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The final document will be formatted.

Ice Drilling Program

Long Range Science Plan 2021-2031

Prepared by the U.S. Ice Drilling Program in collaboration with its
Science Advisory Board and with input from the research community

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Ice Drilling Program (IDP)

Mary R. Albert, Executive Director, Dartmouth
Louise Huffman, Director of Education and Public Outreach, Dartmouth
Kristina Slawny, Director of Operations, University of Wisconsin Madison
Joseph Souney, Project Manager, University of New Hampshire
Blaise Stephanus, Program Manager, Dartmouth
Mark Twickler, Director of Digital Communications, University of New Hampshire

Science Advisory Board to the IDP

Chair: Jill Mikucki, University of Tennessee Knoxville
T.J. Fudge, University of Washington
Brent Goehring, Tulane University
Bess Koffman, Colby College
Erin Pettit, Oregon State University
Martin Truffer, University of Alaska Fairbanks
Trista Vick-Majors, Michigan Technological University
Paul Winberry, Central Washington University

Snapshot of updates in the 2021 Long Range Science Plan

The Long Range Science Plan 2021-2031 includes updated information in the following areas:

- *The listing of Recommended Technology Investments was updated based on discussions with the community and prioritized by the IDP Science Advisory Board.*
- *Description of the current state of science and associated references were updated throughout.*
- *The Science Planning Matrices (Tables 4, 5, 6) were updated in accordance with community input*

Executive Summary

The rapid rate of current climate change creates urgency to understand the context of abrupt changes in the past and also the need to predict rates of sea level rise. Glaciers, ice sheets, and subglacial environments contain evidence of past atmospheric composition and climate which enables understanding current and future climate evolution. In addition, understanding glacial dynamics, stability of ice sheets and ice sheet response to climate change are imperative for predicting rapid sea level rise. The under-explored subglacial realm preserves unique biological, geochemical, and geological environments. For all of these areas of science, extracting this evidence involves drilling and coring into and through glaciers and the polar ice sheets, a specialized and challenging endeavor that requires extensive planning, technology, and logistics.

The Ice Drilling Program (IDP) was established by the National Science Foundation (NSF) to lead integrated planning for ice coring and drilling, and provision of drills and drilling expertise. The IDP and its Science Advisory Board (SAB) update this Long Range Science Plan (LRSP) annually, in consultation with the broader research community. The purpose of this plan is to articulate goals and make recommendations for the direction of U.S. ice coring and drilling science for a wide variety of areas of scientific inquiry, and to make recommendations for the development of drilling technology, infrastructure and logistical support needed to enable the science. A companion document, the Long Range Drilling Technology Plan, is available on the Icedrill.org website and it provides details about drills and drilling expertise available through IDP. Specific recommendations for the next decade include the following areas of science, as described in more detail within the report:

Past Climate change, from the recent past to several million years ago: Drilling of spatially-distributed ice cores and boreholes spanning the past 200 to 40,000 years provides evidence for a variety of scientific questions. Shallow ice coring enables understanding of the climate signals in remotely-sensed data and assessment of regional environmental changes. Determining patterns of hydroclimate variability, aerosol deposition, climate feedbacks, and the past extent of glaciers and ice sheets requires ice coring in high-altitude and high-latitude regions. Targeted ice coring at coastal domes and coastal ice caps and along the dynamic Amundsen Sea Coast of West Antarctica, and along the Greenland coasts facilitates investigation of current ice, ocean, and atmospheric dynamics. Deeper cores in Northwest Greenland (Qaanaaq) and South Dome, Greenland are in the planning stages, while preparation for drilling at Hercules Dome, Antarctica, are ongoing and will address questions of transient and equilibrium response of ice sheets to warmer-than-present climates. Global-scale questions about the drivers of Earth's climate system and past atmospheric composition back to the Mid-Pleistocene drive the need for retrieving older ice from Antarctica. Blue-ice paleoclimate records from Mt Moulton, Taylor

Glacier, and Allan Hills may provide unlimited samples for atmospheric and ultra-trace component studies and enable access to ice older than a million years.

Ice dynamics and glacial history: Rapid changes in the speed of fast-flowing outlet glaciers and ice streams observed over the past decade create an urgency to understand the dynamics of outlet glaciers and ice sheets. Efforts to improve understanding of ice-ocean interaction, measurement of subglacial geothermal fluxes, basal properties, subglacial hydrology, ice rheology, variation of surface accumulation, and retrieval of short cores of subglacial bedrock at targeted sites for cosmogenic and luminescence dating are all important and described in more detail within this plan. Ice-sheet models that incorporate realistic physics and dynamics at appropriate spatial and temporal scales are needed to predict the "tipping point" when ice-loss becomes irreversible, resulting in ice-sheet collapse and rapid sea-level rise. Observational data are needed to develop and validate the models.

Subglacial geology, sediments, and ecosystems: Bedrock, sediments, and ecosystems existing within and beneath ice sheets have been unexplored in the past due to the lack of rapid access drills. The IDP Agile Sub-Ice Geological Drill (ASIG) has retrieved rock from beneath ice at Pirrit Hills. Development of the Rapid Access Ice Drill (RAID) is underway to retrieve rock cores under very deep ice. Direct sampling of the bedrock is needed to validate models of cratonic growth related to supercontinent assembly in the Mesoproterozoic between about 2.0 and 1.1 billion years ago and for constraining the Phanerozoic geological, tectonic and exhumation history of the Antarctic continent. Direct measurements at grounding zones of fast-flowing ice streams and outlet glaciers, and data from sub-ice-shelf ocean cavities are crucial for predicting future ice sheet dynamics and sea level rise. Significant wet environments exist below ice sheets and glaciers; sampling of subglacial sediments and ecosystems will establish the diversity, and physiology of microbes and their relationships to past climates and their current ecosystem function below the ice.

Ice as a scientific observatory: Polar ice sheets and mid-latitude ice caps also serve as a unique platform to conduct observations and experiments concerning seismic activity, planetary sciences and experimental astrophysics. Specifically, borehole logging of both fast-access holes and boreholes originally drilled for ice cores are needed to fully exploit the histories of climate and ice dynamics preserved within the ice. In-ice physics and astrophysics experiments (e.g., IceCube) make use of polar ice as a clean, highly stable, low-background, and transparent detection medium for observation of sub-atomic particle interactions. Future planned projects (e.g. the Askaryan Radio Array (ARA) and Generation-2 Ice Cube) require multiple boreholes drilled to at least 150 m deep (ARA) and 2,500 m deep (G-2IC) and significant calibration studies of the surrounding ice volume. Ice sheets are a quiet platform for seismic monitoring; the South Pole Remote Earth Science and Seismological Observatory has seismic equipment installed in boreholes about 300 m

below the surface. A similar seismic observation network is being initiated on the Greenland Ice Sheet.

Recommended life cycle cost and logistical principles

Although drills already exist that can achieve some science goals, new drilling technologies are needed to accomplish science goals planned for the next decade. In the past decade there has been an increase in research proposed by the ice science community but the NSF budget has been generally flat. The following principles guiding development of new drills and technologies are recommended:

- Designs require that the supporting logistical needs do not impede execution of the science.
- While developing the science requirements, logistical issues such as weight, size, costs, and time for development, must be clearly defined and transparent at the initial stage of planning. Scientists and engineers working together through IDP must assess the impact of changes as they arise during the engineering design and fabrication process.
- Drills, major drilling subsystems, and accompanying technology must be developed with consideration of potential use in future projects. The drills and technology must be versatile and well documented so that they can be used, maintained, and repaired by other engineers.
- Major drilling systems (e.g., sondes, winches, control and other major electronics systems) should be fungible to the maximum extent possible. Major component interchangeability and logistical agility should be essential deliverables for all new drilling technology projects.
- Engineering design teams must include individuals with field experience using appropriate ice drilling technology and/or other relevant field experience.
- Heavy traversing capability is urgently needed to improve access to many scientifically important regions of the Antarctic and Greenland Ice Sheets.

Recommended Technology Investments

The following investments in drilling technologies are needed to accomplish science goals planned for the next decade. Investments prioritized by time, from consensus of the IDP Science Advisory Board, include:

Priority 1 (needed this year):

- Maintain and upgrade agile equipment in inventory, including: Hand Augers, Sidewinders, the 4" Electromechanical Drills, the 3" Electrothermal Drill, the 3.25" Badger-Eclipse Drills, the Stampfli Drill, Logging Winches, the Small Hot Water Drills, the Blue Ice Drill, the Prairie Dog, the Agile Sub-Ice Geological Drill (ASIG), the Rapid Air Movement Drill (RAM) Drill, and the Winkie Drill.

- Develop the Detailed Design for clean hot water basal ice coring mechanism for a hot water drill.
- Finish a feasibility white paper on logistically effective methods for interdisciplinary projects that seek to retrieve rock, basal ice, sediment, and water from West Antarctic (e.g., Mt. Resnik) and East Antarctic sites (e.g., within Wilkes Basin)
- Develop IDP Science Requirements for collecting a small amount (e.g. chips or less than 10 cm) of sub-ice rock using a lightweight tethered ice core drills, for example the Stampfli drill.
- Develop IDP Science Requirements for collecting a small amount (e.g. 10 cm to 1 meter) of sub-ice rock using an intermediate or deep ice core drill in a fluid-filled hole, for example the Foro 3000 drill.
- Begin construction of the 700 drill.
- Develop the updated IDP Conceptual Design and Detailed Design for a clean Scalable Hot Water drill that minimizes its logistical footprint including fuel supply.
- Establish the IDP Science Requirements for identification and planning of borehole maintenance and fluid maintenance over time.
- Establish the IDP Science Requirements for removing (or lowering) drilling fluid from a borehole (for example for freezing in a sensor).
- Evaluate options for new drilling fluids for Herc Dome and other ice and rock drilling projects.
- Investigate a lighter weight source of power to replace generators for drilling systems, in order to ease demand on logistics, including renewable energy.
- Finish building a stand-alone Foro 3000 Drill as per the IDP Science Requirements.

Priority 2 (needed in the next 3 years)

- Build a Scalable Hot Water Access drill for creating access holes in ice that has modular capability for clean access.
- Identify procurement source and cost for potential purchase of a rapid hole qualifier (temperature and caliper) for field scientist use in borehole logging applications.
- Resolve logging winch electrical noise issues.
- Finish building a second Blue Ice Drill for wide-diameter drilling to 200 m.
- Continue to evaluate options for exploring/testing shallow drill fluid columns.

Priority 3 (needed in 3 to 5 years)

- Continue investigation and modifications of the RAM 2 Drill to achieve the 100 m depth goal reflected in the system Science Requirements.
- Acquire components for a stand-alone Intermediate Depth Drill (1,850 m) with updated control system and other repairs so that it can be deployed at the same time as the Foro 3000 Drill.

Introduction

The rapid rate of current climate change creates urgency to understand the context of abrupt changes in the past and also the need to predict rates of sea level rise.

A more sophisticated and predictive understanding of the mechanisms of climate change and the effects on sea level change are needed to plan for the future. Glaciers, ice sheets, and subglacial environments contain records of past climate and ice thickness, which provide evidence crucial to understanding future climate.

Ice core records have led to many important discoveries; for example, the discovery from the Greenland Ice Sheet Project 2 ice core showed that dramatic changes in climate can occur abruptly, in less than ten years (NRC, 2002); this revolutionized climate science and also has important implications for policy. This finding contributed to the fundamental understanding of the climate system, and was a contributing factor to the 2007 award of the Nobel Peace Prize to the Intergovernmental Panel on Climate Change (IPCC) for climate science. The WAIS Divide Core established the benchmark carbon dioxide record for the most recent glaciation. U.S. and U.K. scientists are collaborating on current studies of the unstable Thwaites Glacier in West Antarctica to investigate the possibility of large sea level change in the near future (e.g., Joughin et al., 2014; Scambos et al., 2017). A recent study of bedrock below the summit of the Greenland Ice Sheet raises questions about the ice sheet's resilience to climate change (Schaefer et al., 2016; Christ et al., 2020). Many other basic questions about Earth's climate system remain unresolved, and new scientific plans, in both Antarctica and Greenland, will address a variety of questions.

Rapid changes in the speed of fast-flowing outlet glaciers and ice streams observed over the past decade have created an urgency to understand the dynamics of outlet glaciers and ice sheets. It has long been recognized that basal conditions exert strong control on the flow of glaciers and ice sheets; and boreholes drilled to the bed have been used to deploy instruments to measure basal properties (e.g., Engelhardt et al., 1990; Engelhardt and Kamb, 1998; Kamb, 2001; Truffer et al., 1999, 2006). These fundamental observations have advanced our understanding, and it is clear that spatial and temporal distribution of sediments and hydraulic conditions at the bed are key to understanding rapid changes in speed of glacial flow. Furthermore, in cases where the bed of outlet glaciers is slippery, perturbations at the grounding line propagate inland over short timescales (order of decades), which has the potential for rapid drawdown of inland ice (Payne et al., 2004; Shepherd et al., 2004; Price et al., 2008; Joughin et al., 2014; Rignot et al., 2014). Perturbations at grounding lines are triggered by changing ocean temperature and circulation (Jenkins et al., 2010), and/or subglacial hydrology or sediment dynamics (Anandakrishnan et al., 2007; Alley et al., 2007; Carter & Fricker, 2012; Christianson et al., 2012; Horgan et al., 2012). Defining the processes that control the dynamic stability of glaciers and ice sheets is crucial for predicting their response to future possible greenhouse gas emission scenarios. The greatest uncertainties in sea level rise projections for the 21st century are associated with the possibility of rapid dynamic responses of the ice sheets to climate and sea level change.

Subglacial environments represent a resource of deep time understanding; a resource that remains largely untapped. Most of our knowledge about subglacial environments comes from geophysical remote sensing

and sparse data retrieved from access holes drilled to the bed, or sub-ice-shelf cavities. More detailed observations are needed to map and understand the variety and complexity of deep ice, subglacial geology, and the interface between them. The lithosphere under the Antarctic and Greenland ice sheets remains unknown except by extrapolation from coastal outcrops and remotely-sensed geophysical data. Subglacial environments also house records of past ice sheet dynamics and longer-term paleoclimatic histories in their sediment and rock basin archives. Recovering these records for intervals of past warm periods will contribute to our understanding of future ice sheet behavior under a warming climate.

New and emerging studies show that subglacial environments harbor unique microbial ecosystems and that these microbial communities are metabolically active and thus play a critical role in subglacial weathering (Martinson et al., 2013; Christner et al., 2014; Michaud et al., 2016). Recent studies have illustrated the potential for subglacial water to supply critical carbon and nutrients to sub-ice-shelf waters (Vick-Majors et al., 2020, Hawkings et al., 2020), however, the extent to which microbial activity alters the chemistry of subglacial efflux and the effects of that efflux on global processes remain outstanding questions. There is considerable scientific and public interest in subglacial environments, particularly in relation to the discoveries of subglacial lakes beneath the Antarctic Ice Sheet and the unique life forms they may harbor. Microorganisms that exist under permanently dark and cold subglacial conditions have broadened our understanding of the phylogenetic and metabolic diversity of life on Earth (Achberger et al., 2016; Mikucki et al. 2016), and may help inform our search for extraterrestrial life.

Technological developments are required to integrate geological drilling technologies with those of ice drilling, including clean access. The U.S. Antarctic Program complies with the Antarctic Treaty and other treaties to uphold protection of the environment, including activities that involve drilling through the ice. Challenges with this drilling approach include keeping access holes open for long periods and operating under conditions of differential ice flow movement. Given the pristine nature of Antarctic subglacial environments in particular, the Scientific Committee on Antarctic Research (SCAR) has developed a Code of Conduct for access in order to *“recognize the value of these environments and the need to exercise wise environmental stewardship.”*

The U.S. ice coring and drilling community has led and participated in fundamental and vital scientific discoveries for more than 60 years. These discoveries require drilling and coring of glaciers and the polar ice sheets, a specialized and challenging endeavor that requires extensive planning, technology, and logistics. This Long Range Science Plan was established by the U.S. Ice Drilling Program (IDP), working with its Science Advisory Board (SAB), associated IDP working groups, and the broader research community, to articulate direction for U.S. ice coring and drilling science for the next decade. The science direction provides a foundation as well as direction for the Long Range Drilling Technology Plan for developing some of the new drills and technology. These paired plans enable the community to develop well-coordinated proposals while allowing the NSF to plan for budgets and logistics to facilitate the science. SAB-recommended updates to the IDP Long Range Science Plan are posted to the icedrill.org website each spring, with listserv invitations for comments and suggestions to enable broad community input. The document is then revised, approved by the SAB and the final version for the year is posted to the icedrill.org website in summer.

