the Amundsen Sea Embayment may be able to  
530 provide a context for more recent observed changes in ice dynamics, particularly accelerated thinning in  
531 the most recent several decades.

532  
533 Cosmogenic nuclides in bedrock beneath ice sheets can tell us about their former extent, and the timing  
534 and duration of past exposure periods. Techniques to estimate the size and shape of ice sheets during  
535 colder periods are well established (e.g., Mercer, 1968, Denton et al., 1989, Todd et al., 2010; Bentley et  
536 al., 2010; Stone et al., 2003; Hall et al., 2004; Anderson et al., 2014; Schaefer et al., 2016); determining  
537 their extent and thickness under warmer climates is more problematic. Much of the evidence is hidden  
538 beneath the present ice sheets. Recovery of cores of basal ice for gas analysis is very useful when obtained  
539 in conjunction with basal rock cores. Under shallow ice, nimble methods for reconnaissance recovery of  
540 short rock cores for cosmogenic nuclide techniques to quantify periods of exposure (ice free) and burial  
541 (ice cover) have been developed, for example the Agile Sub-Ice Geological Drill (ASIG) and ice-enabled  
542 Winkie drill for use near the ice margins. Under very deep ice, rapid access drilling using the RAID drill may  
543 open up new perspectives on ice-climate linkages in a warmer world.

544  
545 (image here)

546 Caption: John Goodge and a colleague collecting specimens in the Transantarctic Mountains. Photo credit: *John*  
547 *Goodge/University of Minnesota-Duluth.*

548  
549 Depth profile measurements on short (1-5 m) subglacial bedrock cores will be used to confirm that  
550 cosmogenic nuclides were produced in situ, and identify surfaces that constrain subglacial landscape  
551 evolution by subglacial erosion. Erosion reduces and ultimately erases the nuclide profile, so eroded  
552 surfaces must be avoided by targeting surfaces where ice is frozen to the bed. Note, however, that small  
553 amounts of erosion can be identified and the effects constrained using combinations of nuclides with  
554 different production profiles (Liu et al., 1994). With rapid access to subglacial bedrock in which  
555 cosmogenic nuclides can be measured, key problems can be addressed, such as the vulnerability of the  
556 West Antarctic, parts of the East Antarctic, and Greenland Ice Sheets to future climate warming, Pliocene  
557 ice-sheet collapse, and the onset of continental glaciation in Antarctica. Potential targets to address the  
558 interglacial extent of West Antarctic glaciation include Mt. Resnik, a subglacial peak which rises to within

559 330 m of the surface near the WAIS divide (e.g., Morse et al., 2002), ice rises particularly along the Siple  
560 Coast such as the Crary and Steershead ice rises (e.g., Scherer et al., 1998), the subglacial roots of nunataks  
561 (rocks emerging above the ice) in the Pine Island and Weddell Sea catchments, and a variety of sites in  
562 Greenland including both interior sites (e.g., Schaefer et al., 2016) and peripheral sites that border key  
563 areas (e.g., North Greenland, NEGIS). In addition, ongoing international studies of past ice thickness  
564 variations evidenced from multiple cosmogenic nuclides on nunatak altitude transects in Dronning Maud  
565 Land could benefit from future sampling of subglacial nunatak slopes. Data from beneath high-altitude  
566 domes and plateaus in the Transantarctic Mountains, , and also from subglacial debris-rich ice from outlet  
567 glaciers that drain marine basin margins (e.g. Wilkes and Aurora basins), or from subglacial strata directly,  
568 could shed new light on the long-running debate over ice-sheet collapse in the Pliocene (e.g., Webb et al.,  
569 1984; Denton et al., 1993). A variety of isotopes with varying half-lives (half-life:  $t_{1/2}$ ) can be used to  
570 constrain long-term ice sheet stability (e.g.,  $^{37}\text{Cl}$ ,  $^{26}\text{Al}$ ,  $^{10}\text{Be}$ ), and new application of in situ  $^{14}\text{C}$  can constrain  
571 Holocene ice sheet changes. In Greenland for example, in situ  $^{14}\text{C}$  measurements from periphery ice  
572 drilling sites would provide ice sheet models with direct measures of ice sheet presence/absence during  
573 smaller-than-present ice sheet conditions during the Holocene thermal optimum. Eventually,  
574 measurements of long-lived radionuclides such as  $^{53}\text{Mn}$  ( $t_{1/2} = 3.7$  million years) and  $^{129}\text{I}$  ( $t_{1/2} = 16.7$  million  
575 years) paired with stable  $^3\text{He}$  and  $^{21}\text{Ne}$  may even provide constraints on the early Neogene onset of  
576 Antarctic glaciation, targeting samples from the subglacial Gamburtsev Mountains.

577

#### 578 **Summary**

579 Understanding present and past behaviors of glaciers and ice sheets is essential for improving predictions  
580 of future behavior of ice sheets and sea level. Improved understanding requires fast-access drilling, such  
581 as those from the IDP Agile Sub-Ice Geological Drill, the IDP ice-enabled Winkie drill, or the Rapid Access  
582 Ice Drill to enable fundamental measurements of: (i) physical conditions, including geothermal flux, and  
583 processes at the beds of glaciers and ice sheets; (ii) physical properties of the ice that affect ice flow and  
584 folding, (iii) physical processes at grounding lines and grounding zones of fast-moving ice streams and  
585 outlet and tidewater glaciers; (iv) ice-ocean interactions at grounding lines. Past responses of glaciers and  
586 ice sheets to climate and sea level change also offer clues to future possible responses. Depth profiles of  
587 age and temperature from ice cores can be used to reconstruct past thickness and extent of ice sheets as  
588 well as climate. Basal ice cores in condition suitable for trapped gas analysis provide dating of the basal  
589 ice. Recovery of subglacial sediments that record paleoclimate and paleoenvironmental proxy information  
590 provide evidence of past conditions in area now covered by ice, including interglacial conditions, and in  
591 the case of marine basin settings like West Antarctica (Thwaites Embayment and Ross Sea) or East  
592 Antarctica (Wilkes/Aurora basins and Totten Glacier), to reveal polar region paleoenvironments in  
593 warmer-than-present past scenarios, that guide model projections of future warmth scenarios.  
594 Intermediate depth (~1,000 m) cores at targeted coastal domes are needed to constrain the extent and  
595 timing of deglaciation. For reconnaissance of sites for bedrock coring under shallow ice, the use of an very  
596 agile ice coring drill (e.g, Foro or Eclipse drill) along with a method of sampling rock at the ice-rock interface  
597 using the same or slightly modified drilling apparatus, would provide a logistically agile way of site  
598 selection for subsequent rock drilling of meters-long rock cores. The collection of basal ice cores for  
599 trapped gas analysis, and subglacial bedrock for cosmogenic nuclides from both strategic periphery sites  
600 and also from the ice sheet interior can provide direct constraints on past ice sheet history.

601

602 **Table 2. Requirements of drills needed for studies of ice dynamics and glacial history.**

	Diam (cm)	Ice Depth (km)	Core or hole	Ambie nt temp (°C)	Clean access?	Transport type	Site occupanc y	Int'l Aspects	Drill Name
Bed conditions	8	1-3	Hole	-50	maybe	Twin otter/ helo/lt traverse/Herc	<4 weeks	US & others	CHWD
Geothermal flux	5-8	1-3	Hole	-50	no	Twin otter/ helo/lt traverse/Herc	<4 weeks	US & others	RAID
Geologic coring for cosmogenic samples and for basal ice samples	6-10	0.1- 2.5	Basal ice core, Rock core	-50	no	Helo sling load/ Baseler/ traverse	4-8 weeks	US	Winkie/ ASIG/ RAID
Nimble geologic coring and basal ice coring under shallow ice	3-5	<.5	Basal ice core, Rock core	-30	no	Twin otter/helo/ lt traverse	<4 weeks	US	Winkie/ ASIG
Rheological properties	8	<3k	Hole	-40	no	Herc/ traverse	<4 weeks	US & others	RAID/ CHWD
Internal layering	8-10	<3k	Hole	-40	no	Herc/ traverse	<4 weeks	US & others	RAID/ CHWD
Sub-ice shelf/ice stream instrumentation	10- 25	<1k	Hole	-30	shelf- no; stream- yes	Twin otter/ helo/ herc/ traverse	2 weeks	US & others	CHWD/ SchWD
Ice shelf ROV deployment	100	<1k	Hole	-30	no	Twin otter/ helo/Herc/ traverse	2-4 weeks	US & others	CHWD/ SchWD
Grounding zone	8-75	<1k	Hole	-30	no	Helo/Herc/ traverse	2 weeks	US	CHWD/ SchWD
Seismic imaging	5-10	~100 m	Hole	-40	no	Twin otter/ helo	Hours/da ys	US	RAM/ SHWD

603

604 Note: CHWG is the Clean Hot Water Drill owned by the University of Nebraska-Lincoln. In addition to the  
605 drills in this table, the University of Nebraska-Lincoln owns the following drills that may be available by  
606 request directly to the university: the WISSARD/SALSA clean access drill, the Roving Drill, and Shot-hole  
607 Drill. The IDP Long Range Drilling Technology Plan discusses IDP drills available for retrieving cores or  
608 creating access holes in ice sheets.  
609