Science goals articulated in this document are all interconnected, but for convenience in associating science endeavors with appropriate drilling technology, they are described in four categories: climate change; ice dynamics and glacial history; subglacial geology, sediments, and ecosystems; and ice as a scientific observatory. These four goals and objectives are described below, together with an outline of their respective needs for drilling technologies. Planning matrices are also developed to provide a timeline for the development of technologies, so that the support for the science will be ready when needed.

Ice Coring and Drilling Science Goals

I. Past Climate Change

Earth's climate system involves local, regional, hemispheric, and global phenomena. It is impossible to understand global climate without understanding both individual components of the system and the system as a whole, as evidenced by data from a large number of locations and over a range of time scales. Issues articulated by many U.S. scientists (e.g., ICWG, 2003) were central to the themes in the International Partnership in Ice Core Sciences (IPICS) white papers (Brook and Wolff, 2006); hence a number of the categories below reflect those themes.

1. Industrial and Instrumental Period: Spatially distributed evidence from ice cores spanning the industrial (last 200 years) and instrumental (last 100 years) period is needed to establish ice core records of human impacts on the climate, cryosphere and atmosphere, study modern surface processes, and calibrate models and remote sensing data with in situ data. As shallow ice cores (generally <200 m), these records are relatively easy to recover and consequently more records can be collected to evaluate spatial patterns of change.

Over the past 200 years, human activities have had significant impact on atmospheric composition and climate, yet the impacts in polar and remote high-latitude and high-elevation regions are not fully understood. Shallow ice coring programs have been, and will continue to be done through individual or small-group projects at targeted sites (e.g., ice coring in mid-latitude temperate glaciers or in selected areas of the Arctic and Antarctic such as Summit and Disko Bay Greenland) and internationally coordinated scientific traverses (e.g., International Trans-Antarctic Science Expedition, Norwegian-U.S. Scientific Traverse of East Antarctica). While shallow coring has been done in several locations, more cores are needed in order to understand whether observed patterns are regional, hemispheric, or global. Through a combination of over-snow science traverses and coordinated individual site efforts, an extensive array of relatively easy-to-recover ice core records, driven by individual and group proposals, is a mainstay of the ice coring community with the following objectives:

- Determine accumulation rate and temperature changes on the Greenland and Antarctic ice sheets and in alpine regions where instrumental records are rare.

- Understand changes in the chemistry and isotopic composition of the atmosphere during the Industrial Period, including greenhouse gases, acidic species, oxidants, toxic metals, and trace species such as carbon monoxide and hydrocarbons.
- Understand stability and rapid changes along coastal areas of the West Antarctic Ice sheet (WAIS). Ice core records, strategically placed on ice domes along the Amundsen Sea coast, will provide high-resolution (annual) records of natural variability in ice, ocean, and atmospheric dynamics in which to place the recent observations in context.
- Constrain surface mass balance processes including accumulation, surface melt, runoff and refreezing, and evaluate areas of water retention in perched water tables and aquifers in Greenland and Antarctica. These data can also be used to ground-truth high-resolution climate models.
- Improve understanding of relevant physical and chemical processes related to snow deposition and post-depositional changes (including metamorphism, in situ chemical processes, interactions with cosmic rays, etc.) and their effects on atmospheric chemistry preservation and interpretation of geochemical signals (including atmospheric) at larger depths.
- Calibrate snow/firn/ice properties measured remotely (e.g., borehole, ground, airborne, and satellite-based measurements) with in situ data, thereby allowing interpolation based on remote sensing data.
- Produce detailed temporal and spatial (regional-scale) maps of climate and environmental parameters (e.g., temperature, accumulation rate, atmospheric and snow chemistry), and anthropogenic impacts.
- Develop an inventory of microbes within ice to improve understanding of the role of microbes related to geological, chemical, and climatological changes.
- Improve records of global and local volcanism for climate forcing and geohazard studies.

--photo here--

Caption: Part of an ice core retrieved from Mt Hunter Plateau of Denali exhibits layering and dust carried to the area from afar.
Photo credit: Brad Markle, Univ. Washington

Individuals and small groups conduct studies of these types across glaciological settings ranging from the Greenland and Antarctic ice sheets, to ice caps and alpine glaciers in low, mid, and high latitudes. Versatile drills required for 200-year ice coring exist in the current U.S. inventory, and are in high demand; they need to be upgraded and continuously maintained so that they are functional and can be quickly deployed to the field. Requirements for drills to achieve these and other ice coring goals are listed in table one. The Long Range Drilling Technology Plan describes the agile drills in the IDP inventory in detail and discusses their current condition.

2. Pre-Industrial Baseline: The late Holocene (ca. the last two millennia) is an important temporal focus because it is long enough to allow investigation of annual to centennial climate variability, yet short enough that relevant climate boundary conditions have not changed appreciably. Thus, this period represents a critical pre-industrial baseline against which to compare 20th century changes in climate, the cryosphere, and atmospheric composition and chemistry. Existing quantitative reconstructions of climate spanning the past two millennia continue to be debated, in part due to a lack of annually-resolved records prior to 1600 B.P. in many areas, and due to the highly regional nature of many climate processes. A

coordinated international effort to recover a spatial array of annually resolved and calibrated 2,000-year ice core records has several primary objectives:

- Determining regional and high-resolution temporal patterns of temperature, precipitation, sea ice extent, and atmospheric composition and chemistry to better understand climate forcing and particularly climate feedbacks that will also operate in the near future.
- Evaluating 20th century warming, precipitation, atmospheric circulation, sea ice, and atmospheric composition and chemistry changes in the context of the past 2,000 years.
- Establishing the extent and regional expression of the so-called Little Ice Age and Medieval Climate Anomaly phenomena, and constraining their relationships with regional climate patterns like the North Atlantic Oscillation (NAO), Arctic Oscillation (AO), El Niño Southern Oscillation (ENSO), and Monsoons.
- Calibrating local, regional, and global climate models against a recent but sufficiently long pre-anthropogenic period.
- Determining the sensitivity of alpine glaciers and ice sheet margins to the relatively warm Medieval Climate Anomaly and relatively cold Little Ice Age, with implications for the impact of future warming on water resource availability and sea level rise.
- Quantifying spatial and temporal patterns of climate-forcing mechanisms that are regionally variable (e.g., greenhouse and reactive gases, sulfate, terrestrial dust and associated biological material, black carbon aerosols), and the record of solar variability.
- Assessing the relative roles of anthropogenic and natural forcing on climate evolution prior to and into the industrial period.
- Quantifying anthropogenic pollution sources and emission levels prior to the industrial revolution, from early metal smelting activities.

-Photos here --

Caption: Scientific drilling on the Mt. Hunter Plateau of Denali provides a 2,000-year record of precipitation and atmospheric circulation in Central Alaska. Drilling at this site was accomplished by wind and solar energy without the need for gas-fueled generators. Photo credit: Top) *Seth Campbell, CRREL*; Bottom) *Dom Winski, Dartmouth*

New coring associated with this effort will include Arctic, Antarctic, and mid-latitude sites. Recent and desired future U.S. or U.S./International efforts include Central Alaska Range; British Columbia (Mt. Waddington); Detroit Plateau on the Antarctic Peninsula; multiple locations on the coastal WAIS ice shelves and ice domes; the Aurora Basin in Antarctica; Hercules Dome (the 2,000-year record would be part of a deeper core); Greenland coastal ice caps, high accumulation rate sites in Greenland, and Northwest (Qaanaaq) and South Dome sites (the 2000-year record would be part of a deeper core). This list is not exclusive, but illustrates the diversity of discussions within the research community.

1. Large-Scale Global Climate Change: Large changes in global climate driven by external forcing have involved significant interactions between ice sheets, carbon cycle, vegetation, dust, ocean and atmospheric circulation resulting in rapid changes in regional climate. Understanding Earth system

dynamics especially in times of rapid transitions is critical to making improvements in current Earth system models for assessment of future climate. Incomplete evidence about processes and dynamics of abrupt millennial scale change and regional impacts requires evidence from ice cores to develop more complete knowledge of the underlying mechanisms. Ice cores are uniquely placed to provide the contrasting polar elements of climate in high resolution and are the only source of past atmosphere allowing measurements of greenhouse and other trace gases. Scientific challenges include the following (mainly from the International Partnerships in Ice Core Sciences (IPICS) terminations and seesaws initiative):

- Further develop ice core proxies for different aspects of the Earth system, for example to reconstruct conditions at the ocean surface (e.g., sea ice, marine biological productivity, ocean evaporation conditions), and in the boreal continental biosphere (e.g. forest fires, land ice extent).
- Continue to improve understanding of climate forcings on millennial timescales (greenhouse gases, solar irradiance, aerosols)
- Improve the absolute and relative chronologies of individual ice cores, in both the ice and gas phase, and construct consistent multi-ice-core chronologies.
- Develop methods to synchronise ice core records to those from other palaeoclimate archives, in particular for previous glacial cycles where radiocarbon is unavailable. [2]
- Identify further sites to complement the spatial picture and implement plans to fill the “gaps”: examples include further cores at coastal domes around Antarctica.
- Apply newer methods to improve the resolution of data from some existing sites. [2] Quantify and understand the spatial and temporal evolution of rapid climate changes, and assess how this varies with background climate (orbital forcing, greenhouse gas concentration, land ice masses).
- Construct, using ice cores carefully synchronised to other records, the sequence of events (including forcings and responses) through several glacial-interglacial transitions at the highest resolution possible.
- Use these reconstructions with Earth system modelling to provide a stringent test of mechanisms. This will require an increase in modeling capability to assess changes at sufficient resolution through multiple terminations

Under the auspices of IPICS, the international scientific community aspires to a network of ice cores. Specific U.S. contributions to this network include the completed WAIS Divide core and the South Pole ice (SPICE) core, and a future core at Hercules Dome. In Greenland, two potential sites in the Northwest (Qaanaaq) and at South Dome will also contribute to IPICS goals. The projects may vary in scope and logistical needs, but many are envisioned to be drilling campaigns conducted in two or three seasons with minimal logistics. Site-specific records of climate and environmental change are the primary objective; it will not be necessary to undertake the full suite of measurements possible in an ice core, although clearly such measurements provide data for a variety of future projects. The Foro 1800 drill was used to successfully drill the South Pole Ice Core to a depth of 1,751 m (age ~ 49,000 years), and the Foro 3000 drill is being readied for use at Hercules Dome. Individual and small group projects targeting specific aspects of climate and atmospheric variability on millennial to glacial timescales can be conducted at low-accumulation sites such as Dome C, Antarctica or at ice margin sites with agile drill systems and moderate logistics requirements.

4. High-resolution Records of the Last Interglacial; A Warm-Earth Analog: The last interglacial (LIG) period (~130k to 110k years ago) was warmer than present due to differences in Earth's orbital configuration, and can provide clues about how the Earth will respond as human activities continue to force global warming. Critical science priorities for ice cores spanning the Last Interglacial include:

- Determining whether the West Antarctic Ice Sheet experienced partial or total collapse during the LIG, and determining the extent of the Greenland Ice Sheet during this warmer time. These objectives are critical for constraining sea-level rise estimates in a warmer world.
- Quantifying the temperature, precipitation, atmospheric circulation, and sea-ice extent of Greenland and Antarctica during the LIG.
- Establishing whether rapid climate change events occurred during the warmer world of the LIG.
- Determining whether the lack of an abrupt climate change during the deglacial warming (i.e. Bolling Allerod warming and Antarctic Cold Reversal) contributes to a climate "overshoot" and a warmer interglacial.
- Comparing the evolution of the LIG with our present interglacial period, the Holocene.
- Investigating glacial inception at the end of the LIG.

Existing ice core records of the last interglacial are primarily from low accumulation sites in East Antarctica are insensitive to changes in the marine segments of the ice sheets. The detailed behavior of polar climate, greenhouse gases, ice sheet size, and other earth system attributes recorded by ice cores are not well known for this period, and require high-accumulation conditions. Results from the North Greenland Eemian (NEEM) ice core in Greenland, and similar results from other Greenland ice cores, have shown that the Eemian record located there is at least partially recoverable, but not in stratigraphic order. Large volumes of ice from the last interglacial have been shown to outcrop at the surface of Taylor Glacier, Antarctica; however a complete and undisturbed stratigraphic sequence of the warming from the climate period MIS-6 has yet to be recovered.

- Photo here -

Caption: The bubbles visible in this piece of ice from an Antarctic ice core contain carbon dioxide and other gases that were trapped in the ice when formed many thousands of years ago. Ice cores provide the only natural archive of ancient air. Photo credit: Oregon State University.

The search for sites with unfolded ice will continue in both polar regions along with efforts to interpret folded ice; likely targets are relatively high accumulation sites in Antarctica, such as Hercules Dome, where last interglacial ice is likely to be preserved, and possible new sites in Greenland, including near the Camp Century site in Northwest Greenland.

The U.S. community, represented by the IDP Ice Core Working Group (ICWG), has prioritized Hercules Dome as the next deep ice core site in Antarctica, due to its likely preservation of ice from the last interglacial period (Jacobel et al., 2005) and its sensitivity to a potential collapse of the WAIS (Steig et al., 2015), as well as its potential to provide bubble-free ice (below the problematic bubble-to-clathrate transition zone) for gas studies during the last glacial-interglacial transition.

5. Evidence from the ice sheet prior to 800,000 years B.P.: Each time ice cores have extended further back in time they have revealed new facets of climate dynamics. The record from the EPICA core at Dome C extends back to just over 800,000 years, and shows that different styles of glacial-interglacial cycles occur even under superficially similar external forcing. The Dome C site was selected to recover old, but not the oldest ice. Antarctic ice sheet inception is thought to have occurred 35 million years ago, and although basal processes may have removed or altered the very oldest ice in many places, it is likely that ice older than 800,000 years is preserved in East Antarctica.

The primary reason to seek this older ice is to further understand one of the major puzzles of climate system history: Why did the climate system change from a dominantly 41k- to a 100k- year glacial cycle about one million years ago? Numerous research objectives related to this transition, and the earlier time period, could be addressed with ice core records extending back ~ 1.5 million years, including:

- Evaluating the CO₂-climate relationship prior to 800 ka, to determine whether the change to 100-kyr cycles and/or the long-term cooling trend from 1.5 – 0.8 Ma was related to changes in greenhouse gas concentrations.
- Clarifying whether 23k-year climate cycles are present in ice core records prior to the transition to 100-kyr cycles around 1 Ma. The 23k-year cycles are not present in marine proxy records of this age, but are present in both marine records and ice cores after the transition.
- Investigating the high-resolution nature of glacial transitions during the 41k-year world.
- Determining if rapid climate change events like Dansgaard-Oeschger events were present during the 41k-year world.

- Photo here -

Caption: 100,000-year 'sawtooth' variability in Antarctic climate over the last 800,000 years is mirrored by generally, similar variability in atmospheric carbon dioxide (as well as methane and nitrous oxide, not shown) and global ice volume inferred from deep ocean oxygen isotope records from marine calcium carbonate. Whether Antarctic climate followed the ice volume record prior to this time, when ice volume records are dominated by a 40,000-year period, is not known, neither are the mean levels of greenhouse gases and the temporal variability of those levels. *Figure from Severinghaus et al. (2013).*

There are two complementary, but very different, ways of accessing ice older than 800,000 years. The first is drilling at very low accumulation rate sites in interior East Antarctica. This has the advantage of recovering a continuous record, which, in the younger part, can be compared to other ice cores (an important consideration for drilling at very low accumulation sites where record integrity may be an issue).

The second method is to make use of "blue ice" sites such as Taylor Glacier (Aciego et al., 2007), Mt. Moulton (Dunbar et al., 2008) and Allan Hills (Spaulding et al., 2013; Higgins et al., 2015) where old ice may outcrop at the surface via slow ablation or be present in the shallow subsurface. Continuous records require careful site selection, however discoveries are possible from sites with easier access, through smaller and less expensive projects. A site in the Allan Hills (Kehrl et al, 2018) has been shown by ice penetrating radar to likely have a continuous record to ~250 ka with several hundred more meters of ice below. Different drilling requirements are needed for the two approaches. Development of blue ice sampling techniques should continue, given the potential for large volume sampling, very old ice (see

below) and the possibility that continuous ice core records will not be discovered. Consideration of sites where only old ice might be preserved (for example areas where there is no accumulation today but has been in the past) should also continue.

The IPICS “Oldest Ice” workshop resulted in a paper (Fischer et al., 2013) describing the state of knowledge of possible oldest ice sites; although it is possible to use modeling to identify possible locations (Liefvering and Pattyn, 2013) it was the general conclusion that more reconnaissance was needed before choosing a site. The European project “Beyond EPICA” has selected a site near Dome C for deep drilling. The site is shallower than the existing Dome C ice core site to limit the chance of basal melting. Locations near Dome Fuji are also under consideration for a Japanese led deep core. Choosing a location with confidence is still difficult; mainly due to poorly-known geothermal heat flux. Determination of the spatial variability of geothermal heat flux is critical to the identification of potential drilling sites for oldest ice. Large regions which may have the optimal site conditions for Oldest Ice, remain unexplored. The Rapid Access Ice Drill, currently in a testing phase, should be able to quickly create access holes for spatially-distributed measurements of geothermal heat flux to facilitate site selection. New and ongoing radar, laser altimetry, gravity and magnetic data from ICECAP and Antarctica’s Gamburtsev Province (AGAP) airborne surveys are helping identify potential sites, but additional observations and model calculations are needed. In Greenland, locations on the west side of the east mountain range where the first ice sheet originated might result in ice more than one million years old. Since the stratigraphy is likely to be disturbed in that area, methods for dating ice that is not in order stratigraphically should be further developed before drilling for ice older than 800,000 years in Greenland.

Rapid sampling of or access to the near basal region of the East Antarctic ice sheet is needed for site selection for the oldest ice project because temperature and heat flow measurements are needed to constrain models of ice sheet dynamics that are needed to predict potential locations of old ice. The Rapid Access Ice Drill (RAID) would be useful for this purpose. In addition, a more agile drill that could create holes as deep as 1,000 m would accelerate discovery. There are complementary international efforts to explore for oldest ice sites; these include a European oldest ice site selection program that involves rapid access with several different new tools under development, including SUBGLACIOR, a novel hybrid mechanical and thermal drill with on board gas concentration and water isotope capability (Alemany and others, 2014).

6. Pre-Quaternary atmosphere: The possibility that very old ice (>1.5 million years) is preserved in special environments (for example, in debris-laden glaciers) in Antarctica (e.g., Yau et al., 2015) is exciting because it would provide a window into the composition of the atmosphere and climate during times when global environmental conditions were very different from today. Such sites will likely range from blue ice locations, where drilling issues are essentially identical to those mentioned above, to debris-laden glaciers or similar environments, which will require specialized drilling equipment; for example the Agile Sub-Ice Geologic Drill (ASIG) has proven to be useful in some cases.

7. Large-volume sampling of climatic intervals and tracers of high interest: Rare isotopes, ultra-trace species, micro-particles, biological materials, and other measurements that have not yet been fully

exploited in ice core research offer new opportunities for discovery if large volumes of ice are made available. Examples include ^{14}C of CH_4 to trace methane hydrate destabilization during past warming events (Petrenko et al., 2017, nano-diamonds, ^3He , and micrometeorites as tracers of extraterrestrial impacts, and ^{14}C of CO as a tracer for atmospheric oxidizing capacity or past cosmic ray flux (depending on site characteristics). In the case of traditional drill sites, multiple cores or replicate coring technology are needed to obtain larger sample sizes, and in situ melting has been suggested (but not yet successfully used) as a means of sampling large volumes of air from deep ice core sites. Blue ice areas such as Taylor Glacier and Allan Hills currently provide the best opportunities for rapid collection of large samples of ancient ice with relatively light logistics (e.g., Buizert et al., 2014). Ice sections ranging in age from Early Holocene to 1 M-year have already been clearly identified at these sites (e.g., Higgins et al., 2015; Korotkikh et al., 2011) and are ready for access/sampling by future projects. Continued studies at these sites that would provide more detailed and complete age maps of the desired outcropping ice areas.

Depending on the site and scientific target, a range of ice drilling and sampling tools may be appropriate. The Blue Ice Drill, Eclipse Drill, Foro 400 drill, 4" drill, hand augers, and chainsaws have all been successfully used. Continuing to maintain the capability to explore and utilize the ice at these sites is desired.

- Photo here -

Caption: Scientists have drilled a large-volume ice core on the Taylor Glacier ablation zone, Antarctica. Bubbles in the ice at the site contain evidence of ancient atmospheric composition. The Blue Ice Drill is an easily-transportable drill capable of retrieving quality firn cores of approximately 9.5 inches in diameter as well as quality solid ice cores of the same diameter up to 70 meters below the firn-ice transition. Photo credit: Jeff Severinghaus.

8. Ancient microbial life: Ice sheets provide chronological reservoirs of microbial cells entombed during atmospheric deposition and studies have shown that microbial DNA and viable organisms can be recovered from ice cores collected from both Greenland and the Antarctic as well as temperate glaciers (e.g., Christner et al., 2001, 2003; Miteva et al., 2004). In addition, the distributions of microbial cells themselves can serve as climatic records in deep ice (Santibáñez and others, 2018). Many questions remain regarding how these organisms survive in deep ice for tens to hundreds of thousands of years, the origin of these airborne microorganisms and what their diversity and biogeographic distribution reveals about climate during deposition. The ability to obtain larger volumes in conjunction with advances in molecular techniques such as metagenomic analyses (Simon et al., 2009) and methods that can amplify smaller quantities of nucleic acids will enable more detailed study of the genomic potential of resident microbes and available preserved organic carbon material (D'Andrilli et al. 2017a,b) and how they integrate with our understanding of ice core ecology. There is interest in investigating the physiology of microorganisms recovered from ice cores to elucidate unique physiological properties that enable them to survive in ice for extended periods of time and that may offer important biotechnological applications (Cavicholi et al., 2002). For example, studies have shown novel, ultra small microbial isolates from deep Greenland glacier ice that may inform on how organisms survive energy deprivation for extended periods of time (Miteva, 2005).

Built on this foundation, recent studies have characterized organic carbon materials within various Antarctic ice cores and shown changing carbon signals measurable from different climates (spanning back to 27,000 years ago; D'Andrilli et al., 2017a, b) and within the Holocene of Arctic and the Antarctic ice cores (Grannas et al., 2006; Xu et al., 2018; King et al. 2019; Vogel et al., 2019; D'Andrilli & McConnell, 2020: In Review). High temporal resolution organic carbon data from the WAIS Divide ice core emphasized the highly complementary nature of carbon surveys with routinely surveyed geochemical assays in paleoclimate atmospheric reconstructions (D'Andrilli et al., 2017a). Notably, the preservation paradigm of geochemicals in ice cores also extends to biological and other organic materials, therefore it will become increasingly important to characterize their concentrations and qualitative nature now, in the ice, before it melts to reconstruct our past, learn about our present, and help better predict their impacts in the future in a warming climate. The inclusion of organic material (OM) chemical characterizations in ice coring efforts improves geochemical and biological paleoclimate atmospheric composition interpretations at broad and fine scales, ice sheet carbon storage assessments and comparisons with other ecosystem OM, our understanding of potential signatures arising from preservation mechanisms, and the ability to predict the fate of carbon in ice sheets in the Arctic and Antarctic.

9. Borehole Array for Spatial Variations in Climate: Although borehole observations do not provide a detailed climate history, an array of boreholes linked to an ice core can provide information on the spatial variability in climate history for any of the ice cores mentioned above. Notably, the WATSON instrument has thus far shown a unique microscale approach to organic matter and microbial distribution within the ice sheets and in the future requires a joining of the microscale, fine scale, and broad scale ice core researchers to best understand small to large spatial variations of materials in ice. (Eshelman et al., 2019; Malaska et al., 2020). See also section IV.1 below.

Summary

Advances in understanding climate require arrays of ice cores with depths ranging from tens of meters to 3,000 m, and the requirements for the coring or sampling vary. Agile drills currently at IDP-Wisconsin should be continually maintained in good working condition so that they can be used for new projects. Clean hand augers and agile drills are needed for biological studies in glaciers. The Foro 3000, capable of coring up to 3,000 m, is being created by modifying the Foro 1650 (aka Intermediate Depth Drill for coring up to 1,650 m). The large-diameter Blue Ice Drill for blue ice areas was used successfully on Taylor Glacier, Antarctica, and elsewhere; with continued science attention to blue ice areas as well as to large-volume sampling in general, an additional Blue Ice Drill may be useful. Estisol-140 has replaced the IsoparK-HCFC-141b combo for use as drilling fluid in deep drilling since HCFC-141b is no longer produced. While Estisol-140 had some issues in the first season of use, changes to the drill and handling procedures in the second season mitigated many of the issues. Neither Isopar-K nor Estisol-140 are "clean" enough for biological and organic matter chemistry measurements in ice cores. For those types of research, the outer layers of the ice cores must be shaved off manually or within the continuous flow analysis melting system, since they introduce large contaminants into the measurements (Christner et al., 2005; D'Andrilli et al., 2017b). Table one lists characteristics for drills needed for the areas of the science outlined above.

691 **Table 1. Requirements of drills for studies of climate change.**

692 Additional information on the drills is given in the IDP Long Range Drilling Technology Plan.

	Diam (cm)	Depth (m)	Ambien t temp(C)	Clean coring ?	Transport type	Site occupancy	Int'l aspects	Drill Name
<200 years	5-7	Horiz.	-20	yes	Backpack	Days	US	-
<200 years	5	15	-30	somet imes	Backpack	Days	US	Hand auger
200 year	7-10	400	-50	no	Twin otter/ lt traverse	Days/ weeks	US	Eclipse, 4", Foro
200 year	7-10	400	-5 warm ice	no	Twin otter/ lt traverse	Days/ weeks	US	Thermo- mechanical
2k array	7-10	<1,500	-50	somet imes	Twin otter/ lt traverse	Weeks/ month	US part of IPICS	700 Drill, Foro 1650
40k array	10+	1-3k	-50	no	Twin otter/ Herc	1-2 seasons	US or shared	Foro 1650, Foro 3000
Interglacial	10+	1-3k	-50	no	Herc	Multiple seasons	US only or US-led	Foro 1650, Foro 3000
>800k years (oldest ice)	10+	3k	-50	no	Herc & traverse	Multiple seasons	IPICS	Foro 3000
Site selection oldest ice	2-4	<1,000	-50	no	Herc & traverse	2 days	IPICS	RAID
>800k years (blue ice)	25	5-20	-40	no	Twin otter	1-2 seasons	US	Blue Ice Drill
Pre- Quaternary atmosphere	7-25 rock-ice mix	200	-40	no	Helicopter	1-2 seasons	US	ASIG
Tracers; large diameter cores	10-25	200 ⁺	-40	no	Helicopter/ Basler/ Traverse	1-2 seasons	US	Blue Ice Drill
Ancient microbial life/ other organic material	25	200 ⁺	-40	somet imes	Helicopter twin otter, herc	1-2 seasons	US	ScHWD

Borehole Array	8	200 to 3.5k	-40	no	Twin Otter/Lt Travers	Week	US	RAID
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II. Ice Dynamics and Glacial History

Rapid changes in speed of fast-flowing tidewater glaciers, outlet glaciers, and ice streams observed over the past decade create an urgency to understand their dynamics. In West Antarctica, ongoing rapid loss of ice in the region of Thwaites Glacier and the Amundsen Sea is occurring, with possible accelerated loss due to ocean-driven melting at the grounding zone and nearby areas (e.g., Scambos et al., 2017). A complete retreat of the Thwaites Glacier basin over the next few centuries would raise global sea level by more than three meters. It is possible that processes such as hydrofracture and ice cliff failure could lead to a more rapid collapse of Thwaites Glacier within the next few decades. Reducing uncertainty in the projected contribution of Thwaites Glacier to sea level rise requires substantial and coordinated collaborations involving a multidisciplinary, international scientific community. The Ross Sea sector of the Antarctic Ice Sheet, with its spatially diverse and changing natural environment, can facilitate process studies through field sampling and data which will be used in evaluating the current state of the ice sheet, quantifying the glaciological and oceanographic processes that may play a role in rapid decay of the ice sheet, and interpreting past ice sheet changes from subglacial and ice-proximal geologic records to understand ice sheet sensitivity to climate forcing on different timescales. In general, predicting responses of glaciers and ice sheets to future possible environmental change requires models that incorporate realistic ice dynamics (Alley and Joughin, 2012). Ice loss on the Greenland Ice Sheet is also happening at a dramatic rate, and contains an additional 7.4 meters of sea level equivalent. Predicting dynamic ice loss of major ice streams, like the northeast Greenland ice stream is a major challenge to the international community. For both the Antarctic and Greenland Ice Sheets, understanding the history of past ice sheet change is key for pinpointing ice sheet sectors most sensitive to climate change. Measurements and observations of present day conditions are needed to develop and validate such models. Properties of the ice and the ice-bed interface exert strong control on the flow of glaciers and ice sheets. Instruments deployed down boreholes drilled to the bed are needed to collect basic data concerning the spatial and temporal distribution of ice properties, sediments, and subglacial hydrology.

Another approach to understand future ice-sheet response to local and global climate is to reconstruct its history. Histories of ice dynamics (thinning and divide location) and climate (accumulation and temperature) can be inferred from observations from ice cores, basal ice samples, and boreholes near ice divides. Ice core and borehole data, including basal ice samples for gas analysis, depth-profiles of age, layer thickness, temperature, ice fabric, and bubble density all provide constraints for ice flow models. For example, the depth-age relationship contains information about past accumulation and past thinning; a thin annual layer at depth could imply either low accumulation in the past or ice sheet thinning (Waddington et al., 2005; Price et al., 2007). Radar-detected layers can also be used to infer the flow history of glaciers and ice sheets and the history contained in the layers is much richer if their age is known (Waddington et al., 2007, Dahl-Jensen et al. 2013); ice cores can be used to date intersecting radar layers. The high quality radio echo sounding data from the Center for Remote Sensing of Ice Sheets (CReSIS) and Operation IceBridge both in Antarctica and Greenland make it possible to detect internal layers reaching to the bedrock. Disturbances, folding, and larger structures are observed that strongly influence the local ice dynamics and point towards the need for more complex and anisotropic ice deformation relations.

Specific observational data needed to improve and validate models of ice sheet response to environmental change include:

1. Basal conditions and geothermal flux: Direct measurements of bed conditions including frozen/thawed bed, basal pore water pressure, slip, and sediment properties and deformation are needed to develop and test realistic models of the controls on the fast flow of ice streams and outlet glaciers. Determination of whether a bed is frozen or thawed requires coupled thermo-mechanical flow models. A necessary boundary condition is a realistic realization of the geothermal flux. Geothermal flux has been determined at a few locations from borehole thermometry, but we expect the geothermal flux to vary significantly over spatial scales of less than 25 km (Fahnestock et al., 2001). In Greenland borehole temperature reconstructions imply low values in south Greenland ($<40 \text{ mW/m}^2$, values of 50 mW/m^2 at Greenland Ice Core Project (GRIP) and Camp Century and higher values at NEEM (80 mW/m^2) and North-GRIP (130 mW/m^2). Until recently the only measurement in West Antarctica was from Siple Dome (69 mW/m^2), but recent borehole temperature measurements from the WAIS Divide borehole indicate a geothermal flux of $\sim 230 \text{ mW m}^{-2}$ (Clow, 2012), and $\sim 285 \text{ mW m}^{-2}$ at Subglacial Lake Whillans (Fisher et al, 2015). Additional measurements are needed to provide boundary conditions for ice sheet models. Based on the data to date, geothermal flux values vary considerably throughout West Antarctica and further investigation is required to provide boundary conditions for ice sheet modeling.

Measurements at the bed of glaciers and ice sheets are hampered because of difficulties accessing the bed, and keeping boreholes open long enough to deploy sensors. Rapid-access drills that are portable and capable of drilling to the bed of glaciers and ice sheets in less than one field season are needed to make basic measurements including temperature, heat flux, pressure, deformation, and slip, and to sample basal sediments and bedrock. The U.S. RAID drill is a step in this direction. Hot-water drills capable of accessing the bed through 500 m to 2,500 m of ice are urgently needed. Logging tools to detect temperature, diameter, inclination, azimuth, and pressure are needed in connection with the production of ice boreholes.

2. Remote sensing of basal conditions: Remote sensing such as active and passive seismic arrays and radio echo sounding complement in situ measurements of bed conditions and englacial properties. Seismic imaging requires arrays of shallow holes for emplacing sources. The capability for producing large numbers of shallow holes (25 - 100 m depth, 5 - 10 cm diameter) should be maintained within IDP. The Rapid Air Movement (RAM) drill is being refurbished and enhancements for increased portability, and increased efficiencies (drilling rate, reduced size, and power consumption) are needed to improve the agility of the RAM drill for creating shot holes.