### 610 **III. Subglacial Geology, Sediments, and Ecosystems**

611 Bedrock, sediments, and ecosystems existing within and beneath ice sheets remain largely unexplored  
612 because of the lack of rapid access. Rapid access to subglacial environments is needed to address a wide  
613 range of science questions. Specifically:  
614

615 **1. Bedrock geology:** The Antarctic continent and its lithospheric plate, play important but poorly  
616 understood roles in global tectonic architecture, leading to contradictory hypotheses. Antarctica is  
617 considered aseismic, but if so, it would be unique among all of the continents. Its plate is surrounded by  
618 mid-ocean-ridges, and hence should be under compression, yet there are active extensional regimes. The  
619 West Antarctic Rift System is one of the largest on Earth, and currently known attributes are unique, by  
620 having only one rift shoulder and by being largely below sea level. Fundamental questions about the  
621 Antarctic Ice Sheet persist. What is the origin of the enigmatic Gamburtsev Subglacial Mountains and how  
622 have they influenced the overlying ice sheet? What are the composition, geothermal heat flux, and  
623 geotectonic histories of East Antarctica, and how does it influence ice-sheet behavior? What were the  
624 dominant factors controlling the spatial extent and temporal variability of ice sheets during warm climate  
625 periods in the past? What is the role and history of subglacial sediments in the interior? What are the  
626 physical conditions at the base of the East Antarctic ice sheet? Constraints on composition and age of  
627 basement rocks of interior East Antarctica would place better constraints on Precambrian provinces and  
628 evolution of the Antarctic shield for verifying current models. The state of stress in basement rocks is  
629 required for evaluating seismicity and extensional regimes. Boreholes through the ice into crustal rocks  
630 are needed to conduct passive and active seismic experiments for delineating crustal structure.  
631 Continental topography is a significant control on glaciation; rising mountains and higher elevations focus  
632 snow accumulation and become nivation centers for ice sheets. Sampling bedrock to determine its age  
633 and constrain its cooling history using thermochronology is important for supercontinent reconstruction,  
634 understanding the tectonic history of the continent as well as reconstructing paleotopography for  
635 glaciological modeling of Antarctic Ice Sheet history. Access boreholes to the ice sheet bed are required  
636 to recover short rock and sediment cores for these studies. Locations should be based on best estimates  
637 of bedrock geology, bed paleotopography, and plausible ice sheet extents based on models. In Greenland,  
638 the ice sheet has waxed and waned during the last 2.5 million years. Erosion of mountains and ice sheet  
639 modeling has simulated past changes, but access to old ice and basal rocks/material is needed for  
640 verification and full understanding.  
641

642 **2. Subglacial basins and sedimentary records:** The records of glaciation and their variations in Antarctica  
643 are found in scattered terrestrial deposits and sedimentary basins and can be compared with offshore  
644 records that have been collected near the margins. Interior subglacial basins also likely contain proxy  
645 records of paleoclimate and ice sheet history to complement these records from the continental margins.

646 Four main categories of sedimentary targets are: subglacial lakes, ice rises, West Antarctic sedimentary  
647 basins, and East Antarctic marine and terrestrial basins. Each category may have a variety of origins and  
648 histories because of differing locations relative to the ice sheet margin and magnitudes of past ice sheet  
649 fluctuations. Thus, they may provide valuable archives of paleo-ice sheet and paleoclimatic changes.  
650 Subglacial lakes occur throughout the continent, the largest being subglacial Lake Vostok. Subglacial lakes  
651 contain sedimentary records; sediments have been collected at Lake Ellsworth, the Whillans Subglacial  
652 Lake (Hodson et al, 2016), and Mercer Subglacial Lake (Rosenheim et al. 2023).

653  
654 (image here)

655 Illustration showing the aquatic system that scientists believe is buried beneath the Antarctic ice sheet. Photo credit: *National*  
656 *Science Foundation, Photo Gallery.*

657

658 Subglacial ice rises can cause locally grounded “pinning points” that play an important role in buttressing  
659 the discharge of streaming ice from the ice sheet. Recovery of these sediments will provide Neogene and  
660 Quaternary paleo-environmental archives, but may also provide insights on till deformation processes  
661 downstream of the Whillans and Kamb ice streams (e.g., Scherer et al., 1988). Shallow drilling of ice rises  
662 and acquisition of oversnow seismic reflection profiles radiating away from core sites will allow the deeper  
663 geometry of the strata to be evaluated for locating deeper drilling and recovery of long, continuous  
664 records in adjacent marine basins. In West Antarctica, the stratigraphic record in various basins and  
665 probable rifted grabens may contain a mid-late Mesozoic and Cenozoic history of West Antarctic evolution  
666 and paleoclimate history. Two low regions within the Wilkes Land sector of East Antarctica (Aurora and  
667 Wilkes Subglacial Basins) appear as broad down-warped basins filled by marine and non-marine strata.  
668 They may contain evidence of the much debated past dynamics and paleoclimate of the East Antarctic Ice  
669 Sheet. Recently, Mengel and Levermann (2014) suggested that only a narrow, low coastal rim holds the  
670 portion of the East Antarctic ice sheet overlaying the Wilkes Subglacial Basin back, raising concern about  
671 ice sheet stability.

672

673 Access holes are also needed to recover longer sedimentary cores comparable to those from the  
674 continental margins. Also, the basins on the interior of the Transantarctic Mountains may be sites for good  
675 proxy records of past ice sheet dynamics. These are also excellent sites to measure geothermal heat flux  
676 to help constrain ice sheet bed conditions.

677

678 **3. Sub-ice microbial ecosystems and biogeochemistry:** Aqueous and sedimentary subglacial  
679 environments in Antarctica and Greenland are inhabited by microorganisms (e.g. Christner et al., 2014;  
680 Dubnick et al., 2017) and are a potentially large planetary reservoir of microbes and (microbially derived)  
681 organic carbon, perhaps of the same magnitude as that in the surface oceans. Modeling and direct  
682 measurements (Wadham and others 2012; Michaud and others, 2017) suggests these environments could  
683 contain large volumes of the greenhouse gas, methane, which could impact atmospheric concentrations  
684 in response to rapid deglaciation. It has also been hypothesized that the flux of dissolved elements and  
685 sediments in subglacial waters can enhance primary productivity in the marine environments they drain  
686 into (Hawkings et al., 2020; Vick-Majors et al., 2020). Elucidating the spatial and temporal distribution and

687 dynamics of these aqueous environments, including their physical and chemical properties (such as  
688 temperature, salinity, and pressure) and associated biogeochemical processes (i.e. microbial communities  
689 and organic carbon material fluxes) is key to understanding ice sheet stability and the role of large  
690 continental ice sheets in global biogeochemical cycles. The rapid changes anticipated in the size of polar  
691 ice sheets may trigger significant reorganization of subglacial hydrologic conditions, which may feed back  
692 into acceleration of ice sheet retreat and may force adaptation of subglacial biota, whose metabolism is  
693 linked to subglacial hydrology (Vick-Majors et al., 2016), to rapidly changing conditions.

694  
695 The long timescale of microbial entrapment in sub-ice environments relative to the lifetimes of microbial  
696 cells provides an opportunity to explore questions concerning rates of evolution, and constraints on  
697 biodiversity. Microbial cells and their genomic material should also provide valuable information that can  
698 be linked to paleoclimatic change; such life forms may be the only biological survivors in areas covered by  
699 glaciations for millions of years. Icy systems on Earth also may provide crucial terrestrial analogs for  
700 extraterrestrial life surviving and persisting on icy planetary bodies in our solar system, such as Mars,  
701 Europa, or Enceladus. Exploration of life within subglacial lakes and their sediment has begun; the first  
702 reports on the microbiology of Subglacial Lake Whillans have been published (Christner et al, 2014) and a  
703 second subglacial lake (Subglacial Lake Mercer) was accessed in 2019. Of particular interest is the  
704 distribution and ecological function of the resident microbes, the extent to which biogeochemical  
705 weathering occurs, and the genetic diversity of microbial communities in subglacial lakes and sediments.  
706 Furthermore, the forward motion of thick layers of water-saturated till beneath fast-flowing ice streams  
707 may provide a pathway for transportation of subglacial biological and diagenetic materials and weathering  
708 products to the surrounding ocean, as does the movement of debris-rich basal ice. Some subglacial  
709 meltwater is also transported over long distances within basal drainage systems, which again, likely  
710 discharge subglacial nutrients, microbes and their metabolic products into circum-Antarctic seawater.  
711 Access holes through the ice and the acquisition of basal ice cores are needed for this science, and, for  
712 scientific and environmental integrity, these studies must be conducted with clean technology both during  
713 access and sample acquisition. This science is progressing but is at an early stage, and it is best to conduct  
714 studies first at sites where the ice is not thick and logistics issues can be readily addressed.

715  
716 (image here)

717 Caption: Microbial ecosystems have been found under the West Antarctic Ice sheet (Christner et al, 2014). This photo shows a  
718 coccoid-shaped microbe with an attached sediment particle from subglacial Lake Whillans. Photo credit: *Trista Vick-Majors*.