3. Sub-ice shelf mass balance: Ice shelves buttress discharge from ice sheets and ice sheets grounded below sea level can become unstable after their buttressing ice shelves disintegrate. Recent work indicates that ocean temperatures control rates at which the ice shelves melt, and emerging observations (Jenkins et al., 2010; Stanton et al., 2013) and model results (Favier et al. 2014; Pattyn et al., 2013; Gagliardini et al., 2010; Pollard and DeConto, 2009) indicate that sub-shelf melting exerts strong control

on the mass balance of ice sheets. Although measurements near the grounding line have been made and more are being conducted, coverage is still sparse. The melting process is determined by boundary layer physics that operate on spatial scales of centimeters. Access holes large enough for deploying instruments on moorings, autonomous underwater vehicles, and remotely operated vehicles are needed to acquire short-term spatially-distributed data. Additionally, long-term observatories at targeted sites are needed to document temporal variability. All these experiments should be directly related to grounding-zone studies and linked to oceanographic campaigns beyond the ice shelves.

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Heat and mass exchange in sub-ice shelf cavities impact ice flow and ice sheet mass balance. Image credit: *WISSARD project*.

4. Grounding zone processes: Improved understanding of processes in grounding zones is needed to assess the role of fast-flowing ice streams and outlet glaciers on the stability of ice sheets. Conceptual geological models of grounding-line environments have been inferred from stratigraphic successions. Remote sensing studies using satellite observations and geophysical surveys have been conducted at grounding lines of major ice streams, and studies of processes at modern grounding lines are underway (Anandakrishnan et al., 2007; Alley et al., 2007; Horgan et al., 2013; Christianson et al., 2013). Few direct measurements or materials have been collected at grounding lines and grounding zones of fast-flowing ice streams and outlet glaciers (Begeman et al., 2018; Venturelli et al., 2020). Small diameter access holes are needed to deploy instruments to measure spatial and temporal changes in these critical areas.

5. Rheological properties of ice: Rheological properties of ice depend strongly on temperature, impurities, and texture, including grain size and fabric (Cuffey and Paterson, 2010). Improved understanding of the controls on the rheology is needed to develop realistic models of deformation of ice sheets. These models are needed to help develop depth-age relationships in ice cores, understanding flow and shear, and also to establish past, present, and future responses to possible environmental changes. Folding of deep ice and large structures forming at the base of the ice are believed to be related to the rheological structure of ice. Studies at Siple Dome (Pettit et al., 2011, Bay et al., 2001) and Dome C (Pettit et al., 2011), for example, have shown that strong vertical gradients in the effective viscosity of ice are likely present at depth in the ice sheets. These strong variations in ice rheology have the potential to lead to folding (such as at NEEM, Dahl-Jensen et al., 2013) or the formation of shear bands. Sensors that measure depth profiles of temperature, fabric, optical stratigraphy, tilt, and borehole diameter in boreholes are now available and can be calibrated against ice core properties. Rapid-access drills that can drill through ice up to 3,000 m thick are needed to deploy such sensors. In particular, the ability to drill multiple holes along a flow line can provide key spatial changes in ice properties. In addition, a system to rapidly access the ice sheet and then extract ice cores from selected depths would allow analyses of ice properties at depths of special interest; such a drill does not yet exist but should be planned.

6. Glacial history: Defining the extent and volume of ice sheets under paleoclimatic conditions warmer than the present (Eemian, Marine Isotope Stage-14, Pliocene, and mid-Holocene warm periods in Greenland) is an important indicator of future ice sheet vulnerability. Although a variety of indirect approaches have been used to constrain the history of ice sheets (glacial geology, paleoceanography,

etc.), the most direct method is to determine the age of basal ice across an ice sheet bed. Basal ice age can be modeled with age-depth flow models, or more directly by dating trapped air in basal ice. Slow-moving ice in the vicinity of ice divides contains a record of past ice dynamics (thinning and divide location). Depth profiles of age and temperature from ice cores and boreholes can be used to extract histories of accumulation and ice dynamics (Waddington et al., 2005; Price et al., 2007). Records from coastal domes and coastal ice caps are of special interest because they can be used to infer past extents of ice sheets and the history of deglaciation (Conway et al., 1999). Intermediate depth (~1,500 m) cores to measure depth-profiles of age and temperature at targeted coastal domes are needed to help constrain the deglaciation of ice sheets. Coring on ice domes near the Amundsen Sea Embayment may be able to provide a context for more recent observed changes in ice dynamics, particularly accelerated thinning in the most recent several decades.

Cosmogenic nuclides in bedrock beneath ice sheets can tell us about their former extent, and the timing and duration of past exposure periods. Techniques to estimate the size and shape of ice sheets during colder periods are well established (e.g., Mercer, 1968, Denton et al., 1989, Todd et al., 2010; Bentley et al., 2010; Stone et al., 2003; Hall et al., 2004; Anderson et al., 2014; Schaefer et al., 2016); determining their extent and thickness under warmer climates is more problematic. Much of the evidence is hidden beneath the present ice sheets. Recovery of cores of basal ice for gas analysis is very useful when obtained in conjunction with basal rock cores. Under shallow ice, nimble methods for reconnaissance recovery of short rock cores for cosmogenic nuclide techniques to quantify periods of exposure (ice free) and burial (ice cover) have been developed, for example the Agile Sub-Ice Geological Drill (ASIG) and ice-enabled Winkie drill for use near the ice margins. Under very deep ice, rapid access drilling using the RAID drill may open up new perspectives on ice-climate linkages in a warmer world.

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Caption: John Goodge and a colleague collecting specimens in the Transantarctic Mountains. Photo credit: John Goodge/University of Minnesota-Duluth.

Depth profile measurements on short (1-5 m) subglacial bedrock cores will be used to confirm that cosmogenic nuclides were produced in situ, and identify surfaces that constrain subglacial landscape evolution by subglacial erosion. Erosion reduces and ultimately erases the nuclide profile, so eroded surfaces must be avoided by targeting surfaces where ice is frozen to the bed. Note, however, that small amounts of erosion can be identified and the effects constrained using combinations of nuclides with different production profiles (Liu et al., 1994). With rapid access to subglacial bedrock in which cosmogenic nuclides can be measured, key problems can be addressed, such as the vulnerability of the West Antarctic, parts of the East Antarctic, and Greenland Ice Sheets to future climate warming, Pliocene ice-sheet collapse, and the onset of continental glaciation in Antarctica. Potential targets to address the interglacial extent of West Antarctic glaciation include Mt. Resnik, a subglacial peak which rises to within 330 m of the surface near the WAIS divide (e.g., Morse et al., 2002), ice rises particularly along the Siple Coast such as the Crary and Steershead ice rises (e.g., Scherer et al., 1998), the subglacial roots of nunataks (rocks emerging above the ice) in the Pine Island and Weddell Sea catchments, and a variety of sites in Greenland including both interior sites (e.g., Schaefer et al., 2016) and peripheral sites that border key

areas (e.g., North Greenland, NEGIS). In addition, ongoing international studies of past ice thickness variations evidenced from multiple cosmogenic nuclides on nunatak altitude transects in Dronning Maud Land could benefit from future sampling of subglacial nunatak slopes. Data from beneath high-altitude domes and plateaus in the Transantarctic Mountains could shed new light on the long-running debate over ice-sheet collapse in the Pliocene (e.g., Webb et al., 1984; Denton et al., 1993). A variety of isotopes with varying half-lives can be used to constrain long-term ice sheet stability (e.g., ^{36}Cl , ^{26}Al , ^{10}Be), and new application of in situ ^{14}C can constrain Holocene ice sheet changes. In Greenland for example, in situ ^{14}C measurements from periphery ice drilling sites would provide ice sheet models with direct measures of ice sheet presence/absence during smaller-than-present ice sheet conditions during the Holocene thermal optimum. Eventually, measurements of long-lived radionuclides such as ^{53}Mn ($t_{1/2} = 3.7$ million years) and ^{129}I (16.7 million years) paired with stable ^3He and ^{21}Ne may even provide constraints on the early Neogene onset of Antarctic glaciation, targeting samples from the subglacial Gamburtsev Mountains.

Summary

Understanding present and past behaviors of glaciers and ice sheets is essential for improving predictions of future behavior of ice sheets and sea level. Improved understanding requires fast-access drilling, such as those from the IDP Agile Sub-Ice Geological Drill, the IDP ice-enabled Winkie drill, or the Rapid Access Ice Drill to enable fundamental measurements of: (i) physical conditions, including geothermal flux, and processes at the beds of glaciers and ice sheets; (ii) physical properties of the ice that affect ice flow and folding, (iii) physical processes at grounding lines and grounding zones of fast-moving ice streams and outlet and tidewater glaciers; (iv) ice-ocean interactions at grounding lines. Past responses of glaciers and ice sheets to climate and sea level change also offer clues to future possible responses. Depth profiles of age and temperature from ice cores can be used to reconstruct past thickness and extent of ice sheets as well as climate. Basal ice cores in condition suitable for trapped gas analysis provide dating of the basal ice. Intermediate depth (~1,000 m) cores at targeted coastal domes are needed to constrain the extent and timing of deglaciation. For reconnaissance of sites for bedrock coring under shallow ice, the use of an very agile ice coring drill (e.g., Foro or Eclipse drill) along with a method of sampling rock at the ice-rock interface using the same or slightly modified drilling apparatus, would provide a logistically agile way of site selection for subsequent rock drilling of meters-long rock cores. The collection of basal ice cores for trapped gas analysis, and subglacial bedrock for cosmogenic nuclides from both strategic periphery sites and also from the ice sheet interior can provide direct constraints on past ice sheet history.

Table 2. Requirements of drills needed for studies of ice dynamics and glacial history. The IDP Long Range Drilling Technology Plan discusses drills available for retrieving cores or creating access holes in ice sheets.

	Diam (cm)	Ice Depth (km)	Core or hole	Ambie nt temp (°C)	Clean access?	Transport type	Site occupanc y	Int'l Aspects	Drill Name
Bed conditions	8	1-3	Hole	-50	maybe	Twin otter/ helo/lt traverse/Herc	<4 weeks	US & others	CHWD
Geothermal flux	5-8	1-3	Hole	-50	no	Twin otter/ helo/lt traverse/Herc	<4 weeks	US & others	RAID
Geologic coring for cosmogenic samples and for basal ice samples	6-10	0.1- 2.5	Basal ice core, Rock core	-50	no	Helo sling load/ Baseler/ traverse	4-8 weeks	US	Winkie/ ASIG/ RAID
Nimble geologic coring and basal ice coring under shallow ice	3-5	<.5	Basal ice core, Rock core	-30	no	Twin otter/helo/ lt traverse	<4 weeks	US	Winkie/ ASIG
Rheological properties	8	<3k	Hole	-40	no	Herc/ traverse	<4 weeks	US & others	RAID/ CHWD
Internal layering	8-10	<3k	Hole	-40	no	Herc/ traverse	<4 weeks	US & others	RAID/ CHWD
Sub-ice shelf/ice stream instrumentation	10- 25	<1k	Hole	-30	shelf- no; stream- yes	Twin otter/ helo/ herc/ traverse	2 weeks	US & others	CHWD/ SchWD
Ice shelf ROV deployment	100	<1k	Hole	-30	no	Twin otter/ helo/Herc/ traverse	2-4 weeks	US & others	CHWD/ SchWD
Grounding zone	8-75	<1k	Hole	-30	no	Helo/Herc/ traverse	2 weeks	US	CHWD/ SchWD
Seismic imaging	5-10	~100 m	Hole	-40	no	Twin otter/ helo	Hours/da ys	US	RAM/ SHWD

Note: CHWG is the UNL Clean Hot Water Drill

III. Subglacial Geology, Sediments, and Ecosystems

Bedrock, sediments, and ecosystems existing within and beneath ice sheets remain largely unexplored because of the lack of rapid access. Rapid access to subglacial environments is needed to address a wide range of science questions. Specifically:

1. Bedrock geology: The Antarctic continent and its lithospheric plate, play important but poorly understood roles in global tectonic architecture, leading to contradictory hypotheses. Antarctica is considered aseismic, but if so, it would be unique among all of the continents. Its plate is surrounded by mid-ocean-ridges, and hence should be under compression, yet there are active extensional regimes. The West Antarctic Rift System is one of the largest on Earth, and currently known attributes are unique, by having only one rift shoulder and by being largely below sea level. Fundamental questions about the Antarctic Ice Sheet persist. What is the origin of the enigmatic Gamburtsev Subglacial Mountains and how have they influenced the overlying ice sheet? What are the composition, geothermal heat flux, and geotectonic histories of East Antarctica, and how does it influence ice-sheet behavior? What were the dominant factors controlling the spatial extent and temporal variability of ice sheets during warm climate periods in the past? What is the role and history of subglacial sediments in the interior? What are the physical conditions at the base of the East Antarctic ice sheet? Constraints on composition and age of basement rocks of interior East Antarctica would place better constraints on Precambrian provinces and evolution of the Antarctic shield for verifying current models. The state of stress in basement rocks is required for evaluating seismicity and extensional regimes. Boreholes through the ice into crustal rocks are needed to conduct passive and active seismic experiments for delineating crustal structure. Continental topography is a significant control on glaciation; rising mountains and higher elevations focus snow accumulation and become nivation centers for ice sheets. Sampling bedrock to determine its age and constrain its cooling history using thermochronology is important for supercontinent reconstruction, understanding the tectonic history of the continent as well as reconstructing paleotopography for glaciological modeling of Antarctic Ice Sheet history. Access boreholes to the ice sheet bed are required to recover short rock and sediment cores for these studies. Locations should be based on best estimates of bedrock geology, bed paleotopography, and plausible ice sheet extents based on models. In Greenland, the ice sheet has waxed and waned during the past 2.5 million years. Erosion of mountains and ice sheet modeling has simulated past changes, but access to old ice and basal rocks/material is needed for verification and full understanding.

2. Subglacial basins and sedimentary records: The records of glaciation and their variations in Antarctica are found in scattered terrestrial deposits and sedimentary basins and can be compared with offshore records that have been collected near the margins. Interior subglacial basins also likely contain proxy records of paleoclimate and ice sheet history to complement these records from the continental margins. Four main categories of sedimentary targets are: subglacial lakes, ice rises, West Antarctic sedimentary basins, and East Antarctica basins. Each category may have a variety of origins and histories because of

differing locations relative to the ice sheet margin and magnitudes of past ice sheet fluctuations. Thus, they may provide valuable archives of paleo-ice sheet and paleoclimatic changes. Subglacial lakes occur throughout the continent, the largest being subglacial Lake Vostok. Subglacial lakes contain sedimentary records; sediments have been collected at Lake Ellsworth and the Whillans Subglacial Lake (Hodson et al, 2016).

- Photo here -

Caption: Illustration showing the aquatic system that scientists believe is buried beneath the Antarctic ice sheet. Photo credit: National Science Foundation, Photo Gallery.

Subglacial ice rises can cause locally grounded “pinning points” that play an important role in buttressing the discharge of streaming ice from the ice sheet. Recovery of these sediments will provide Neogene and Quaternary paleo-environmental archives, but may also provide insights on till deformation processes downstream of the Whillans and Kamb ice streams (e.g., Scherer et al., 1988). Shallow drilling of ice rises and acquisition of oversnow seismic reflection profiles radiating away from core sites will allow the deeper geometry of the strata to be evaluated for locating deeper drilling and recovery of long, continuous records in adjacent marine basins. In West Antarctica, the stratigraphic record in various basins and probable rifted grabens may contain a mid-late Mesozoic and Cenozoic history of West Antarctic evolution and paleoclimate history. Two low regions within the Wilkes Land sector of East Antarctica (Aurora and Wilkes Subglacial Basins) appear as broad down-warped basins filled by marine and non-marine strata. They may contain evidence of the much debated past dynamics and paleoclimate of the East Antarctic Ice Sheet. Recently, Mengel and Levermann (2014) suggested that only a narrow, low coastal rim holds the portion of the East Antarctic ice sheet overlaying the Wilkes Subglacial Basin back, raising concern about ice sheet stability.

Access holes are also needed to recover longer sedimentary cores comparable to those from the continental margins. Also, the basins on the interior of the Transantarctic Mountains may be sites for good proxy records of past ice sheet dynamics. These are also excellent sites to measure geothermal heat flux to help constrain ice sheet bed conditions.

3. Sub-ice microbial ecosystems and biogeochemistry: Aqueous and sedimentary subglacial environments in Antarctica and Greenland are inhabited by microorganisms (e.g. Christner et al., 2014; Dubnick et al., 2017) and are a potentially large planetary reservoir of microbes and (microbially derived) organic carbon, perhaps of the same magnitude as that in the surface oceans. Modeling and direct measurements (Wadham and others 2012; Michaud and others, 2017) suggests these environments could contain large volumes of the greenhouse gas, methane, which could impact atmospheric methane concentrations in response to rapid deglaciation. It has also been hypothesized that the flux of dissolved elements and sediments in subglacial waters can enhance primary productivity in the marine environments that they drain into (Hawkings et al., 2020; Vick-Majors et al., 2020). Elucidating the spatial and temporal distribution and dynamics of these aqueous environments, including their physical and chemical properties (such as temperature, salinity, and pressure) and associated biogeochemical

processes (i.e. microbial communities and organic carbon material fluxes) is key to understanding ice sheet stability and the role of large continental ice sheets in global biogeochemical cycles. The rapid changes anticipated in the size of polar ice sheets may trigger significant reorganization of subglacial hydrologic conditions, which may feed back into acceleration of ice sheet retreat and may force adaptation of subglacial biota, whose metabolism is linked to subglacial hydrology (Vick-Majors et al., 2016), to rapidly changing conditions.

The long timescale of microbial entrapment in sub-ice environments relative to the lifetimes of microbial cells provides an opportunity to explore questions concerning rates of evolution, and constraints on biodiversity. Microbial cells and their genomic material should also provide valuable information that can be linked to paleoclimatic change; such life forms may be the only biological survivors in areas covered by glaciations for millions of years. Icy systems on Earth also may provide crucial terrestrial analogs for extraterrestrial life surviving and persisting on icy planetary bodies in our solar system, such as Mars, Europa, or Enceladus. Exploration of life within subglacial lakes and their sediment has begun; the first reports on the microbiology of Subglacial Lake Whillans have been published (Christner et al, 2014) and a second subglacial lake (Subglacial Lake Mercer) was accessed in 2019. Of particular interest is the distribution and ecological function of the resident microbes, the extent to which biogeochemical weathering occurs, and the genetic diversity of microbial communities in subglacial lakes and sediments. Furthermore, the forward motion of thick layers of water-saturated till beneath fast-flowing ice streams may provide a pathway for transportation of subglacial biological and diagenetic materials and weathering products to the surrounding ocean, as does the movement of debris-rich basal ice. Some subglacial meltwater is also transported over long distances within basal drainage systems, which again, likely discharge subglacial microbes and their metabolic products into circum-Antarctic seawater. Access holes through the ice and the acquisition of basal ice cores are needed for this science, and, for scientific and environmental integrity, these studies must be conducted with clean technology both during access and sample acquisition. This science is at an early stage, and it is best to conduct studies first at sites where the ice is not thick and logistics issues can be readily addressed.

- Photo here -

Caption: Microbial ecosystems have been found under the West Antarctic Ice sheet (Christner et al, 2014). This photo shows a coccoid-shaped microbe with an attached sediment particle from subglacial Lake Whillans. Photo credit: *Trista Vick-Majors*.

4. Subglacial lakes and hydrological systems: Subglacial hydrodynamics are an important yet poorly understood factor in ice sheet dynamics in both Antarctica and Greenland. The volume and distribution of water exert strong influences on the resistance of the bed to ice flow and therefore, is an important control over ice velocities. More than 400 subglacial lakes have been discovered in Antarctica. Measurements to quantify present-day lakes and subglacial hydrological systems are important for understanding ice dynamics, weathering and erosion of subglacial rock, sediment transport and jökulhlaup events, microbial ecosystems, and maintaining systems of subglacial lakes. Of particular interest is to establish the diversity of life in subglacial lakes, the degree of hydrological interconnectivity between lakes and the Southern Ocean, and their influence on the rest of the subglacial hydrological

system. The lakes also house sedimentary evidence of ice sheet and geological histories as well as climate change. Access holes and the ability to collect samples of water and sediments are necessary to understand these systems. In addition, data from Subglacial Lake Whillans suggests that in active hydrological systems, water geochemistry, microbial life, and hydrology are intimately connected (Vick-Majors et al., 2016). Understanding the temporal dynamics and geochemical ramifications of subglacial processes requires installation of sensor strings capable of collecting subglacial hydrological data including dissolved oxygen concentrations, current velocity and direction, and salinity.

Russian drillers accessed Subglacial Lake Vostok during the 2011-12 season, and then during 2012-13 successfully recovered an ice core (~30 m) of the frozen lake water that entered the borehole the year before. The British attempted to access subglacial Lake Ellsworth in the interior of West Antarctica in 2012-13 but unfortunately were stopped due to operational problems during drilling. The U.S. successfully penetrated and sampled subglacial Lake Whillans upstream from the Siple Coast grounding line during the 2012-13 season. The new drill built for drilling Lake Whillans includes a filtration unit and UV-treatment system to decrease contaminants in the drilling water and provide clean access to the subglacial environment (Priscu et al. 2013) and was used again to access Subglacial Lake Mercer in 2019. The filtration technology was successful at reducing microbial bioload in the drilling fluid as per the Antarctic Treaty Code of Conduct.

Summary

Subglacial environments contain biologic, climatic, geologic, and glaciologic materials and information, much of which cannot be obtained elsewhere. Drills to create access holes are urgently needed to sample basal ice, subglacial water and sediments, and bedrock. Hole diameter requirements vary depending on instrumentation needed; clean technology is required (NRC, 2007), as is strict environmental review where the bed is wet, except for ice shelves and grounding zones at the end of drainage basins. Successful sampling will require that access holes receive regular maintenance, allowing the holes to remain open for several days. Differential ice motion may be a complicating factor, especially if the ice sheet is sliding at the bed. A conceptual design is also needed for a drill that can provide clean access large enough to deploy subglacial rovers; this design should strive to minimize supporting logistical requirements.

The desired characteristics of the drills needed to create clean access holes for the science of the sub-ice environment are provided in table 3 below. For accessing sensitive targets such as subglacial lakes, hot water drills should have temperature and depth sensors and “smart” drill heads; this technology needs to be developed. Other subglacial access requirements are also covered above in Table 2.

Table 3. Requirements of drills needed for studies of subglacial geology, sediments, and ecosystems.
The IDP Long Range Drilling Technology Plan discusses hot water and mechanical rapid-access drills that could provide clean access holes for the projects described above. Clean mechanical rapid-access drills do not currently exist; conceptual and engineering development is needed.

	Diam. (cm)	Depth (km)	Core or hole	Ambient temp (°C)	Transport type	Site occupa ncy	Int'l aspects	Environ restrictions	Drill Name
Sediments/i ce sheet dynamics (Wet bed)	10-25	0.2-3	Hole, sediment core	-50	Helo/Twin Otter/travers e/Herc	weeks	U.S. & others	Clean access	CHWD/ SchWD
Biogeochem (Wet bed)	3-25	<3	Hole, sediment/r ock, basal ice core	-50	Helo/Twin Otter/travers e/Herc	weeks	U.S. & others	Clean access	CHWD/ SchWD
Bedrock geology/ Tectonics (Frozen bed)	5-10	1-3	Icehole, rock core	-50	Herc/travers e	4-8 weeks	U.S.	None (dry bed only)	RAID/ ASIG
Geology/ ice sheet history (Wet bed)	5-20	<4k	Hole, rock core	-50	Herc/travers e	weeks	U.S. & others	Clean access	-
Deep Subglacial lake biogeochem (Wet bed)	50-100	3-4k	Hole, sediment, basal ice core	-50	Herc/TwinOt ter /traverse	4-8 weeks	U.S. & others	Clean access	CHWD

IV. Ice as a Scientific Observatory

Polar ice sheets and mid-latitude ice caps archive evidence of past climate and ice dynamics and also serve a variety of endeavors that use the ice as a platform for science. Borehole access to the interior of the ice sheet enables wide-ranging observations, from glaciology, climatology, and planetary science to experimental astroparticle physics.

1. Borehole logging for past climate and ice dynamics: Borehole logging of both fast-access holes and boreholes originally drilled for ice cores greatly enhance evidence of climate and ice dynamics preserved in the ice. These analyses are difficult or impossible to obtain by other methods, and complement observations from ice cores and remote sensing platforms. Borehole logging is nondestructive, continuous, and immune to core damage or drill depth errors and permits study of a large volume of ice in situ. Ice sheet boreholes serve as enduring scientific observatories. For example, borehole paleothermometry probes provide the most direct measurement of temperature histories and can be used to calibrate other paleoclimatic indicators. Optical borehole probes can rapidly obtain stratigraphic records, which are more coherent and detailed than can be reconstructed from core measurements. Borehole sonic loggers can provide continuous records of ice fabric that are difficult or impractical to obtain using thin sections of core. Repeated measurements of fabric, tilt, and hole deformation improve modeling of ice sheet behavior and stability over time as an ice sheet flows over uneven terrain. Logging multiple nearby rapid access holes permits advanced studies of climate history and ice flow.

1.1 Winches: Winch platforms that can support borehole-logging projects are important community resources. IDP has three winches in inventory, one for intermediate depth (1.5 km) and two for deep (4 km) applications. IDP has adopted a standard wireline for all community winches, a 3/16" four-conductor armored oil-patch cable with a 1" Gearhart-Owen cable head. IDP has also established a policy of deploying a trained operator to the field along with the IDP winches, particularly the deep winches. Although this cost is not directly reflected in proposal budgets, a cost estimate is included with each proposal requiring IDP resources, for NSF budgeting purposes, as is the case with ice and rock coring drills. In certain cases, the PI or members of the PI's team may be trained and certified to operate the winches, particularly the intermediate depth winch.

- Photo here -

Caption: Ryan Bay and Elizabeth Morton deploy borehole-logging instruments at Siple Dome. Photo credit: *Joseph Talghader*.

Pre-deployment winch telemetry testing of all logging tools is essential for successful fieldwork. Whenever possible, logging tools should be tested over the winch that will be used in the field. In some cases IDP leaves winches deployed to save logistical cost and effort, and tools must instead be tested on winch-cable systems that are electrically similar.

Pressure testing of new borehole tools prior to deployment is performed at an IDP facility in Madison, WI. IDP maintains a pressure chamber for testing tools up to pressures of 6 kpsi. The chamber is approximately a 10-foot cylinder with an inside diameter of 10 inches. Pressure testing is especially

important with Estisol drill fluid, since it is more aggressive than other drill liquids and even small leaks may damage internal components.

1.2 Borehole preservation: Where practical, drilling practices and materials should be chosen to produce and maintain clean uniform boreholes, and to keep the boreholes accessible. Anticipated failure modes of glacial boreholes include:

- “Natural” end-of-life borehole collapse: Depending on the strain regime, complete collapse of even a borehole fully compensated with fluid occurs over years and is largely unavoidable.
- Borehole collapse due to removal or failure of borehole casing: Premature collapse can be avoided by leaving the casing in place, proper casing design, and maintenance.
- Borehole burial: Burial of borehole casing by snow accumulation.
- Ice plug: An ice plug can form at the fluid level when a partial casing failure permits snow and ice to accumulate in the well.

Over time borehole-drilling fluids can become turbid, degrading optical measurements. Best practices should include avoiding the introduction of substances such as heavy greases in the borehole, and materials that can be dissolved by solvents used as drill fluids. IDP also provides towers and sheave wheels needed for borehole access. If requested and resourced by NSF, IDP could preserve, maintain, extend borehole casings, and maintain the proper level of drill fluid compensation for existing boreholes. The Borehole Logging Working group will work with the ice borehole logging community to prioritize the boreholes requiring preservation.

1.3 Recent logging projects: WAIS Divide: Several groups have logged the WAIS Divide ice core borehole in the 2014 to 2017 time frame. Measurements included temperature, optical, and seismic profiling, and an acoustic caliper along with a kHz acoustic fabric logger for the kHz range. WAIS Divide drilling included five replicate coring deviations and this logging activity was the first to be done in a borehole with deviation channels. The Replicate Coring System for the DISC drill was designed in order to make all deviations on the uphill side of the main borehole, so that logging tools naturally follow gravity and remain within the parent channel. The deviations did affect logging data and some issues were encountered while passing the deviations, in particular the acoustic logging tool was diverted into the side channel at the deviation with the most borehole damage near 3,000 m, but accessed the main borehole on a following attempt. Three other logging tools followed the main borehole and passed all deviations without incident. The deviation drilling and subsequent borehole logging at WAIS Divide was largely successful.

South Pole Ice (SPICE) core: The Foro 1650 (aka Intermediate-Depth Drill) was deployed to the South Pole for the 2014-15 and 2015-16 field seasons, successfully collecting 1,751 m of ice core. Analysis of the SPICE core may take advantage of and supplement the wealth of existing South Pole data from shallow cores, snow pits, IceCube hot-water boreholes and meteorological observations. The SPICE Core borehole is pressure compensated by Estisol-140, which has caused convective problems in temperature logging because of its high viscosity. Estisol-140 has also exhibited a tendency to cloud, which could affect optical logging. The SPICE core project is a benefit to ongoing South Pole in-ice particle physics projects, by

providing ground truth measurements of ice chemistry, fabric, and particulates for characterization of optical, radio, and acoustic signal propagation. Due to the proximity of the drill site to the IceCube and ARA arrays, the borehole continues to serve as an access point for calibration of existing and future South Pole in-ice physics and astrophysics experiments.

RAID: The RAID (Rapid Access Ice Drill) is a drill in development that will be capable of penetrating a 3,000 m ice sheet, and coring small samples of ice and subglacial basal rock. RAID is expected to produce five (5) boreholes every season, and these boreholes will potentially serve as scientific observatories for the study of ice and climate. RAID will require a dedicated logging winch integrated with the drilling platform, capable of reaching 3,000 m for logging immediately following the drill. The system could partly serve as a hole qualifier for evaluating the performance of the drill during development. Measurement of pressure will ensure that the borehole is properly compensated and optical dust logging will provide immediate verification of the depth-age model. Additional measurements could include temperature, diameter, borehole inclination/trajectory, and a camera. It would be desirable to rapidly log temperature and borehole diameter immediately after drilling, possibly at the same time as the optical dating. These preliminary readings could form baselines for subsequent measurements and time evolution studies. Infrastructure will be needed to manage future borehole logging projects that will make use of RAID boreholes.

RAID borehole preservation: RAID has the potential to create many deep boreholes over a number of years. Preserving every RAID borehole indefinitely is impractical. The RAID project, with borehole logging scientists, will need to determine the scope of preservation efforts. The number of holes to preserve, the priority of holes and the duration of the effort will need to be weighed against cost and logistics.

Borehole preservation effort could be separated into short-term (<5 years) and long-term time horizons. Preservation of each RAID borehole for 3 - 5 years will allow for repeat measurements, particularly in studies of borehole temperature and deformation. Uncased and under-balanced boreholes could be of interest for deformation studies, although removal of the casing and fluid head will limit the lifetime of the borehole to a few years.

RAID may also select a subset of holes for long-term preservation, to serve as observatories and to allow for future technology developments. Preservation would require leaving a sturdy casing in place, maintaining, and periodically extending the casing above the snow surface, as well as removal of ice plugs when necessary. Holes near ice divides could be kept open for decades in principle. In off-axis zones, shearing could severely limit borehole lifetime and closure may occur at discrete depths. In higher accumulation areas, it may be possible to use an extended casing supported by a lightweight tower to relieve maintenance effort. Qualifying tools (borehole diameter, inclination/trajectory, camera) could be useful for assessing borehole condition prior to fielding a more substantial logging mission. Holes selected for long-term preservation would likely be chosen to form a geographically diverse set.

1.4 Borehole qualifying: IDP does not currently maintain logging tools for verifying borehole parameters such as inclination, diameter, depth, roundness, temperature, etc. There is growing consensus in the

logging community that IDP should develop this capability. A hole qualifying system could be deployed each season as a hole is drilled or upon hole completion. The information provided by such a logging system could be crucially important for drillers, particularly for drills with little or no down-hole sensing capacity, such as the Intermediate Depth Drill or the RAID. These logging measurements could also provide a baseline for longer-term borehole deformation studies.

1.5 Borehole Allocation Committee: The IDP Borehole Logging Working Group (BLWG) is currently exploring formation of a special committee to advise IDP on management of community resources as the logging community continues to grow. These resources include winch and winch operators, logging tools and accessories, and borehole time. Pre-deployment reviews of logging projects, with participation by IDP engineers, will ensure that new tools are safe and ready to deploy.

2. Ice as platform for physics and astrophysics: Efforts are under way to use glacial ice as a platform for study of fundamental physics and astrophysics. These experiments make use of polar ice as an abundant, clean, stable, low-background and transparent (to radio and optical waves) detection medium for observation of sub-atomic particle interactions. For example, the now completed IceCube telescope uses ice at South Pole to detect high-energy neutrinos traveling to Earth from cosmic sources. The Enhanced Hot Water Drill (EHWD) developed for IceCube is a powerful and fast access drill capable of creating 2,500 m deep, half-meter diameter boreholes at a rate of about three per week.

IceCube-Gen2 is a proposed facility for future Antarctic neutrino astronomy. IceCube-Gen2 will aim to increase the effective volume of IceCube by an order of magnitude, while only doubling the amount of in-ice instrumentation. The IceCube inter-string spacing of 125 m would be increased to 250 - 300 m, taking advantage of the long absorption lengths of optical photons in Antarctic ice, particularly South Pole ice from the early stages of the Last Glacial Period. This expanded array would improve detection capability in the PeV energy range and provide high statistics samples of extraterrestrial neutrinos, for better characterization of source distribution, spectrum and flavor composition. IceCube-Gen2 will require improvements to the EHWD, including a more mobile and efficient hot water plant, and a modular sled-mounted drill system, which is less complex and requires a smaller operations crew.