719  
720 **4. Subglacial lakes and hydrological systems:** Subglacial hydrodynamics are an important yet poorly  
721 understood factor in ice sheet dynamics in both Antarctica and Greenland. The volume and distribution  
722 of water exert strong influences on the resistance of the bed to ice flow and therefore, is an important  
723 control over ice velocities. More than 400 subglacial lakes have been discovered in Antarctica.  
724 Measurements to quantify present-day lakes and subglacial hydrological systems are important for  
725 understanding ice dynamics, weathering and erosion of subglacial rock, sediment transport and  
726 jökulhlaup events, microbial ecosystems, and maintaining systems of subglacial lakes. Of particular  
727 interest is to establish the diversity of life in subglacial lakes, the degree of hydrological interconnectivity  
728 between lakes and the Southern Ocean, and their influence on the rest of the subglacial hydrological

729 system. The lakes also house sedimentary evidence of ice sheet and geological histories as well as climate  
730 change. Access holes and the ability to collect samples of water and sediments are necessary to  
731 understand these systems. In addition, data from Subglacial Lake Whillans suggests that in active  
732 hydrological systems, water geochemistry, microbial life, and hydrology are intimately connected (Vick-  
733 Majors et al., 2016), similar to many other terrestrial and marine biospheres around the world.  
734 Understanding the temporal dynamics and geochemical ramifications of subglacial processes requires  
735 installation of sensor strings capable of collecting subglacial hydrological data including dissolved oxygen  
736 concentrations, current velocity and direction, and salinity.  
737

738 Russian drillers accessed Subglacial Lake Vostok during the 2011-12 season, and then during 2012-13  
739 successfully recovered an ice core (~30 m) of the frozen lake water that entered the borehole the year  
740 before. The British attempted to access subglacial Lake Ellsworth in the interior of West Antarctica in  
741 2012-13 but unfortunately were stopped due to operational problems during drilling. The U.S. successfully  
742 penetrated and sampled subglacial Lake Whillans upstream from the Siple Coast grounding line during the  
743 2012-13 season. The new drill built for drilling Lake Whillans includes a filtration unit and UV-treatment  
744 system to decrease contaminants in the drilling water and provide clean access to the subglacial  
745 environment (Priscu et al. 2013) and was used again to access Subglacial Lake Mercer in 2019. The  
746 filtration technology was successful at reducing microbial bioload and other contaminants in the drilling  
747 fluid as per the Antarctic Treaty Code of Conduct.  
748

#### 749 **Summary**

750 Subglacial environments contain biologic, climatic, geologic, and glaciologic materials and information,  
751 much of which cannot be obtained elsewhere. Drills to create access holes are urgently needed to sample  
752 basal ice, subglacial water and sediments, and bedrock. Hole diameter requirements vary depending on  
753 instrumentation needed; clean technology is required (NRC, 2007), as is strict environmental review  
754 where the bed is wet, except for ice shelves and grounding zones at the end of drainage basins. Successful  
755 sampling will require that access holes receive regular maintenance, allowing the holes to remain open  
756 for several days. Differential ice motion may be a complicating factor, especially if the ice sheet is sliding  
757 at the bed. A conceptual design is also needed for a drill that can provide clean access large enough to  
758 deploy subglacial rovers; this design should strive to minimize supporting logistical requirements.  
759

760 The desired characteristics of the drills needed to create clean access holes for the science of the sub-ice  
761 environment are provided in Table 3. For accessing sensitive targets such as subglacial lakes, hot water  
762 drills should have temperature and depth sensors and “smart” drill heads; this technology needs to be  
763 developed.  
764

765  
766

767 **Table 3. Requirements of drills needed for studies of subglacial geology, sediments, and ecosystems.**  
 768 The IDP Long Range Drilling Technology Plan discusses hot water and mechanical rapid-access drills that  
 769 could provide clean access holes for the projects described above. Clean mechanical rapid-access drills  
 770 do not currently exist; conceptual and engineering development is needed. Note that CHWD is the Clean  
 771 Hot Water Drill owned by the University of Nebraska-Lincoln; in addition, the University of Nebraska-  
 772 Lincoln owns the following drills that may be available by request directly to the university (D.  
 773 Harwood): the WISSARD/SALSA clean access drill, the Roving Drill, and Shot-hole Drill.  
 774

	Diam. (cm)	Depth (km)	Core or hole	Ambient temp (°C)	Transport type	Site occupa ncy	Int'l aspects	Environ restrictions	Drill Name
Sediments/i ce sheet dynamics (Wet bed)	10-25	0.2-3	Hole, sediment core	-50	Helo/Twin Otter/travers e/Herc	weeks	U.S. & others	Clean access	CHWD/ SchWD
Biogeochem (Wet bed)	3-25	<3	Hole, sediment/r ock, basal ice core	-50	Helo/Twin Otter/travers e/Herc	weeks	U.S. & others	Clean access	CHWD/ SchWD
Bedrock geology/ Tectonics (Frozen bed)	5-10	1-3	Icehole, rock core	-50	Herc/travers e	4-8 weeks	U.S.	None (dry bed only)	RAID/ ASIG
Geology/ ice sheet history (Wet bed)	5-20	<4k	Hole, rock core	-50	Herc/travers e	weeks	U.S. & others	Clean access	-
Deep Subglacial lake biogeochem (Wet bed)	50-100	3-4k	Hole, sediment, basal ice core	-50	Herc/TwinOt ter /traverse	4-8 weeks	U.S. & others	Clean access	CHWD

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779 **IV. Ice as a Scientific Observatory**

780 Polar ice sheets and mid-latitude ice caps archive evidence of past climate and ice dynamics and also serve  
781 a variety of endeavors that use the ice as a platform for science. Borehole access to the interior of the ice  
782 sheet enables wide-ranging observations, from glaciology, climatology, and planetary science to  
783 experimental astroparticle physics.

784  
785 **1. Borehole logging for past climate and ice dynamics:** Borehole logging of both fast-access holes and  
786 boreholes originally drilled for ice cores greatly enhance evidence of climate and ice dynamics preserved  
787 in the ice. These analyses are difficult or impossible to obtain by other methods, and complement  
788 observations from ice cores and remote sensing platforms. Borehole logging is nondestructive, non-  
789 contaminating, continuous, and immune to core damage or drill depth errors and permits study of a large  
790 volume of ice in situ. Ice sheet boreholes serve as enduring scientific observatories. For example, borehole  
791 paleothermometry probes provide the most direct measurement of temperature histories and can be  
792 used to calibrate other paleoclimatic indicators. Optical borehole probes can rapidly obtain stratigraphic  
793 records, which are more coherent and detailed than can be reconstructed from core measurements.  
794 Borehole sonic loggers can provide continuous records of ice fabric that are difficult or impractical to  
795 obtain using thin sections of core. Repeated measurements of fabric, tilt, and hole deformation improve  
796 modeling of ice sheet behavior and stability over time as an ice sheet flows over uneven terrain. Logging  
797 multiple nearby rapid access holes permits advanced studies of climate history and ice flow.

798  
799 **1.1 Winches:** Winch platforms that can support borehole-logging projects are important community  
800 resources. IDP has three winches in inventory, one for intermediate depth (1.5 km) and two for deep (4  
801 km) applications. IDP has adopted a standard wireline for all community winches, a 3/16" four-conductor  
802 armored oil-patch cable with a 1" Gearhart-Owen cable head. IDP has also established a policy of  
803 deploying a trained operator to the field along with the IDP winches, particularly the deep winches.  
804 Although this cost is not directly reflected in proposal budgets, a cost estimate is included with each  
805 proposal requiring IDP resources, for NSF budgeting purposes, as is the case with ice and rock coring drills.  
806 In certain cases, the PI or members of the PI's team may be trained and certified to operate the winches,  
807 particularly the intermediate depth winch.

808  
809 (image here)  
810 Ryan Bay and Elizabeth Morton deploy borehole-logging instruments at Siple Dome. Photo credit: *Joseph Talghader*.

811  
812 Pre-deployment winch telemetry testing of all logging tools is essential for successful fieldwork. Whenever  
813 possible, logging tools should be tested over the winch that will be used in the field. In some cases IDP  
814 leaves winches deployed to save logistical cost and effort, and tools must instead be tested on winch-  
815 cable systems that are electrically similar.

816 Pressure testing of new borehole tools prior to deployment is performed at an IDP facility in Madison, WI.  
817 IDP maintains a pressure chamber for testing tools up to pressures of 6 kpsi. The chamber is  
818 approximately a 3-meter cylinder with an inside diameter of 25.4 cm. Pressure testing is especially  
819 important with Estisol-140 drill fluid, since it is more aggressive than other drill liquids and even small  
820 leaks may damage internal components.

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**1.2 Borehole preservation:** Where practical, drilling practices and materials should be chosen to produce and maintain clean uniform boreholes, and to keep the boreholes accessible. Anticipated failure modes of glacial boreholes include:

- “Natural” end-of-life borehole collapse: Depending on the strain regime, complete collapse of even a borehole fully compensated with fluid occurs over years and is largely unavoidable.
- Borehole collapse due to removal or failure of borehole casing: Premature collapse can be avoided by leaving the casing in place, proper casing design, and maintenance.
- Borehole burial: Burial of borehole casing by snow accumulation.
- Ice plug: An ice plug can form at the fluid level when a partial casing failure permits snow and ice to accumulate in the well.

Over time borehole-drilling fluids can become turbid, degrading optical measurements. Best practices should include avoiding the introduction of substances such as heavy greases in the borehole, and materials that can be dissolved by solvents used as drill fluids. IDP also provides towers and sheave wheels needed for borehole access. If requested and resourced by NSF, IDP could preserve, maintain, extend borehole casings, and maintain the proper level of drill fluid compensation for existing boreholes. The IDP Englacial and Subglacial Access Working Group will work with community scientists to develop a strategy to make best use of available boreholes.

**1.3 Recent logging projects: WAIS Divide:** Several groups have logged the WAIS Divide ice core borehole in the 2014 to 2017 time frame. Measurements included temperature, optical, and seismic profiling, and an acoustic caliper along with a kHz acoustic fabric logger for the kHz range. WAIS Divide drilling included five replicate coring deviations and this logging activity was the first to be done in a borehole with deviation channels. The Replicate Coring System for the DISC drill was designed in order to make all deviations on the uphill side of the main borehole, so that logging tools naturally follow gravity and remain within the parent channel. The deviations did affect logging data and some issues were encountered while passing the deviations, in particular the acoustic logging tool was diverted into the side channel at the deviation with the most borehole damage near 3,000 m, but accessed the main borehole on a following attempt. Three other logging tools followed the main borehole and passed all deviations without incident. The deviation drilling and subsequent borehole logging at WAIS Divide was largely successful.

**South Pole Ice (SPICE) core:** The Foro 1650 (aka Intermediate-Depth Drill) was deployed to the South Pole for the 2014-15 and 2015-16 field seasons, successfully collecting 1,751 m of ice core. The SPICE Core borehole is pressure compensated by Estisol-140, which has caused convective problems in temperature logging because of its high viscosity. Estisol-140 has also exhibited a tendency to cloud, which could affect optical logging. The SPICE core project is a benefit to ongoing South Pole in-ice particle physics projects, by providing ground truth measurements of ice chemistry, fabric, and particulates for characterization of optical, radio, and acoustic signal propagation. Due to the proximity of the drill site to the IceCube and ARA arrays, the borehole continues to serve as an access point for calibration of existing and future South Pole in-ice physics and astrophysics experiments.

863 **Rapid Access Ice Drill (RAID):** The RAID is a drill in development that will be capable of penetrating a  
864 3,000 m ice sheet, and coring small samples of ice and subglacial basal rock. RAID is expected to produce  
865 five boreholes every season, and these boreholes will potentially serve as scientific observatories for the  
866 study of ice and climate. RAID will require a dedicated logging winch integrated with the drilling platform,  
867 capable of reaching 3,000 m for logging immediately following the drill. The system could partly serve as  
868 a hole qualifier for evaluating the performance of the drill during development. Measurement of pressure  
869 will ensure that the borehole is properly compensated and optical dust logging will provide immediate  
870 verification of the depth-age model. Additional measurements could include temperature, diameter,  
871 borehole inclination/trajectory, and a camera. It would be desirable to rapidly log temperature and  
872 borehole diameter immediately after drilling, possibly at the same time as the optical dating. These  
873 preliminary readings could form baselines for subsequent measurements and time evolution studies.  
874 Infrastructure will be needed to manage future borehole logging projects that will make use of RAID  
875 boreholes.

876  
877 **RAID borehole instrumentation and preservation:** RAID has the potential to create many deep boreholes  
878 over a number of years. Instrumenting and preserving every RAID borehole indefinitely is impractical.  
879 The RAID project, with glacial and subglacial research community, will need to determine the scope of  
880 instrumentation and preservation efforts. The science goals and the number of holes to instrument,  
881 preserve, the priority of holes and the duration of the effort will need to be weighed against cost and  
882 logistics.

883  
884 Borehole preservation effort could be separated into short-term (<5 years) and long-term time horizons.  
885 Preservation of each RAID borehole for 3 - 5 years will allow for repeat measurements, particularly in  
886 studies of borehole temperature and deformation. Uncased and under-balanced boreholes could be of  
887 interest for deformation studies, although removal of the casing and fluid head will limit the lifetime of  
888 the borehole to a few years.

889  
890 RAID may also select a subset of holes for long-term preservation, to serve as observatories and to allow  
891 for future technology developments. Preservation would require leaving a sturdy casing in place,  
892 maintaining, and periodically extending the casing above the snow surface, as well as removal of ice plugs  
893 when necessary. Holes near ice divides could be kept open for decades in principle. In off-axis zones,  
894 shearing could severely limit borehole lifetime and closure may occur at discrete depths. In higher  
895 accumulation areas, it may be possible to use an extended casing supported by a lightweight tower to  
896 relieve maintenance effort. Qualifying tools (borehole diameter, inclination/trajectory, camera) could be  
897 useful for assessing borehole condition prior to fielding a more substantial logging mission. Holes selected  
898 for long-term preservation would likely be chosen to form a geographically diverse set.

899  
900 **1.4 Borehole qualifying:** IDP does not currently maintain logging tools for verifying borehole parameters  
901 such as inclination, diameter, depth, roundness, temperature, etc. There is growing consensus in the  
902 logging community that IDP should develop this capability. A hole qualifying system could be deployed  
903 each season as a hole is drilled or upon hole completion. The information provided by such a logging  
904 system could be crucially important for drillers, particularly for drills with little or no down-hole sensing

905 capacity, such as the Foro 3000, the Foro 1650, or the RAID. These logging measurements could also  
906 provide a baseline for longer-term borehole deformation studies.

907

908 **1.5 Borehole Allocation Committee:** The IDP Englacial and Subglacial Access Working Group (ESAWG)  
909 will be exploring formation of a special committee to advise IDP on management of community borehole  
910 resources as the research community continues to grow. These resources include winch and winch  
911 operators, logging tools and accessories, and borehole time. Pre-deployment reviews of logging projects,  
912 with participation by IDP engineers, will ensure that new tools are safe and ready to deploy.

913

914 **2. Ice as platform for physics and astrophysics:** Efforts are under way to use glacial ice as a platform for  
915 study of fundamental physics and astrophysics. These experiments make use of polar ice as an abundant,  
916 clean, stable, low-background and transparent (to radio and optical waves) detection medium for  
917 observation of sub-atomic particle interactions. For example, the now completed IceCube telescope uses  
918 ice at South Pole to detect high-energy neutrinos traveling to Earth from cosmic sources. The Enhanced  
919 Hot Water Drill (EHWD) developed for IceCube is a powerful and fast access drill capable of creating 2,500  
920 m deep, half-meter diameter boreholes at a rate of about three per week.

921

922 IceCube-Gen2 is a proposed facility for future Antarctic neutrino astronomy. IceCube-Gen2 will aim to  
923 increase the effective volume of IceCube by an order of magnitude, while only doubling the amount of in-  
924 ice instrumentation. The IceCube inter-string spacing of 125 m would be increased to 250 - 300 m, taking  
925 advantage of the long absorption lengths of optical photons in Antarctic ice, particularly South Pole ice  
926 from the early stages of the Last Glacial Period. This expanded array would improve detection capability  
927 in the PeV energy range and provide high statistics samples of extraterrestrial neutrinos, for better  
928 characterization of source distribution, spectrum and flavor composition. IceCube-Gen2 will require  
929 improvements to the EHWD, including a more mobile and efficient hot water plant, and a modular sled-  
930 mounted drill system, which is less complex and requires a smaller operations crew.

931

932 (image here)

933 A Digital Optical Module (DOM) is lowered into a hole in the ice at Amundsen-Scott South Pole Station for the IceCube project. IceCube detects  
934 neutrinos from distant astrophysical sources. Photo credit: *Ethan Dicks, National Science Foundation.*

935

936 The proposed low-energy sub-array physics experiments such as PINGU (Precision IceCube Next  
937 Generation Upgrade) would be embedded within the IceCube array, in order to use the existing detector  
938 as an active shield. PINGU objectives include the study of neutrino oscillations and mass hierarchy, dark  
939 matter, supernovae, and neutrino tomography of Earth's core. PINGU will deploy a relatively high density  
940 of photocathode (light sensors) in a small ice volume, requiring hot-water drills capable of making deep  
941 access holes at small spacing. The currently proposed PINGU geometry will have inter-string spacing of  
942 ~20 m and a three to five meters vertical spacing between sensors. These projects will enable research  
943 and development on the next generation of low-light photodetectors and the optical properties of in situ  
944 ice over short distance scales. Hot-water drill upgrades are aimed at improving the optical clarity of the  
945 refrozen water column, including filtration of large-particle impurities and degassing to avoid bubble  
946 formation.