- Photo here -

Caption: A Digital Optical Module (DOM) is lowered into a hole in the ice at Amundsen-Scott South Pole Station as part of the IceCube project. IceCube will search for neutrinos from distant astrophysical sources. Photo credit: *Ethan Dicks, National Science Foundation*.

The proposed low-energy sub-array physics experiments such as PINGU (Precision IceCube Next Generation Upgrade) would be embedded within the IceCube array, in order to use the existing detector as an active shield. PINGU objectives include the study of neutrino oscillations and mass hierarchy, dark matter, supernovae, and neutrino tomography of Earth's core. PINGU will deploy a relatively high density of photocathode (light sensors) in a small ice volume, requiring hot-water drills capable of making deep access holes at small spacing. The currently proposed PINGU geometry will have inter-string spacing of

~20 m and a three to five meters vertical spacing between sensors. These projects will enable research and development on the next generation of low-light photodetectors and the optical properties of in situ ice over short distance scales. Hot-water drill upgrades are aimed at improving the optical clarity of the refrozen water column, including filtration of large-particle impurities and degassing to avoid bubble formation.

Experiments to detect extremely high-energy neutrinos will make use of large areas of the polar ice sheet. The ARA experiment (Askaryan Radio Array), in early development at South Pole, is planning to instrument on the order of 100 km² of ice with radio antennas to detect radio pulses from so-called Greisen-Zatsepin-Kuzmin-scale (GZK) neutrinos. The Radio Neutrino Observatory (RNO) at the South Pole is a proposed, distributed, detector array for measuring the highest energy neutrinos. Detector stations are spaced 1.25 km apart in rectangular grid spaced away from the South Pole station along a power and communications backbone. Drilling would consist of 4-5 holes per station, drilled down to a nominal depth of 60 m, with the possibility of 100 m evaluation holes. Drilling would be conducted with the ASIG auger drill setup on a sled along with integrated weather sheltering. Holes are used for radio antennas to detect the radio pulses from high energy neutrinos. Field work over four Austral Summer seasons has been proposed.

Due to its proximity to the IceCube and ARA detectors, the SPICE core borehole could serve as an access point for calibration beacons or standard candles, as part of the South Pole facility and infrastructure. These beacons could be operated at multiple depths and hence different ice temperatures, densities, fabrics and impurity levels. These unique measurements would have implications for radio and optical detection of high-energy neutrinos and also provide opportunities for basic glaciology research. Radio-illuminating beacons could provide signals in the 100 - 1,000 MHz frequency range out to a radius of 20 km, thus permitting studies of neutrino detection over areas up to 1,000 km², and also help in understanding anomalous features seen in ice-penetrating radar surveys.

The ARIANNA experiment (Antarctic Ross Ice shelf Antenna Neutrino Array) proposes to deploy a large array of surface radio antennas on the Ross Ice Shelf to observe cosmogenic GZK neutrinos. Log-periodic dipole antennas will be buried in pits approximately 6' x 6' x 0.5' and controlled by solar-powered relay stations. An effective and efficient means for digging and backfilling many such pits will need to be developed. The ARIANNA also anticipates taking two shallow (~100 m), 4"-6" cores from nearby for study of the firn-to-ice transition and for borehole-to-borehole radio tomography, as well as drilling one deep (~500 m), 4"-6" borehole.

3. Seismic studies: The Global Seismographic Network includes seismic monitoring stations for earthquakes and other events such as emissions from calving and sliding glaciers and ice sheets. The South Pole Remote Earth Science and Seismological Observatory has seismic equipment installed ~300 m deep within boreholes. A similar observation network is planned for Greenland.

4. Ice sheet as an archive of recent past atmospheric composition: In the very cold areas of ice sheets where snow rarely melts, many decades of snowfall create a porous network of firn in the top many tens of meters of the ice sheet. The firn serves as an archive of atmospheric composition, with the oldest air

existing at depth. Sampling firn air from various depths within boreholes drilled in the ice sheet enables, for example, observation of the extent of anthropogenic emissions and patterns of increase or decrease.

5. Exploration of basal ice formation processes: Radar imaging of basal conditions under the Antarctic and Greenland ice sheets reveals structures that have been proposed to result from accretion ice grown onto the base of the ice sheet. In order to acquire the ice to test this hypothesis, drilling at sites in Greenland, or near Dome A in East Antarctica, could access these ice features with the 1,500 m Intermediate Depth Drill.

6. Meteorite collection: Glaciers and ice sheets are sites for efficient collection of meteorites and micrometeorites. Micrometeorites yield clues to the birth and evolution of the solar system. Some are visible to the human eye on the surface of some blue ice areas, while others may be swept up inside melted water wells created in the ice at established field stations.

Summary

Ice sheets serve as a platform for a wide range of observations spanning many areas of science. In some areas, for example, firn-air studies and seismic monitoring proven-drills already exist for making the necessary access holes. Dedicated hot water drills have proven to be effective in creating deep boreholes in rapid succession. Other areas are at an early stage and will require further development of RAM drills or reverse circulation drills. A rapid access drill, with the capability to bore through several kilometers of ice to retrieve rock cores is in development. The borehole logging community is a strong proponent for repairing and maintaining boreholes at Greenland Ice Sheet Project 2 (GISP2) site at Summit, Siple Dome, and other sites. Identifying which boreholes need maintenance, prioritizing those with highest scientific value for future logging, and determining methods of repair are activities that need urgent attention. The IDP Borehole Logging Working Group will prepare a list of boreholes in the U.S. program and will work with the community to create a prioritized list for maintenance and repair.

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1313 **Science Planning Matrices**

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1315 Goals to advance the frontiers of the science in ways that enable evidence-based decision making and
1316 that inspire the next generation of scientists are described in the sections above. Community planning
1317 for the execution of the science is important for providing coordinated scientific investigations, and also
1318 for planning the associated logistical and funding requirements. For each area described above, matrices
1319 below identify the current plans for timing of the field research. In cases where new technologies are
1320 needed, a timeline for the development of technologies is provided. Black lettering in a matrix indicates
1321 projects that are currently funded, and blue lettering indicates those in the planning phase.

1322

1323 In tables 4 – 7 below the letters denoting specific drills to be used are: A: Agile sub-ice geological drill; b:
1324 Badger-eclipse; B: Blue ice drill; D: DISC drill; f: Foro 400 drill; F: Foro3000 drill; I: Intermediate depth drill
1325 (Foro1650); L: Borehole logging; It: Logging tower; R: RAID drill; r RAM drill; Sc: Scalable hot water drill; S:
1326 Stamphli drill; T: Thermomechanical drill; U: UNL CHWD drill; W: Winkie drill; 4: 4” drill; 7: 700 m coring
1327 drill, x: hand auger, sidewinder, or prairie dog.

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Table 4: Past Climate Change Planning Matrix 2021-2031

	2021				2022				2023				2024				2025				2026				2027				2028				2029				2030				2031			
Past Climate	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4				
Industrial period and glaciology																																												
Hand auger drilling projects ¹		x	x		x	x			x	x			x	x			x	x																										
SE Greenland mass balance ²		x	x																																									
Greenland climate & human history ³		f	f																																									
Mt Waddington Canada firn aquifer ⁴					T	T																																						
GreenTracs2 ¹²					f	f			f	f																																		
Taylor Dome firnification ⁵								b	b																																			
Pre-industrial baseline & dynamics																																												
Eclipse Icefield Canada ⁶								7	7																																			
Greenland coastal ice caps ²								x	x			x	x																															
Andes ⁷					T	T																																						
Amundsen Sea coastal dome ⁴												x	x																															
Dome C - past cosmic ray flux ⁹											x	x																																
Detroit Plateau Ant. Peninsula ¹⁰																																												
Large scale gobal climate change																																												
GISP2.1 Central Greenland (near Summit) ¹¹																																												
Intermediate depth coring to 1650 m																					I	I				I	I																	
Borehole logging																											L		L															
NW Greenland ¹²																																												
Prudhoe Dome 700 m core					7	7																																						
Borehole logging Prudhoe Dome										L			L																															
South Dome Greenland ¹⁷																																												
South Dome 700 m core																																												
Borehole logging South Dome																																										L	L	
Hercules Dome ¹³																																												
Ice coring at Herc Dome														F	F		F	F		F	F				F	F																		
Borehole logging at Herc Dome																											L		L	L														
2 Ma ice & last interglacial																																												
Allan Hills (4" & BID) ¹⁴					B																																							
Blue Ice Paleo Ice Project ¹⁰																																												
Blue Ice tbd COLDEX (BID & 4") ¹⁵								B	B		B	B		I	I																													
U.S. Oldest Ice deep core																																												

Point of Contact for projects in Table 4: ¹Das, Holland; ²Das; ³McConnell; ⁴Neff; ⁵Keegan; ⁶Kreutz; ⁷Meyewski; ⁹Petrenko; ¹⁰Kurbatov; ¹¹Winski; ¹²Osterberg; ¹³Steig; ¹⁴Higgins; ¹⁵Brook.

Table 5: Ice Dynamics and Glacial History Planning Matrix 2021-2031

	2021				2022				2023				2024				2025				2026				2027				2028				2029				2030				2031			
	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4				
Ice Dynamics & Glacial History																																												
Ice Dynamics																																												
Thwaites-MELT ¹					x	x																																						
Flask glacier melt dynamics ²					x	x			x	x																																		
Greenland ice sheet acoustics ³								x	x																																			
Borehole logging of RAID holes ¹⁰																					L	L			L	L			L	L														
Glacial history																																												
NW Greenland ⁴								A	A				A	A																														
Mt Woollard ⁵										A	A																																	
Thwaites grounding line ⁶	w																																											
WAIS interglacial ⁷					w	w			w	w																																		
Thwaites - seismic sounding ⁸									r	r			r	r																														
Mt Waesche unconformities ⁹					b	b																																						
Taylor Glacier Antarctica ¹¹																					Sc	Sc																						
Northern Victoria Land ¹²												w	w			A	A																											
Continental RAID drilling Antarctica ¹³																					R	R			R	R			R	R														

Point of Contact for projects in Table 5: ¹Holland; ²Kingslake; ³Lipofsky; ⁴Schaeffer; ⁵Stone; ⁶Goehring; ⁷Mitrovika; ⁸Anandakrsihnan; ⁹Campbell; ¹⁰Pettit; ¹¹Mikucki; ¹²Balco; ¹³Goodge.

Table 6: Subglacial Geology, Sediments and Ecosystems Planning Matrix 2021-2031

	2021				2022				2023				2024				2025				2026				2027				2028				2029				2030				2031			
	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4				
Subglacial Geology, Sediments, & Ecosystems																																												
Bedrock geology																																												
Continental RAID drilling Antarctica ¹																																												
WAIS ²																																												
Thwaites ³																																												
NW Greenland ⁴																																												
Subglacial hydrology & sediment dynamics																																												
Taylor Glacier Antarctica ⁵																																												
Mount Resnick Antarctica ⁶																																												
West Antarctica / Siple Coast ⁷																																												
Microbial ecosystems & biogeochem																																												
Taylor Glacier Antarctica ⁵																																												
Mount Resnick Antarctica ⁶																																												
West Antarctica / Siple Coast ⁷																																												

Point of Contact for projects in Table 6: ¹Goodge; ²Mitrovika; ³Goehring; ⁴Schaefer; ⁵Goehring/Mikucki; ⁶Mikucki; ⁷Vick-Majors.

Table 7: Ice as a Scientific Observatory Planning Matrix 2021-2031

Note: These projects either already have holes in place, or else are in need of holes but the drill has not yet been identified.

	2021				2022				2023				2024				2025				2026				2027				2028				2029				2030				2030				2031			
	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4	1	2	3	4				
<u>Ice as a Scientific Observatory</u>																																																
Neutrino detection & seismic network	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x		
South Pole Global Seismic Network	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x	x		

Associated logistical challenges

In addition to planning the science and associated drilling technology, logistical challenges impact the timing and possibilities of the field science. Challenges to conducting the field activities include:

- There is an especially urgent need to increase tractors and sleds for scientific traverses in both Greenland and Antarctica in order to address the priority science goals identified in science community consensus documents, including the IDP Long Range Science Plan 2021-2031 (this document). In addition, the 2015 National Academies report “A Strategic Vision for NSF Investments in Antarctic and Southern Ocean Research” articulated the need for improved ground-based access to the deep field in Antarctica. Access to sites on the Antarctic and Greenland Ice Sheets has become very restricted due to limited air support and aging scientific traverse vehicles. With multiple science communities requesting flights or traverses to support their science on the ice sheet, access to the deep field is now a limiting factor in executing new scientific endeavors.
- The NSF Ice Core Facility in Denver is the key location for processing and archiving of U.S. ice cores. Although some infrastructure upgrades and improvements have been made, it is an aging facility that will soon reach full capacity. Expanding the ice core storage facility requires a major investment in infrastructure, which is currently in planning. The ice core science community should be involved in all stages of the design process.
- The community wishes to maintain key boreholes as long-term observatories for conducting measurements with existing and new instruments. GISP2 at Greenland summit is one of the most influential and widely cited records in paleoclimatology, but measurements have shown that the borehole casing is collapsing and already not navigable by most logging instruments. The GISP2 casing should be repaired and maintained for current and future science, as should the casings of boreholes at Siple Dome and Taylor Dome. The IDP Borehole Logging Working Group will produce a prioritized list of existing U.S. boreholes and the scientific reasons for their preservation, and a list of boreholes where instrumentation may be permanently frozen-in, and a process for identifying the fate of future boreholes.
- Drilling ice cores deeper than ~300 m generally requires a drilling fluid mixture that has a density similar to ice to maintain core quality and prevent borehole closure. The fluid must also have a viscosity that is low enough to permit passage of the drill sonde through the fluid many times during the drilling process. Estisol-140 is a fluid that was identified by international partners; it was used at South Pole for the SPICE core project, and it will likely be used for future drilling projects until an improved fluid is identified.

Recommendations

Recommended science goals

1. Past Climate change: Present-day climate change can only be fully understood in context of the past; well-dated histories of climate and the atmosphere over a wide range of time scales are needed to understand climate forcing and response. U.S. scientists are leading investigations in national teams and also in international circles, including generation of international science goals through the International Partnerships in Ice Core Sciences (IPICS). The U.S. ice coring community has always been intimately involved in establishing the IPICS goals. Some of the goals below include U.S. involvement in IPICS targets, while some of the goals below are primarily from members of the U.S. community who, for example, are leading efforts to gain critical samples of ice prior to 800,000 years ago, for evidence of the atmosphere from times when the Earth had 40,000-year climate cycles.

- Drilling of spatially-distributed ice cores and boreholes at many locations to investigate past climate and atmosphere over the past 200 to 40,000 years should continue. Understanding climate signals in remotely-sensed data, understanding climate impacts on the transition from snow to firn to ice on ice sheets, and calibrating high-resolution models, all require arrays of shallow cores covering a range of accumulation and melt rates both in Greenland and in Antarctica; these efforts should continue. Spatially-distributed shallow coring for records ranging from the recent past to 2,000 years will include multiple scientific traverses in Greenland for study of the ice sheet under the currently changing climate. Recent projects in the Arctic include the 1,000-2,000 year annual record from Denali (Mt. Hunter), Alaska, which is providing important constraints on North Pacific climate and tropical teleconnections, during the Medieval Climate anomaly, Little Ice Age, and modern warming. In Antarctica, one example is the recently-completed shallow ice coring at Law Dome, aimed at reconstructing changes in atmospheric oxidizing capacity. Proposed Antarctic science includes an ice coring at Dome C to investigate Holocene changes in the cosmic ray flux.
- Determining patterns of hydroclimate variability, climate feedbacks, and past extent of high-altitude glaciers and aerosol deposition requires ice coring in the Sub-Antarctic Islands, North Pacific coastal mountain ranges, and the Karakoram in Asia.
- Determining the amount of meltwater retained and refrozen in the near surface firn (top ~60 m) on the Greenland Ice Sheet and on the Antarctic Peninsula is critical for improving estimates of surface mass balance under current warming conditions.
- Targeted ice coring to investigate ice, ocean, and atmospheric dynamics in WAIS coastal domes and coastal ice caps and along the dynamic Amundsen Sea Coast of Antarctica, and near Camp Century along the northwest coast of Greenland, are in the planning stages.
- A climate record from the last interglacial period (the Eemian, ~130k to 110k years ago) is key to predicting the response of glaciers and ice sheets to future warming. The search for sites from which to extract Eemian ice in Greenland, both by coring and through horizontal sampling of blue ice ablation zones, should continue. Eemian ice was recovered from the Camp Century core in the 1960's, and an effort to retrieve an intermediate depth ice core from this region is in the planning stages. In Antarctica, extracting a record from Eemian ice is especially important for helping constrain climate and glacial histories of the WAIS during the last interglacial, and is the primary motivation for planned deep drilling at Hercules Dome. An ice core from Hercules Dome would lead to understanding whether the WAIS collapsed during the last interglacial period (MIS5e), and if it did not collapse, then under what climate conditions was it stable? Hercules Dome is the highest-priority next deep ice core for the US community. WAIS history during the Eemian is

poorly known; because large sea level rise due to current climate warming may occur if the WAIS becomes destabilized, an understanding of the WAIS during the last interglacial is urgent.