947  
948 Experiments to detect extremely high-energy neutrinos will make use of large areas of the polar ice sheet.  
949 The ARA experiment (Askaryan Radio Array), in early development at South Pole, is planning to instrument  
950 on the order of 100 km<sup>2</sup> of ice with radio antennas to detect radio pulses from so-called Greisen-Zatsepin-  
951 Kuzmin-scale (GZK) neutrinos. The Radio Neutrino Observatory (RNO) at the South Pole is a proposed,  
952 distributed, detector array for measuring the highest energy neutrinos. Detector stations are spaced 1.25  
953 km apart in rectangular grid spaced away from the South Pole station along a power and communications  
954 backbone. Drilling would consist of 4-5 holes per station, drilled down to a nominal depth of 60 m, with  
955 the possibility of 100 m evaluation holes. Drilling would be conducted with the ASIG auger drill setup on  
956 a sled along with integrated weather sheltering. Holes are used for radio antennas to detect the radio  
957 pulses from high energy neutrinos. Field work over four Austral Summer seasons has been proposed.  
958

959 Due to its proximity to the IceCube and ARA detectors, the SPICE core borehole could serve as an access  
960 point for calibration beacons or standard candles, as part of the South Pole facility and infrastructure.  
961 These beacons could be operated at multiple depths and hence different ice temperatures, densities,  
962 fabrics and impurity levels. These unique measurements would have implications for radio and optical  
963 detection of high-energy neutrinos and also provide opportunities for basic glaciology research. Radio-  
964 illuminating beacons could provide signals in the 100 - 1,000 MHz frequency range out to a radius of 20  
965 km, thus permitting studies of neutrino detection over areas up to 1,000 km<sup>2</sup>, and also help in  
966 understanding anomalous features seen in ice-penetrating radar surveys.  
967

968 The ARIANNA experiment (Antarctic Ross Ice shelf Antenna Neutrino Array) proposes to deploy a large  
969 array of surface radio antennas on the Ross Ice Shelf to observe cosmogenic GZK neutrinos. Log-periodic  
970 dipole antennas will be buried in pits approximately 6' x 6' x 0.5' and controlled by solar-powered relay  
971 stations. An effective and efficient means for digging and backfilling many such pits will need to be  
972 developed. The ARIANNA also anticipates taking two shallow (~100 m), 4"-6" cores from nearby for study  
973 of the firn-to-ice transition and for borehole-to-borehole radio tomography, as well as drilling one deep  
974 (~500 m), 4"-6" borehole.  
975

976  
977 **3. Seismic studies:** The Global Seismographic Network includes seismic monitoring stations for  
978 earthquakes and other events such as emissions from calving and sliding glaciers and ice sheets. The South  
979 Pole Remote Earth Science and Seismological Observatory has seismic equipment installed ~300 m deep  
980 within boreholes. A similar observation network is planned for Greenland.  
981

982 **4. Ice sheet as an archive of recent past atmospheric composition:** In the very cold areas of ice sheets  
983 where snow rarely melts, many decades of snowfall create a porous network of firn in the top many tens  
984 of meters of the ice sheet. The firn serves as an archive of atmospheric composition, with the oldest air  
985 existing at depth. Sampling firn air from various depths within boreholes drilled in the ice sheet enables,  
986 for example, observation of the extent of anthropogenic emissions and patterns of increase or decrease.  
987

988 **5. Exploration of basal ice formation processes:** Radar imaging of basal conditions under the Antarctic  
989 and Greenland ice sheets reveals structures that have been proposed to result from accretion ice grown  
990 onto the base of the ice sheet. In order to acquire the ice to test this hypothesis, drilling at sites in  
991 Greenland, or near Dome A in East Antarctica, could access these ice features with the 1,500 m  
992 Intermediate Depth Drill.

993  
994 **6. Meteorite collection:** Glaciers and ice sheets are sites for efficient collection of meteorites and  
995 micrometeorites. Micrometeorites yield clues to the origin and evolution of the solar system. Some are  
996 visible to the human eye on the surface of some blue ice areas, while others may be swept up inside  
997 melted water wells created in the ice at established field stations.

998

999 **Summary**

1000 Ice sheets serve as a platform for a wide range of observations spanning many areas of science. In some  
1001 areas, for example, firn-air studies and seismic monitoring proven-drills already exist for making the  
1002 necessary access holes. Dedicated hot water drills have proven to be effective in creating deep boreholes  
1003 in rapid succession. Other areas are at an early stage and will require further development of RAM drills  
1004 or reverse circulation drills. A rapid access drill, with the capability to bore through several kilometers of  
1005 ice to retrieve rock cores is in development. The borehole logging community is a strong proponent for  
1006 repairing and maintaining boreholes at Greenland Ice Sheet Project 2 (GISP2) site at Summit, Siple Dome,  
1007 and other sites. Identifying which boreholes need maintenance, prioritizing those with highest scientific  
1008 value for future logging, and determining methods of repair are activities that need urgent attention. The  
1009 IDP Borehole Logging Working Group will prepare a list of boreholes in the U.S. program and will work  
1010 with the community to create a prioritized list for maintenance and repair.

1011

1012 **Science Planning Matrices**

1013  
1014 Goals to advance the frontiers of the science in ways that enable evidence-based decision making and  
1015 that inspire the next generation of scientists are described in the sections above. Community planning  
1016 for the execution of the science is important for providing coordinated scientific investigations, and also  
1017 for planning the associated logistical and funding requirements. For each area described above, Tables 4  
1018 through 7 identify the current plans for timing of the field research. In cases where new technologies are  
1019 needed, a timeline for the development of technologies is provided. Black lettering in a matrix indicates  
1020 projects that are currently funded, and blue lettering indicates those in the planning phase.

1021  
1022 In Tables 4 – 7 the letters denoting specific drills to be used are: A: Agile sub-ice geological drill; b: Badger-  
1023 eclipse; B: Blue ice drill; D: DISC drill; f: Foro 400 drill; F: Foro 3000 drill; I: Intermediate depth drill  
1024 (Foro1850); L: Borehole logging; Lt: Logging tower; R: RAID drill; r RAM drill; Sc: Clean Scalable hot water  
1025 drill; S: Staphli drill; T: Thermomechanical drill; U: UNL CHWD drill; W: Winkie drill; 4: 4" drill; 7: 700 m  
1026 coring drill, x: hand auger, sidewinder, or prairie dog.

1027

1028

1029 **Table 4: Past Climate Change Planning Matrix 2023-2033**

	2023				2024				2025				2026				2027				2028				2029				2030				2031				2032				2033							
	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4				
<b>Past Climate</b>																																																
<b>Industrial period and glaciology</b>																																																
Mt Waddington Canada firn aquifer <sup>4</sup>	T	T																																														
AON Greenland firn <sup>1</sup>	x	x			x	x			x	x			x	x																																		
Biogenic sulfur Greenland <sup>2</sup>	x	x																																														
Biogenic CO <sub>2</sub> , CH <sub>4</sub> Greenland <sup>3</sup>	B	B																																														
Greenland outflow chem bio <sup>5</sup>	x	x																																														
Taylor Dome firnification <sup>6</sup>			b	b																																												
Taylor Dome H <sub>2</sub> <sup>7</sup>							7	7																																								
<b>Pre-industrial baseline &amp; dynamics</b>																																																
Eclipse Icefield Canada <sup>8</sup>									s	s			7	7																																		
Andes <sup>9</sup>	T																																															
Guest Peninsula <sup>4</sup>					f	f																																										
Dome C - past cosmic ray flux <sup>10</sup>					4	4																																										
Detroit Plateau Ant. Peninsula <sup>11</sup>															7	7			7	7																												
<b>Large scale global climate change</b>																																																
<i>GISP2.1 Central Greenland (near Summit)</i> <sup>12</sup>																																																
Intermediate depth coring to 1650 m									I	I			I	I																																		
Borehole logging													L				L																															
<i>NW Greenland</i> <sup>13</sup>																																																
Prudhoe Dome 700 m core									7	7							L																															
Borehole logging Prudhoe Dome																																																
<i>South Dome Greenland</i> <sup>14</sup>																																																
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Borehole logging South Dome																																																
<i>Hercules Dome</i> <sup>15</sup>																																																
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Replicate coring at Herc Dome																					R																											
Borehole logging at Herc Dome																							L																									
<i>2 Ma ice &amp; last interglacial</i>																																																
COLDEX Intermed core Allan Hills <sup>16</sup>									I	I			I	I																																		
COLDEX shallow cores (BID & 4") <sup>16</sup>			B	B			B	B			B	B																																				
<i>COLDEX deep core</i>																																																
East Antarctica tbd COLDEX <sup>14</sup>																							F	F			F	F			F																	
Replicate coring for deep core COLDEX <sup>14</sup>																																											R					

1030 Point of Contact for projects in Table 4: <sup>1</sup>Harper; <sup>2</sup>Alexander, <sup>3</sup>Chellman; <sup>4</sup>Neff; <sup>5</sup>Licht; <sup>6</sup>Keegan; <sup>7</sup>Saltzman;  
 1031 <sup>8</sup>Kreutz; <sup>9</sup>Mayewski; <sup>10</sup>Petrenko; <sup>11</sup>Kurbatov; <sup>12</sup>Winski; <sup>13</sup>Osterberg; <sup>14</sup>Severinghaus; <sup>15</sup>Steig; <sup>17</sup>Brook.  
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1035 **Table 5: Ice Dynamics and Glacial History Planning Matrix 2023-2033**

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	2023				2024				2025				2026				2027				2028				2029				2030				2031				2032				2033																							
	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4																				
<b>Ice Dynamics &amp; Glacial History</b>																																																																
<b>Ice Dynamics</b>																																																																
Flask glacier melt dynamics <sup>1</sup>									x x																																																							
Borehole logging of RAID holes <sup>2</sup>													L L				L L				L L																																											
<b>Glacial history</b>																																																																
GreenDrill <sup>3</sup>	A A				A A																																																											
WAIS interglacial <sup>4</sup>					w w																																																											
Thwaites - seismic sounding <sup>5</sup>					x x				x x																																																							
Mt Waesche unconformities <sup>6</sup>					b b																																																											
Taylor Glacier Antarctica <sup>7</sup>													Sc Sc																																																			
Northern Victoria Land <sup>8</sup>													w w				A A				R R				R R				R R																																			
Continental RAID drilling Antarctica <sup>9</sup>																																																																

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Point of Contact for projects in Table 5: <sup>1</sup>Kingslake; <sup>2</sup>Pettit; <sup>3</sup>Schaeffer; <sup>4</sup>Mitrovika; <sup>5</sup>Anandakrsihnan; <sup>6</sup>Campbell; <sup>7</sup>Mikucki; <sup>8</sup>Balco; <sup>9</sup>Goodge.

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**Table 6: Subglacial Geology, Sediments and Ecosystems Planning Matrix 2023-2033**

	2023				2024				2025				2026				2027				2028				2029				2030				2031				2032				2033																			
	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4																
<b>Subglacial Geology, Sediments, &amp; Ecosystems</b>																																																												
<b>Bedrock geology</b>																																																												
Continental RAID drilling Antarctica <sup>1</sup>													R R				R R				R R																																							
WAIS interglacial <sup>2</sup>					w w																																																							
Thwaites - seismic sounding <sup>3</sup>					x x				x x																																																			
GreenDrill <sup>4</sup>	A A				A A																																																							
<b>Subglacial hydrology &amp; sediment dynamics</b>																																																												
Taylor Glacier Antarctica <sup>5</sup>													Sc Sc																																															
Mount Resnick Antarctica <sup>6</sup>													A																																															
West Antarctica / Siple Coast <sup>7</sup>																	Sc Sc																																											
Subglacial sediment experiment <sup>8</sup>													c c																																															
<b>Microbial ecosystems &amp; biogeochem</b>																																																												
Taylor Glacier Antarctica <sup>5</sup>													Sc Sc																																															
Mount Resnick Antarctica <sup>6</sup>													A																																															
West Antarctica / Siple Coast <sup>7</sup>																	U U																																											
Greenland subglacial chem-bio <sup>9</sup>					x				x																																																			

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Point of Contact for projects in Table 6: <sup>1</sup>Goodge; <sup>2</sup>Mitrovika; <sup>3</sup>Anandakrishnan; <sup>4</sup>Schaefer; <sup>5</sup>Mikucki; <sup>6</sup>Mikucki, <sup>7</sup>Vick-Majors, <sup>8</sup>Balco; <sup>9</sup>Licht



1052 **Table 7: Ice as a Scientific Observatory Planning Matrix 2023 – 2033**

1053 Note: These projects either already have holes in place, or else are in need of holes but the drill has not  
1054 yet been identified.  
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	2023				2024				2025				2026				2027				2028				2029				2030				2031				2032				2033			
	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4
<b>Ice as a Scientific Observatory</b>																																												
<b>Neutrino detection &amp; seismic network IceCube Update (Karle)</b>	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x				
<b>South Pole Global Seismic Network</b>	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x				

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1058 **Associated logistical challenges**

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1060 In addition to planning the science and associated drilling technology, logistical challenges impact the  
1061 timing and possibilities of the field science. Challenges to conducting the field activities include:

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- There is an ongoing need for reliable tractors and sleds for scientific traverses in both Greenland and Antarctica in order to address the priority science goals identified in science community consensus documents, including the IDP Long Range Science Plan 2022-2032 (this document). In addition, the 2015 National Academies report “A Strategic Vision for NSF Investments in Antarctic and Southern Ocean Research” articulated the need for improved ground-based access to the deep field in Antarctica. Access to sites on the Antarctic and Greenland Ice Sheets has become very restricted due to limited air support and aging scientific traverse vehicles. With multiple science communities requesting flights or traverses to support their science on the ice sheet, access to the deep field has been a limiting factor in executing new scientific endeavors.
- The NSF Ice Core Facility in Denver is the key location for processing and archiving of U.S. ice cores. Although some infrastructure upgrades and improvements have been made, it is an aging facility that will soon reach full capacity. Expanding the ice core storage facility requires a major investment in infrastructure, which is currently in process. The ice core science community should be involved in all stages of the design process. Fudge and others (2020). The archive of ice in the NSF-ICF and the process for requesting samples for science should be advertised through a variety of venues to early career and other researchers to enable broadening of the community and increase potential collaboration, and to provide an entry point for researchers new to the ice core community.
- The community wishes to instrument and maintain key boreholes as long-term observatories for conducting measurements with existing and new instruments. GISP2 at Greenland summit is one of the most influential and widely cited records in paleoclimatology, but measurements have shown that the borehole casing is collapsing and already not navigable by most logging instruments. The IDP Englacial and Subglacial Working Group will lead community discussions and planning for an approach and a method for prioritizing boreholes for instrumentation and preservation.

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- Drilling ice cores deeper than ~300 m generally requires a drilling fluid mixture that has a density similar to ice to maintain core quality and prevent borehole closure. The fluid must also have a viscosity that is low enough to permit passage of the drill sonde through the fluid many times during the drilling process. Estisol-140 is a fluid that was identified by international partners; it was used at South Pole for the SPICE core project, and it will likely be used for future drilling projects until an improved fluid is identified.

1095 **Broadening participation**

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1097 It is well recognized that increasing diversity within the polar science community has been challenging but

1098 would broaden perspectives and enhance the science. IDP also recognizes this challenge. We strive to

1099 reflect community diversity in our staff, working groups and on the Science Advisory Board. Our Education

1100 and Outreach efforts, for example the IDP School of Ice, are designed to primarily engage

1101 underrepresented groups. More is needed to broaden participation, and we will continue to pursue varied

1102 approaches to increasing diversity. In the near term, one approach may be to encourage use of the ice

1103 cores archived at the NSF Ice Core Facility (NSF-ICF) for different areas of science than are currently

1104 represented in ice core analysis; this could be pursued if knowledge of existence of the NSF-ICF and its ice

1105 were better-known across other areas of science. Broader awareness of NSF-ICF may result in increased

1106 scientific involvement by scientists who do not work in the Polar Regions, and it may help to bring new

1107 people and perspectives into the community. In the coming year, the IDP Science Advisory Board and

1108 working groups will consider how our collective actions could help to broaden awareness about the NSF-

1109 ICF, encourage use of archived ice for a wider array of scientific investigations, and encourage

1110 collaboration with students and scientists from underrepresented groups.

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1112 **Recommendations**

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1114 **Recommended science goals**

1115 **1. Past Climate change:** Present-day climate change can only be fully understood in context of the past;

1116 well-dated histories of climate and the atmosphere over a wide range of time scales are needed to

1117 understand climate forcing and response. U.S. scientists are leading investigations in national teams and

1118 also in international circles, including generation of international science goals through the International

1119 Partnerships in Ice Core Sciences (IPICS). The U.S. ice coring community has always been intimately

1120 involved in establishing the IPICS goals. Some of the goals below include U.S. involvement in IPICS targets,

1121 while some of the goals below are primarily from members of the U.S. community who, for example, are

1122 leading efforts to gain critical samples of ice prior to 800,000 years ago, for evidence of the atmosphere

1123 from times when the Earth had 40,000-year climate cycles.

- 1124 ● Drilling of spatially-distributed ice cores and boreholes at many locations to investigate past climate
- 1125 and atmosphere over the past 200 to 40,000 years should continue. Understanding climate signals in
- 1126 remotely-sensed data, understanding climate impacts on the transition from snow to firn to ice on ice
- 1127 sheets, and calibrating high-resolution models, all require arrays of shallow cores covering a range of
- 1128 accumulation and melt rates both in Greenland and in Antarctica; these efforts should continue.
- 1129 Spatially-distributed shallow coring for records ranging from the recent past to 2,000 years will include
- 1130 multiple scientific traverses in Greenland for study of the ice sheet under the currently changing
- 1131 climate. Recent projects in the Arctic include the 1,000-2,000 year annual record from Denali (Mt.
- 1132 Hunter), Alaska, which is providing important constraints on North Pacific climate and tropical
- 1133 teleconnections, during the Medieval Climate anomaly, Little Ice Age, and modern warming. In

1134 Antarctica, one example is the recently-completed shallow ice coring at Law Dome, aimed at  
1135 reconstructing changes in atmospheric oxidizing capacity. Proposed Antarctic science includes ice  
1136 coring at Dome C to investigate Holocene changes in the cosmic ray flux.

1137 ●Determining patterns of hydroclimate variability, climate feedbacks, and past extent of high-altitude  
1138 glaciers and aerosol deposition requires ice coring in the Sub-Antarctic Islands, North Pacific coastal  
1139 mountain ranges, and the Karakoram in Asia.

1140 ●Determining the amount of meltwater retained and refrozen in the near surface firn (top ~60 m) on  
1141 the Greenland Ice Sheet and on the Antarctic Peninsula is critical for improving estimates of surface  
1142 mass balance under current warming conditions.

1143 ●Targeted ice coring to investigate ice, ocean, and atmospheric dynamics in WAIS coastal domes and  
1144 coastal ice caps and along the dynamic Amundsen Sea Coast of Antarctica, and near Camp Century  
1145 along the northwest coast of Greenland, are in the planning stages.