- Blue-ice paleoclimate records are already providing unlimited samples for atmospheric and ultra-trace component studies and can enable further new types of measurements that have previously been impossible, including analysis of ice older than 800,000 years. Blue-ice studies at Mt. Moulton, Taylor Glacier, and Allan Hills exemplify discoveries from this realm so far; such studies at blue ice sites should continue.
- Ice cores and borehole observations reaching ages between 800,000 years and 1.5 M years (or beyond) are significant, for these data may provide new insight into the effects of greenhouse gases on climate and the observed change in periodicity of glacial cycles during the mid-Pleistocene. The search to identify sites suitable for extracting ancient ice should continue. Extraction of deep ice cores for million-year-old ice, a potential goal for eight to twelve years into the future and a U.S. deep ancient ice site, should be coordinated with international partners through the IPICS "Oldest Ice" project. Currently U.S. scientists are working to retrieve and understand samples of ancient ice from blue ice regions that provide snapshots of climate as it existed more than a million years ago.

2. Ice dynamics and glacial history: Rapid changes in the speed of fast-flowing outlet glaciers and ice streams observed over the past decade create an urgency to understand the dynamics of outlet glaciers and ice sheets. Ice-sheet models that incorporate realistic physics and dynamics at appropriate spatial and temporal scales are needed to predict the "tipping point" when ice-loss becomes irreversible, resulting in ice-sheet collapse and rapid sea-level rise. Observational data are needed to develop and validate the models. Measurements of the ice-bed interface (frozen-thawed, hard-soft bed conditions, sliding, shear), ice-ocean interactions (sub-shelf and basal melting-freezing rates), temperatures and ice deformation properties through the ice, geothermal bedrock conditions and ice-atmosphere interactions (surface mass balance) are key. Another approach to understanding future possible response of ice sheets is to examine their behavior in the past. Dated marine and terrestrial glacial deposits provide information about past ice volume. In regions where such data are not available, histories of ice-sheet thickness and climate can be inferred from radar-detected layers combined with ice core and borehole measurements.

Specific recommendations include:

- Ice-ocean interactions are not fully understood. Boreholes to deploy instruments to measure conditions at ice-ocean interfaces are high priority for investigating ice sheet stability.
- Hydraulic conditions in glaciers and ice sheets exert strong controls on basal motion. Much has been learned through remote sensing methods, but direct measurements through boreholes to the bed are still needed to validate and interpret remote sensing data. Boreholes to the bed at targeted locations are urgently needed to measure geothermal fluxes and basal properties.
- Ice deformation in ice sheets, glaciers, and ice streams depend on temperature and ice rheology. Measurements of ice rheology from ice cores, and borehole logging measurements of temperature, diameter, inclination, and azimuth are needed to provide boundary conditions and constraints for modeling flow of ice sheets and fast-flowing outlet glaciers and ice streams.
- Knowledge of spatial and temporal variations of surface accumulation is critical for quantifying the mass balance of glaciers and ice sheets. Accumulation rate histories derived from short (~200 m) firn and ice cores can be extrapolated spatially to the catchment scale using radar-detected layers. Additional short cores at targeted locations are needed to provide a realistic assessment of surface accumulation over ice-sheet scales.
- Dated ice cores can be used to infer histories of thickness and configuration of ice sheets. Glacial histories contained in coastal ice domes are of particular interest because thickness change near the margins is large. The depth-age relationship from Siple Dome provided key information about

the Holocene deglaciation of the central Ross Embayment, and the depth-age relationship from Roosevelt Island will help constrain the deglaciation of the eastern Ross Embayment. Depth-age profiles from Hercules Dome and other targeted locations are essential for understanding the timing and extent of deglaciation, for example at ice domes near the outflow of the Amundsen Sea Embayment Antarctica, as well as in coastal domes of Greenland.

- The past extent and volume of the Greenland and West Antarctic Ice Sheets is recorded by cosmogenic nuclides in subglacial bedrock. Samples from beneath these ice sheets will provide information on their thickness and configuration during paleoclimates warmer than the present, and help identify their sensitivity to future possible climate change. Short cores of basal ice and bedrock from targeted sites are needed to address questions concerning the extent of the ice sheets during past interglacial climates, and the onset of continental glaciations.

3. Subglacial geology, sediments, and ecosystems: Bedrock, sediments, and ecosystems existing within and beneath ice sheets remain largely unexplored because of the lack of rapid access drills. In particular, the physical conditions at the base of the ice sheets are virtually unknown, but remote sensing of liquid water in subglacial lakes and possibly interconnected hydrologic systems raises concern about thermal conditions and basal slip potential. Likewise, the unknown subglacial geology of Antarctica represents the last continental frontier of geologic exploration, including landscape evolution, past paleoclimates on geological timescales, crustal heat flow, lithospheric stress, ground truth for geophysical imaging, constraints on geodynamical evolution, and relationship with past supercontinents. Information on subglacial biodiversity and biogeochemistry is limited to the Siple Coast, with nothing known about other areas of the Antarctic continent. Subglacial sediments also contain information related to past ice sheet history. Rapid access to subglacial environments is needed to address a wide range of science questions. Specifically,

- Direct sampling of the bedrock is needed to validate models of cratonic growth related to supercontinent assembly in the Mesoproterozoic between about 2.0 and 1.1 billion years ago and for constraining the Phanerozoic geological, tectonic, and exhumation history of the Antarctic continent. Strategic drill-site selection within mapped drainage basins (using products from the BEDMAP2 project) will also allow greater constraints on provenance studies that utilize onshore moraines and offshore glacial strata.
- There exist virtually no heat flow data for Antarctica. Penetration into bedrock provides the first opportunity to accurately measure the geothermal heat flux, which informs us about geotectonic conditions as well as geothermal contributions to ice-sheet temperature.
- Evidence of Cenozoic ice sheet history preserved in sedimentary rocks of subglacial bedrock basins and in sediment deposits within subglacial lakes will provide further dimensions to the records known only from the margins of the continent and will also help to verify paleo-topographic reconstructions for ice sheet modeling. Likewise, access to subglacial bedrock can provide a unique opportunity to study Cenozoic landscape evolution and long-term ice sheet stability using low-temperature thermochronology and cosmogenic-isotope techniques.
- Direct measurements at grounding zones of fast-flowing ice streams and outlet glaciers are badly needed, as are data from sub-ice-shelf ocean cavities in order to provide basic information needed to model ice fluxes near grounding lines and into ice shelves – a critical interface for predicting future ice sheet dynamics.
- Direct measurements of bed conditions including frozen/thawed bed, basal pore pressure, slip, and sediments are needed to develop and test realistic models of the controls on the fast flow of ice streams and outlet glaciers.
- Significant wet environments exist below ice sheets and glaciers; sampling of subglacial habitats including sediments, water, and basal ice is needed to establish the diversity and physiology of

microbes, available nutrients and organic materials, microbial relationships to past climates, and ecosystem function below the ice. Continued support for developing methods and technologies for clean access to subglacial environments and tools for biological and geochemical sampling are needed to investigate these subglacial systems in a clean manner that maintains scientific integrity and environmental stewardship. The recent studies of Whillans and Mercer Subglacial Lakes are steps toward achieving this goal.

4. Ice as a scientific observatory: Polar ice sheets and mid-latitude ice caps archive evidence of past climate and ice dynamics and also serve as a unique platform to conduct observations and experiments concerning seismic activity, planetary sciences and experimental astrophysics, and other novel phenomena. Specifically,

- Borehole logging of both fast-access holes and boreholes originally drilled for ice cores are needed to fully exploit the histories of climate and ice dynamics preserved within the ice. For example, temperature logs are used to infer past temperatures and also the geothermal flux; optical logs yield detailed records of dust and volcanic events and will be important in searches for million year old ice; and sonic logs provide a continuous record of ice fabric and borehole deformation. Community winches to support borehole logging are important assets.
- In-ice physics and astrophysics experiments make use of polar ice as a clean, highly stable, low-background, and transparent (both optically and in the radio frequencies) detection medium for observation of sub-atomic particle interactions. New drilling techniques are under investigation, including cleaner drilling and removal of bubbles from the refrozen water.
- Future planned projects (e.g., the Askaryan Radio Array and Generation-2 Ice Cube) require multiple boreholes drilled to at least 150 m deep (ARA) and 2,500 m deep (G-2IC) and significant calibration studies of the surrounding ice volume. Better understanding of ice attenuation at radio and deep UV wavelengths are particularly desired.
- Ice sheets are a quiet platform for seismic monitoring; the South Pole Remote Earth Science and Seismological Observatory has seismic equipment installed in boreholes about 300 m below the surface. A similar seismic observation network is planned for the Greenland Ice Sheet.
- Novel basal ice structures that have been remotely sensed but whose existence is not well understood should be investigated.

Recommended life cycle cost and logistical principles

Although drills already exist that can achieve some science goals, new drilling technologies are needed to accomplish science goals planned for the next decade. In the past decade there has been an increase in research proposed by the ice science community but the NSF budget has been generally flat. The following principles guiding development of new drills and technologies are recommended:

- Designs require that the supporting logistical needs do not impede execution of the science.
- While developing the science requirements, logistical issues such as weight, size, costs, and time for development, must be clearly defined and transparent at the initial stage of planning. Scientists and engineers working together through IDP must assess the impact of changes as they arise during the engineering design and fabrication process.
- Drills, major drilling subsystems, and accompanying technology must be developed with consideration of potential use in future projects. The drills and technology must be versatile and well documented so that they can be used, maintained, and repaired by other engineers.

- Major drilling systems (e.g., sondes, winches, control and other major electronics systems) should be fungible to the maximum extent possible. Major component inter-changeability and logistical agility should be essential deliverables for all new drilling technology projects.
- Engineering design teams must include individuals with field experience using appropriate ice drilling technology and/or other relevant field experience.
- Increased medium and heavy traversing capability is urgently needed to improve access to many scientifically important regions of the Antarctic and Greenland Ice Sheets.

Recommended Technology Investments

The following investments in drilling technologies are needed to accomplish science goals planned for the next decade. Investments prioritized by time, from consensus of the IDP Science Advisory Board, include:

Priority 1 (needed this year):

- Maintain and upgrade agile equipment in inventory, including: Hand Augers, Sidewinders, the 4" Electromechanical Drills, the 3" Electrothermal Drill, the 3.25" Badger-Eclipse Drills, the Stampfli Drill, Logging Winches, the Small Hot Water Drills, the Blue Ice Drill, the Prairie Dog, the Agile Sub-Ice Geological Drill (ASIG), the Rapid Air Movement Drill (RAM) Drill, and the Winkie Drill.
- Develop the Detailed Design for clean hot water basal ice coring mechanism for a hot water drill.
- Finish a feasibility white paper on logistically effective methods for interdisciplinary projects that seek to retrieve rock, basal ice, sediment, and water from West Antarctic (e.g., Mt. Resnik) and East Antarctic sites (e.g., within Wilkes Basin)
- Develop IDP Science Requirements for collecting a small amount (e.g. chips or less than 10 cm) of sub-ice rock using a lightweight tethered ice core drills, for example the Stampfli drill.
- Develop IDP Science Requirements for collecting a small amount (e.g. 10 cm to 1 meter) of sub-ice rock using an intermediate or deep ice core drill in a fluid-filled hole, for example the Foro 3000 drill.
- Begin construction of the 700 drill.
- Develop the updated IDP Conceptual Design and Detailed Design for a clean Scalable Hot Water drill that minimizes its logistical footprint including fuel supply.
- Establish the IDP Science Requirements for identification and planning of borehole maintenance and fluid maintenance over time.
- Establish the IDP Science Requirements for removing (or lowering) drilling fluid from a borehole (for example for freezing in a sensor).
- Evaluate options for new drilling fluids for Herc Dome and other ice and rock drilling projects.
- Investigate a lighter weight source of power to replace generators for drilling systems, in order to ease demand on logistics, including renewable energy.
- Finish building a stand-alone Foro 3000 Drill as per the IDP Science Requirements.

1624

1625 **Priority 2 (needed in the next 3 years)**

1626 • Build a Scalable Hot Water Access drill for creating access holes in ice that has modular capability for
1627 clean access.

1628 • Identify procurement source and cost for potential purchase of a rapid hole qualifier (temperature
1629 and caliper) for field scientist use in borehole logging applications.

1630 • Resolve logging winch electrical noise issues.

1631 • Finish building a second Blue Ice Drill for wide-diameter drilling to 200 m.

1632 • Continue to evaluate options for exploring/testing shallow drill fluid columns.

1633

1634 **Priority 3 (needed in 3 to 5 years)**

1635 • Continue investigation and modifications of the RAM 2 Drill to achieve the 100 m depth goal
1636 reflected in the system Science Requirements.

1637 • Acquire components for a stand-alone Intermediate Depth Drill (1,850 m) with updated control
1638 system and other repairs so that it can be deployed at the same time as the Foro 3000 Drill.

1639

References

- Achberger, A. M., B. Christner, A. B. Michaud, J. C. Priscu, M. L. Skidmore, and T. J. Vick-Majors. 2016. Microbial community structure of Subglacial Lake Whillans, West Antarctica. *Front. Microbiol.* 7: 1–13. doi:10.3389/fmicb.2016.01457
- Aciego, S.M., K.M. Cuffey, J.L. Kavanaugh, D.L. Morse and J.P. Severinghaus (2007) “Pleistocene ice and paleo-strain rates at Taylor Glacier”, *Antarctica. Quat. Res.*, **68**, 303-313.
- O. Alemany, J. Chappellaz, J. Triest, M. Calzas, O. Cattani, J.F. Chemin, Q. Desbois, T. Desbois, R. Duphil, S. Falourd, R. Grilli, C. Guillerme, E. Kerstel, B. Laurent, E. Lefebvre, N. Marrocco, O. Pascual, L. Piard, P. Possenti, D. Romanini, V. Thiebaut, R. Yamani (2014) “The SUBGLACIOR drilling probe: concept and design”, *Ann. Glaciol.* 55 (68), 233-242.
- Alley, R.B. and I. Joughin, (2012) Modeling Ice-Sheet Flow. *Science*, **336**(6081): 551-552.
- Alley, R.B., S. Anandakrishnan, T.K. Dupont, B.R. Patizek, D. Pollard (2007) “Effect of sedimentation on Ice-sheet grounding-line stability”, *Science*, **315**(5820), 1838-1841.
- Anandakrishnan, S., G. Catania, R.B. Alley, H.J. Hogan (2007) “Discovery of till deposition at the grounding line of Whillans Ice Stream”, *Science*, **315**(5820), 2835-2838.
- Anderson, J.B., H. Conway, P.J. Bart, A.E. Kirshner, S.L. Greenwood, R.M. McKay, B.L. Hall, R.P. Ackert, K. Licht, M. Jakobsson and J.O. Stone. (2014) Ross Sea paleodrainage and deglacial history during and since the LGM. *Quat. Sci. Rev.* <http://dx.doi.org/10.1016/j.quascirev.2013.08.020>
- Bay, R.C., B.F. Price, G.D. Clow, and A.J. Gow. Climate logging with a new rapid optical technique at Siple Dome (2001) *Geophysical Research Letters*, **28**(24), 4635-4638.
- Begeman, C. B., S. M. Tulaczyk, O. J. Marsh, and others. 2018. Ocean Stratification and Low Melt Rates at the Ross Ice Shelf Grounding Zone. *J. Geophys. Res. Ocean.* 123: 7438–7452. doi:10.1029/2018JC013987
- Bentley, M.J., C.J. Fogwill, A.M. Le Brocq, A.L. Hubbard, D.E. Sugden, T.J. Dunai and S.P.H.T. Freeman (2010) “Deglacial history of the West Antarctic Ice Sheet in the Weddell Sea embayment: Constraints on past ice volume change”, *Geology*, **38**(5), 411-414.
- Brook, E. and E. Wolff (2006) “The future of ice coring science”, *EOS Trans. AGU* **87**(4), 39.
- Buizert, C., Baggenstos, D., Jiang, W., Purtschert, R., Petrenko, V.V., Lu, Z.T., Muller, P., Kuhl, T., Lee, J., Severinghaus, J.P., Brook, E.J., (2014). Radiometric Kr-81 dating identifies 120,000-year-old ice at Taylor Glacier, Antarctica. *Proc. Nat. Acad. Sci. U.S.A.* **111**, 6876-6881.
- Carter, S.P. and H.A. Fricker (2012) “The supply of subglacial meltwater to the grounding line of the Siple Coast”, *Annals of Glaciology*, 53 (60) 267-280, doi, 10.3189/2012AoG60A119.
- Christ, A.J., P.R. Bierman, J.M. Schaefer and 15 others (2021) “A multimillion-year-old record of Greenland vegetation and glacial history preserved in sediment beneath 1.4. km of ice at Camp Century”, *PNAS* **118**(13), <https://doi.org/10.1073/pnas.2021442118>.
- Christianson, K., R. W. Jacobel, H. J. Horgan, S. Anandakrishnan and Alley R. B. (2012) “Subglacial Lake Whillans - Ice-penetrating radar and GPS observations of a shallow active reservoir beneath a West Antarctic ice stream”, *Earth and Planetary Science Letters* 331-332(0): 237-245.
- Christianson, Knut, B.R. Parizek, R.B. Alley, H.J. Horgan, R.W. Jacobel, S. Anandakrishnan, B.A. Keisling, B.D. Craig, and A. Muto (2013) Ice sheet grounding zone stabilization due to till compaction. *Geophysical Research Letters*, 40, 5406–5411, doi:10.1002/2013GL057447.
- Christner, B.C., Mikucki, J.A., Foreman, C.M., Denson, J., Priscu, J.C. (2005) Glacial ice cores: A model system for developing extraterrestrial decontamination protocols. *Icarus* 174, 572–584.
- Christner, B. C., Mosley-Thompson, E., Thompson, L. G., & Reeve, J. N. (2001) “Isolation of bacteria and 16S rDNAs from Lake Vostok accretion ice”, *Environmental Microbiology*, 3(9), 570-577.
- Christner, B. C., Mosley-Thompson, E., Thompson, L. G., & Reeve, J. N. (2003) “Bacterial recovery from ancient glacial ice”, *Environmental Microbiology*, 5(5), 433-436.