1146 ●A climate record from the last interglacial period (the Eemian, ~130k to 110k years ago) is key to  
1147 predicting the response of glaciers and ice sheets to future warming. The search for sites from which  
1148 to extract Eemian ice in Greenland, both by coring and through horizontal sampling of blue ice ablation  
1149 zones, should continue. Eemian ice was recovered from the Camp Century core in the 1960's, and an  
1150 effort to retrieve an intermediate depth ice core from this region is in the planning stages. In Antarctica,  
1151 extracting a record from Eemian ice is especially important for helping constrain climate and glacial  
1152 histories of the WAIS during the last interglacial, and is the primary motivation for planned deep drilling  
1153 at Hercules Dome. An ice core from Hercules Dome would lead to understanding whether the WAIS  
1154 collapsed during the last interglacial period (MIS5e), and if it did not collapse, then under what climate  
1155 conditions was it stable? Hercules Dome is the highest-priority next deep ice core for the US  
1156 community. WAIS history during the Eemian is poorly known; because large sea level rise due to  
1157 current climate warming may occur if the WAIS becomes destabilized, an understanding of the WAIS  
1158 during the last interglacial is urgent.

1159 ●Blue-ice paleoclimate records are already providing unlimited samples for atmospheric and ultra-trace  
1160 component studies and can enable further new types of measurements that have previously been  
1161 impossible, including analysis of ice older than 800,000 years. Blue-ice studies at Mt. Moulton, Taylor  
1162 Glacier, and Allan Hills exemplify discoveries from this realm so far; such studies at blue ice sites should  
1163 continue.

1164 ●Ice cores and borehole observations reaching ages between 800,000 years and 1.5 M years (or beyond)  
1165 are significant, for these data may provide new insight into the effects of greenhouse gases on climate  
1166 and the observed change in periodicity of glacial cycles during the mid-Pleistocene. The search to  
1167 identify sites suitable for extracting ancient ice should continue. Extraction of deep ice cores for  
1168 million-year-old ice is currently a goal of the COLDEX program. As part of COLDEX, U.S. scientists will  
1169 continue to need samples of ancient ice from blue ice regions that provide snapshots of climate as it  
1170 existed more than a million years ago.

1171  
1172 **2. Ice dynamics and glacial history:** Rapid changes in the speed of fast-flowing outlet glaciers and ice  
1173 streams observed over the past decade create an urgency to understand the dynamics of outlet glaciers  
1174 and ice sheets. Ice-sheet models that incorporate realistic physics and dynamics at appropriate spatial and  
1175 temporal scales are needed to predict the "tipping point" when ice-loss becomes irreversible, resulting in  
1176 ice-sheet collapse and rapid sea-level rise. Observational data are needed to develop and validate the  
1177 models. Measurements of the ice-bed interface (frozen-thawed, hard-soft bed conditions, sliding, shear),  
1178 ice-ocean interactions (sub-shelf and basal melting-freezing rates), temperatures and ice deformation  
1179 properties through the ice, geothermal bedrock conditions and ice-atmosphere interactions (surface mass  
1180 balance) are key. Another approach to understanding future possible response of ice sheets is to examine  
1181 their behavior in the past. Dated marine and terrestrial glacial deposits provide information about past

1182 ice volume. In regions where such data are not available, histories of ice-sheet thickness and climate can  
1183 be inferred from radar-detected layers combined with ice core and borehole measurements.

1184 Specific recommendations include:

- 1185 ● Ice-ocean interactions are not fully understood. Boreholes to deploy instruments to measure  
1186 conditions at ice-ocean interfaces are high priority for investigating ice sheet stability.
- 1187 ● Hydraulic conditions in glaciers and ice sheets exert strong controls on basal motion. Much has been  
1188 learned through remote sensing methods, but direct measurements through boreholes to the bed are  
1189 still needed to validate and interpret remote sensing data. Boreholes to the bed at targeted locations  
1190 are urgently needed to measure geothermal fluxes and basal properties.
- 1191 ● Ice deformation in ice sheets, glaciers, and ice streams depend on temperature and ice rheology.  
1192 Measurements of ice rheology from ice cores, and borehole logging measurements of temperature,  
1193 diameter, inclination, and azimuth are needed to provide boundary conditions and constraints for  
1194 modeling flow of ice sheets and fast-flowing outlet glaciers and ice streams.
- 1195 ● Knowledge of spatial and temporal variations of surface accumulation is critical for quantifying the  
1196 mass balance of glaciers and ice sheets. Accumulation rate histories derived from short (~200 m) firn  
1197 and ice cores can be extrapolated spatially to the catchment scale using radar-detected layers.  
1198 Additional short cores at targeted locations are needed to provide a realistic assessment of surface  
1199 accumulation over ice-sheet scales.
- 1200 ● Dated ice cores can be used to infer histories of thickness and configuration of ice sheets. Glacial  
1201 histories contained in coastal ice domes are of particular interest because thickness change near the  
1202 margins is large. The depth-age relationship from Siple Dome provided key information about the  
1203 Holocene deglaciation of the central Ross Embayment, and the depth-age relationship from Roosevelt  
1204 Island will help constrain the deglaciation of the eastern Ross Embayment. Depth-age profiles from  
1205 Hercules Dome and other targeted locations are essential for understanding the timing and extent of  
1206 deglaciation, for example at ice domes near the outflow of the Amundsen Sea Embayment Antarctica,  
1207 as well as in coastal domes of Greenland.
- 1208 ● The past extent and volume of the Greenland and West Antarctic Ice Sheets is recorded by cosmogenic  
1209 nuclides in subglacial bedrock. Samples from beneath these ice sheets will provide information on  
1210 their thickness and configuration during paleoclimates warmer than the present, and help identify  
1211 their sensitivity to future possible climate change. Short cores of basal ice and bedrock from targeted  
1212 sites are needed to address questions concerning the extent of the ice sheets during past interglacial  
1213 climates, and the onset of continental glaciations.

1214

1215 **3. Subglacial geology, sediments, and ecosystems:** Bedrock, sediments, and ecosystems existing within  
1216 and beneath ice sheets remain largely unexplored because of the lack of rapid access drills. In particular,  
1217 the physical conditions at the base of the ice sheets are virtually unknown, but remote sensing of liquid  
1218 water in subglacial lakes and possibly interconnected hydrologic systems raises concern about thermal  
1219 conditions and basal slip potential. Likewise, the unknown subglacial geology of Antarctica represents the  
1220 last continental frontier of geologic exploration, including landscape evolution, past paleoclimates on  
1221 geological timescales, crustal heat flow, lithospheric stress, ground truth for geophysical imaging,  
1222 constraints on geodynamical evolution, and relationship with past supercontinents. Information on  
1223 subglacial biodiversity and biogeochemistry is limited to the Siple Coast, with nothing known about other  
1224 areas of the Antarctic continent. Subglacial sediments also contain information related to past ice sheet  
1225 history. .Rapid access to subglacial environments is needed to address a wide range of science questions.

1226 Specifically,

- 1227 ● Direct sampling of the bedrock is needed to validate models of cratonic growth related to  
1228 supercontinent assembly in the Mesoproterozoic between about 2.0 and 1.1 billion years ago and for  
1229 constraining the Phanerozoic geological, tectonic, and exhumation history of the Antarctic continent.

1230 Strategic drill-site selection within mapped drainage basins (using products from the BEDMAP2  
1231 project) will also allow greater constraints on provenance studies that utilize onshore moraines and  
1232 offshore glacial strata.

- 1233 ● There exist virtually no heat flow data for Antarctica. Penetration into bedrock provides the first  
1234 opportunity to accurately measure the geothermal heat flux, which informs us about geotectonic  
1235 conditions as well as geothermal contributions to ice-sheet temperature.
- 1236 ● Evidence of Cenozoic ice sheet history preserved in sedimentary rocks of subglacial bedrock basins  
1237 and in sediment deposits within subglacial lakes will provide further dimensions to the records known  
1238 only from the margins of the continent and will also help to verify paleo-topographic reconstructions  
1239 for ice sheet modeling. Likewise, access to subglacial bedrock can provide a unique opportunity to  
1240 study Cenozoic landscape evolution and long-term ice sheet stability using low-temperature  
1241 thermochronology and cosmogenic-isotope techniques.
- 1242 ● Direct measurements at grounding zones of fast-flowing ice streams and outlet glaciers are badly  
1243 needed, as are data from sub-ice-shelf ocean cavities in order to provide basic information needed to  
1244 model ice fluxes near grounding lines and into ice shelves – a critical interface for predicting future  
1245 ice sheet dynamics.
- 1246 ● Direct measurements of bed conditions including frozen/thawed bed, basal pore pressure, slip, and  
1247 sediments are needed to develop and test realistic models of the controls on the fast flow of ice  
1248 streams and outlet glaciers.
- 1249 ● Significant wet environments exist below ice sheets and glaciers; sampling of subglacial habitats  
1250 including sediments, water, and basal ice is needed to establish the diversity and physiology of  
1251 microbes, available nutrients and organic materials, microbial relationships to past climates, and  
1252 ecosystem function below the ice. Continued support for developing methods and technologies for  
1253 clean access to subglacial environments and tools for biological and geochemical sampling are needed  
1254 to investigate these subglacial systems in a clean manner that maintains scientific integrity and  
1255 environmental stewardship. The recent studies of Whillans and Mercer Subglacial Lakes are steps  
1256 toward achieving this goal.

1257  
1258 **4. Ice as a scientific observatory:** Polar ice sheets and mid-latitude ice caps archive evidence of past  
1259 climate and ice dynamics and also serve as a unique platform to conduct observations and experiments  
1260 concerning seismic activity, planetary sciences and experimental astrophysics, and other novel  
1261 phenomena. Specifically,

- 1262 ● Borehole logging of both fast-access holes and boreholes originally drilled for ice cores are needed to  
1263 fully exploit the histories of climate and ice dynamics preserved within the ice. For example,  
1264 temperature logs are used to infer past temperatures and also the geothermal flux; optical logs yield  
1265 detailed records of dust and volcanic events and will be important in searches for million year old ice;  
1266 and sonic logs provide a continuous record of ice fabric and borehole deformation. Community  
1267 winches to support borehole logging are important assets.
- 1268 ● In-ice physics and astrophysics experiments make use of polar ice as a clean, highly stable, low-  
1269 background, and transparent (both optically and in the radio frequencies) detection medium for  
1270 observation of sub-atomic particle interactions. New drilling techniques are under investigation,  
1271 including cleaner drilling and removal of bubbles from the refrozen water.
- 1272 ● Future planned projects (e.g., the Askaryan Radio Array and Generation-2 Ice Cube) require multiple  
1273 boreholes drilled to at least 150 m deep (ARA) and 2,500 m deep (G-2IC) and significant calibration  
1274 studies of the surrounding ice volume. Better understanding of ice attenuation at radio and deep UV  
1275 wavelengths are particularly desired.

- 1276 ●Ice sheets are a quiet platform for seismic monitoring; the South Pole Remote Earth Science and  
1277 Seismological Observatory has seismic equipment installed in boreholes about 300 m below the  
1278 surface. A similar seismic observation network is planned for the Greenland Ice Sheet.  
1279 ●Novel basal ice structures that have been remotely sensed but whose existence is not well understood  
1280 should be investigated.

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### **Recommended drill life cycle cost and logistical principles**

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1284 Although drills already exist that can achieve some science goals, new drilling technologies are needed to  
1285 accomplish science goals planned for the next decade. The following principles guiding development of  
1286 new drills and associated technologies are recommended:  
1287

- 1288 ●Designs require that the supporting logistical needs do not impede execution of the science.  
1289 ●While developing the science requirements, logistical issues such as weight, size, costs, and time for  
1290 development must be defined and transparent at the initial stages. Scientists and engineers working  
1291 together through IDP must assess the impact of changes as they arise during the development process.  
1292 ●Drills, major drilling subsystems, and accompanying technology must be developed with consideration  
1293 of potential use in future projects. The drills and technology must be versatile and well documented  
1294 so that they can be used, maintained, and repaired by other engineers.  
1295 ●Major drilling systems (e.g., sondes, winches, control and other major electronics systems) should be  
1296 fungible to the maximum extent possible. Major component interchangeability and logistical agility  
1297 should be essential deliverables for all new drilling technology projects.  
1298 ●Engineering design teams must include individuals with field experience using appropriate ice drilling  
1299 technology and/or other relevant field experience.  
1300 ●Increased medium and heavy scientific traversing infrastructure are urgently needed to improve access  
1301 to many scientifically important regions of the Antarctic and Greenland Ice Sheets.  
1302

1303 **Recommended Technology Investments**

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1305 The following investments in drilling technologies are needed to accomplish science goals planned for  
1306 the next decade. Investments prioritized by time (but not prioritized within each Priority level) from  
1307 consensus of the IDP Science Advisory Board, include:

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1309 **Priority 1 (needed this year):**

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1311 ● Maintain and upgrade agile equipment in inventory, including: Hand Augers, Sidewinders, the Foro  
1312 400 drill, the 4" Electromechanical Drills, the 3" Electrothermal Drill, the 3.25" Badger-Eclipse Drills,  
1313 the Stampfli Drill, Logging Winches, the Small Hot Water Drills, the Blue Ice Drill, the Prairie Dog, the  
1314 Agile Sub-Ice Geological Drill (ASIG), the Rapid Air Movement Drill (RAM) Drill, and the Winkie Drill.

1315 ● Adapt the BASE drill rig for retrieving rock core from beneath 200 m of ice.

1316 ● Develop the Conceptual Design for collecting a small amount (chips to several cm) of sub-ice  
1317 rock/mixed media/mud in a frozen regime using an intermediate or deep ice core drill in a fluid filled  
1318 hole, for example with the Foro 3000 drill.

1319 ● Continue construction of the 700 Drill.

1320 ● Develop the updated IDP Conceptual Design and Detailed Design for a clean Scalable Hot Water Drill  
1321 that minimizes its logistical footprint including fuel supply.

1322 ● Investigate a lighter weight source of power to replace generators for drilling systems, in order to  
1323 ease demand on logistics, including renewable energy.

1324 ● Finish the Conceptual Design and begin the Detailed Design for replicate coring for the Foro 3000  
1325 drill.

1326 ● Develop the Detailed Design for clean hot water basal ice coring mechanism for a hot water drill

1327 ● Conduct engineering feasibility study to evaluate and recommend longer-term drilling approaches to  
1328 retrieve ice with good core quality down to 400 m depth in blue ice areas. Possible approaches  
1329 include: replicate coring for Foro 400, large-diameter thermal drill (diameter < 241 mm TBD),  
1330 complete re-design of the BID, and other. For each approach, estimate the anticipated improvement  
1331 in core quality, impacts on associated logistics, and the time required to complete the resulting drill.  
1332 Identify the most promising approach that could be implemented not later than 2028, but earlier if  
1333 possible.

1334 ● Resolve deep logging winch electrical noise issues.

1335

1336 **Priority 2 (needed in the next 3 years)**

1337 ● Build a Scalable Hot Water Access drill for creating access holes in ice that has modular capability  
1338 for clean access.

1339 ● Identify procurement source and cost for potential purchase of a rapid hole qualifier (temperature  
1340 and caliper) for field scientist use in borehole logging applications.

1341 ● Establish the IDP Science Requirements for identification and planning of borehole maintenance  
1342 and fluid maintenance over time, including removing (or lowering) drilling fluid from a borehole  
1343 (for example for freezing in a sensor).

- 1344 ● Create a second, updated Blue Ice Drill.
- 1345 ● Evaluate options for new drilling fluids for future ice and rock drilling projects, in collaboration
- 1346 with international partners.

1347

1348 **Priority 3 (needed in 3 to 5 years)**

- 1349 ● Continue investigation and modifications of the RAM 2 Drill to achieve the 100 m depth
- 1350 goal reflected in the system Science Requirements.
- 1351 ● Establish the Science Requirements for retrieving sidewall ice samples at specific depths in
- 1352 an existing borehole without using an ice coring drill.
- 1353 ● Create a feasibility paper for using shallow drill fluid columns for ice coring.

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1640 **Acronyms**

- 1641 AGAP: Antarctica's Gamburtsev Province  
1642 ANDRILL: Antarctic Drilling Project  
1643 AO: Arctic Oscillation  
1644 ARA: Askaryan Radio Array  
1645 ARIANNA: Antarctic Ross Ice shelf Antenna Neutrino Array  
1646 ASIG: Agile Sub-Ice Geological (drill)  
1647 AUV: Autonomous Underwater Vehicle  
1648 BLWG: Borehole Logging Working Group  
1649 CReSIS: Center for Remote Sensing of Ice Sheets  
1650 DISC: Deep Ice Sheet Coring  
1651 DOSECC: Drilling, Observation, Sampling of the Earths Continental Crust (drilling service)  
1652 EDC: EPICA Dome C  
1653 EGRIP: East Greenland Ice core Project  
1654 EHWD: Enhanced Hot Water Drill  
1655 ENSO: El Niño Southern Oscillation  
1656 EPICA: European Project for Ice Coring in Antarctica  
1657 G-2IC: Generation-2 Ice Cube  
1658 GISP2: Greenland Ice Sheet Program II  
1659 GRIP: Greenland Ice Core Project  
1660 GZK: Greisen-Zatsepin-Kuzmin  
1661 HCFC: Hydrochlorofluorocarbon  
1662 ICECAP: A project name, not an acronym  
1663 ICWG: Ice Core Working Group  
1664 IDP: Ice Drilling Program  
1665 IPCC: Intergovernmental Panel on Climate Change  
1666 IPICS: International Partnerships in Ice Core Sciences  
1667 LIG: Last Interglacial  
1668 LRSP: Long Range Science Plan  
1669 NEEM: North Greenland Eemian Ice Drilling  
1670 NEGIS: Northeast Greenland Ice Stream  
1671 NGRIP: North Greenland Ice Core Project  
1672 NRC: National Research Council  
1673 NSF: National Science Foundation  
1674 PINGU: Precision IceCube Next Generation Upgrade  
1675 RAID: Rapid Access Ice Drill  
1676 RAM: Rapid Air Movement (drill)  
1677 ROV: Remotely Operated Vehicle  
1678 SAB: Science Advisory Board  
1679 SALE: Subglacial Antarctic Lake Environment  
1680 SCAR: Scientific Committee on Antarctic Research  
1681 SHALDRIL: Shallow Drilling on the Antarctic Continental Margin  
1682 SleGE: Sub-Ice Geological Exploration  
1683 SPICE: South Pole Ice  
1684 WAIS: West Antarctic Ice Sheet

1685 WISSARD: Whillans Ice Sheet Subglacial Access Research Drilling