- Christner, B.C., J. Prisco, A.M. Achberger, C. Barbante, S.P. Carter, K. Christianson, A.B. Michaud, J.A. Mikucki, A.C. Mitchell, M.L. Skidmore, T.J. Vick-Majors, and others (2014). "A microbial ecosystem beneath the West Antarctic Ice Sheet", *Nature* 512, 310-313.
- Clow, G., (2012) Personal communication, 2012.
- Conway, H., B.L. Hall, G.H. Denton, A.M. Gades and E.D. Waddington (1999) "Past and future grounding-line retreat of the West Antarctic Ice Sheet", *Science*, 286(5438), 280-283.
- Cuffey, K.M. and W.S.B. Paterson (2010) *"The Physics of Glaciers"* (4th Ed). Elsevier, ISBN: 978-0-12-369461, 704pp.
- D'Andrilli, J., Foreman, C.M., Sigl, M., Prisco, J.C., McConnell, J.R.(2017). "A 21 000-year record of fluorescent organic matter markers in the WAIS Divide ice core." *Climate of the Past* 13(5): 533-544.
- D'Andrilli, J., Smith, H.J., Diesner, M., Foreman, C.M. (2017). "Climate driven carbon and microbial signatures through the last ice age." *Geochemical Perspectives Letters* 4: 29-34.
- D'Andrilli, J. and McConnell, J.R. (2020) "Polar ice core organic matter signatures reveal past atmospheric carbon composition and spatial trends across ancient and modern timescales" *Journal of Glaciology*: In Review.
- Dahl-Jensen et al (2013) "Eemian interglacial reconstructed from a Greenland folded ice core". *Nature* 493(7433). p.489-494
- Denton, G.H., J.C. Bockheim, S.C. Wilson and M. Stuiver (1989) "Late Wisconsin and Early Holocene glacial history, inner Ross Embayment, Antarctica", *Quat. Res.* **31**, 151-182.
- Denton, G.H., D.E. Sugden, D.R. Marchant, B.L. Hall and T.I. Wilch (1993) "East Antarctic Ice Sheet sensitivity to Pliocene climatic change from a Dry Valleys perspective", *Geografiska Annaler*, **75**, 155-204.
- Dubnick, A., S. Kazemi, M. Sharp, J. Wadham, J. Hawkings, A. Beaton, and B. Lanoil (2017), Hydrological controls on glacially exported microbial assemblages, *J. Geophys. Res. Biogeosci.*, 122, 1049-1061, doi:10.1002/2016JG003685.
- Dunbar, N.W., W.C. McIntosh and R.P. Esser (2008) "Physical setting and tephrochronology of the summit caldera ice record at Mount Moulton, West Antarctica", *Geol. Soc. of Amer. Bull.*, **120**(7-8), 796-812.
- Engelhardt, H., N. Humphrey, B. Kamb, M. Fahnestock (1990) "Physical conditions at the base of a fast moving ice stream", *Science*, **248**(4951), 57-59.
- Engelhardt, H., and B. Kamb (1998) "Basal sliding of Ice Stream B, West Antarctica", *J. Glaciol.* 44(147), 223-230.
- Eshelman, E.J., Malaska, M.J., Manatt, K.S., Doloboff, I.J., Wanger, G., Willis, M.C., Abbey, W.J., Beegle, L.W., Prisco, J.C., Bhartia, R. (2019) "WATSON: In Situ organic detection in subsurface ice using Deep-UV fluorescence spectroscopy" *Astrobiology* 19(6): 771-784.
- Fahnestock, M., W. Abdalati, I. Joughin, J. Brozena and P. Goginini (2001) "High geothermal heat flow, basal melt and the origin of rapid ice flow in central Greenland" *Science*, 294(2338), doi:10.1126/science.1065370.
- Favier, L., G. Durand, S.L. Cornford, G.H. Gudmundsson, O. Gagliardini, F. Giller-Chaulet, T. Zwinger, A.J. Payne and A.M. Le Brocq, (2014) Retreat of Pine Island Glacier controlled by marine ice-sheet instability, *Nature Climate Change*, doi:10.1038/nclimate2094
- Fisher, A.T., K.D. Mankoff, S.M. Tulaczyk, S.W. Tyler, N. Foley (2015) High geothermal flux measured below the West Antarctic Ice Sheet, *Science Advances*, 1(60), P.e1500093.
- Gagliardini, O., G. Durand, T. Zwinger, R.C.A. Hindmarsh and E. Le Meur. (2010) Coupling of ice-shelf melting and buttressing is a key process in ice-sheets dynamics. *Geophys. Res. Lett.* **37**, doi:10.1029/2010GL043334.
- Grannas, A.M., Hockaday, W.C., Hatcher, P.G., Thompson, L.G., Mosley-Thompson, E. (2006). "New revelations on the nature of organic matter in ice cores." *Journal of Geophysical Research Atmospheres* 111(D4).

- Hall, B., C. Baroni and G. Denton (2004) "Holocene relative sea-level history of the southern Victoria Land coast", *Antarctica: Glob. Plan. Change*, 42, 241-263.
- Hawkings, J. R., M. L. Skidmore, J. L. Wadham, and others. 2020. Enhanced trace element mobilization by Earth's ice sheets. *Proc. Natl. Acad. Sci. U. S. A.* 117: 31648–31659. doi:10.1073/pnas.2014378117
- Higgins, J.A., Kurbatov, A.V., Spaulding, N.E., Brook, E., Introne, D.S., Chimiak, L.M., Yan, Y.Z., Mayewski, P.A., Bender, M.L., (2015). Atmospheric composition 1 million years ago from blue ice in the Allan Hills, Antarctica. *Proc. Nat. Acad. Sci.* **112**, 6887-6891.
- Hodson, T. O., R. D. Powell, S. A. Brachfeld, S. Tulaczyk, and R. P. Scherer. 2016. Physical processes in Subglacial Lake Whillans, West Antarctica: Inferences from sediment cores. *Earth Planet. Sci. Lett.* 444: 56–63. doi:10.1016/j.epsl.2016.03.036
- Horgan, H. J., S. Anandakrishnan, R. W. Jacobel, K. Christianson, R. B. Alley, D. S. Heeszel, S. Picotti and Walter J. I. (2012) "Subglacial Lake Whillans - Seismic observations of a shallow active reservoir beneath a West Antarctic ice stream. *Earth and Planetary Science Letters* 331-332(0): 201-209.
- Horgan, H.J., K. Christianson, R.W. Jacobel, S. Anandakrishnan and R.B. Alley (2013) Sediment deposition at the modern grounding zone of Whillans Ice Stream, West Antarctica. *Geophysical Research Letters*, 40, 1–6, doi:10.1002/grl.50712.
- Ice Core Working Group – ICWG (2003) "United States Ice Core Science: Recommendations for the future, Univ. of New Hampshire", 48pp. <http://nicl-smo.unh.edu/icwg/ICWG2003.pdf>
- Jacobel R, Welch B, Steig EJ, Schneider DP (2005). Hercules Dome, Antarctica – an optimal site for deep ice core drilling. *Journal of Geophysical Research – Earth Surface* **110**: F01015, doi:10.1029/2004JF000188.
- Jenkins, A., P. Dutrieux, S.S. Jacobs, S.D. McPhail, J.R. Porrett, A.T. Webb and D. White (2010) "Observations beneath Pine Island Glacier in West Antarctica and implications for its retreat", *Nature Geoscience*, doi:10.1038/NGE0890.
- Joughin, I., B.E. Smith & B. Medley (2014). Marine ice sheet collapse potentially underway for the Thwaites Glacier basin, West Antarctica. *Science*, **344**(6185), 735-738 DOI: 10.1126/science.1249055
- Kamb, B. (2001) "Basal zone of the West Antarctic ice streams and its role in lubrication of their rapid motion" *In* Alley, R.B. and R.A. Bindshadler, eds. *The West Antarctic ice sheet: behavior and environment*. Washington, DC, American Geophysical Union, 157–199. (Antarctic Research Series 77).
- Kerhl, L, H.Conway, N. Holschu, S. Campbell, A.V. Kurbatov, N.E. Spaulding (2018). Evaluating the duration and continuity of potential climate records from the Allan Hills Blue Ice Area, East Antarctica. *Geophysical Research Letters* 45, <https://doi.org/10.1029/2018GL077511>.
- King, A.C.F., Thomas, E.R., Pedro, J.B., Markle, B., Potocki, M., Jackson, S.L., Wolff, E., Kalberer, M. (2019). "Organic Compounds in a Sub-Antarctic Ice Core: A Potential Suite of Sea Ice Markers." 46(16): 9930-9939.
- Korotkikh, E.V., Mayewski, P.A., Handley, M.J., Sneed, S.B., Introne, D.S., Kurbatov, A.V., Dunbar, N.W., McIntosh, W.C., (2011). The last interglacial as represented in the glaciochemical record from Mount Moulton Blue Ice Area, West Antarctica. *Quatern. Sci. Rev.* **30**, 1940-1947.
- Liu, B., F.M. Phillips, J.T. Fabryka-Martin, M.M. Fowler and W.D. Stone (1994) "Cosmogenic ³⁶Cl accumulation in unstable landforms 1. Effects of the thermal neutron distribution", *Water Res. Res.*, **30**, 3115-3125.
- Malaska, M.J., Bhartia, R., Manatt, K.S., Priscu, J.C., Abbey, W.J., Mellerowicz, B., Palmowski, J., Paulsen, G.L., Zacny, K., Eshelman, E.J., and D'Andrilli, J. "Subsurface in situ detection of microbes and diverse organic matter hotspots in the Greenland ice sheet", *Astrobiology*: In Press.
- Marteinsson, V. T., Á. Rúnarsson, A. Stefánsson, and others. 2013. Microbial communities in the subglacial waters of the Vatnajökull ice cap, Iceland. *ISME J.* 7: 427–37. doi:10.1038/ismej.2012.97

- 1783
1784 Mercer, J.H. (1968) "Glacial geology of the Reedy Glacier area, Antarctica", *Geol. Soc. Amer. Bull.* **79**, 471-
1785 486.
- 1786 Mengel, M. and Levermann, A., (2014) Ice plug prevents irreversible discharge from East Antarctica,
1787 *Nature Climate Change*, **4**, doi 10.1038/nclimate2226.
- 1788 Michaud, A. B., M. L. Skidmore, A. C. Mitchell, T. J. Vick-Majors, C. Barbante, C. Turetta, W. vanGelder,
1789 and J. C. Priscu. 2016. Solute sources and geochemical processes in Subglacial Lake Whillans, West
1790 Antarctica. *Geology* G37639.1. doi:10.1130/G37639.1
- 1791 Michaud AB, J.E. Dore, A.M. Achberger, B.C. Christner, A.C. Mitchell, M.L. Skidmore, et al. (2017).
1792 Microbial oxidation as a methane sink beneath the West Antarctic Ice Sheet. *Nat. Geosci.* 10:582–586
- 1793 Mikucki, J. A., P. A. Lee, D. Ghosh, and others. 2016. Subglacial Lake Whillans microbial biogeochemistry:
1794 A synthesis of current knowledge. *Philos. Trans. R. Soc. A Math. Phys. Eng. Sci.* 374.
1795 doi:10.1098/rsta.2014.0290
- 1796 Miteva, V. I., Sheridan, P. P., & Brenchley, J. E. (2004) "Phylogenetic and physiological diversity of
1797 microorganisms isolated from a deep Greenland glacier ice core", *Applied and Environmental*
1798 *Microbiology*, **70**(1), 202-213.
- 1799 Miteva, V. I., & Brenchley, J. E. (2005) "Detection and isolation of ultrasmall microorganisms from a
1800 120,000-year-old Greenland glacier ice core", *Applied and Environmental Microbiology*, **71**(12), 7806-
1801 7818.
- 1802 Morse, D.L., Blankenship, D.D., Waddington, E.D., Neumann, T.A. (2002) "A site for deep ice coring in West
1803 Antarctica: results from aerogeophysical surveys and thermo-kinematic modeling", *Annals Glaciol.* **35**,
1804 36-44.
- 1805 NRC (2002) "Abrupt climate change: inevitable surprises", *National Academies Press*, Washington, D.C.
- 1806 NRC (2007) "Exploration of Antarctic Subglacial Aquatic Environments: Environmental and Scientific
1807 Stewardship", *National Academies Press*, Washington, D.C.
- 1808 Pattyn, F. and 27 others (2013) Grounding-line migration in plan-view marine ice-sheet models: results of
1809 the ice2sea MISIMP3d intercomparison, *J. Glaciol.*, **59**, doi:10.3189/2013JoG12J129.
- 1810 Payne, A.J., A. Vieli, A.P. Shepherd, D.J. Wingham, and E. Rignot (2004) "Recent dramatic thinning of
1811 largest West Antarctic ice stream triggered by oceans", *Geophys. Res. Lett.* **31**, L23401, doi:10.1029/
1812 2004GL021284.
- 1813 Petrenko, V.V., A.M. Smith, H. Schaefer, K.Riedel, E.J. Brook, D. Baggenstos, C. Harth, Q. Hua, C. Buizert,
1814 A. Schilt, X. Fain, L. Mitchell, T. Bauska, A. Orsi, R.F. Weiss, J.P. Severinghaus. Minimal geological
1815 methane emissions during the Younger Dryas – Preboreal abrupt warming event. 2017. *Nature* **548**,
1816 443-446.
- 1817 Pettit, E.C., E. Waddington, T. Thorsteinsson, A. Gusmeroli, J. Kennedy, C. Ritz, and R. Carns. (2011) Using
1818 Borehole Sonic Logging to Infer Ice Microstructure and Climate History. EGU General Assembly.
1819 Abstract EGU2011-14160.
- 1820 Pollard, D. and R.M. DeConto (2009) "Modeling West Antarctic ice sheet growth and collapse through the
1821 past five million years", *Nature*, **458**, doi:10.1038/nature07809.
- 1822 Price, S.F., H. Conway and E.D. Waddington (2007) "Evidence for late Pleistocene thinning of Siple Dome,
1823 West Antarctica", *J. Geophys. Res.* **112**, doi:10.1029/2006JF000725.
- 1824 Price, S.F., H. Conway, E.D. Waddington and R.A. Bindshadler (2008) "Model investigations of inland
1825 migration of fast-flowing outlet glaciers and ice streams", *J. Glaciol.* **54**(184), 49-60.
- 1826 Priscu, J. C., Achberger, A. M., Cahoon, J. E., Christner, B. C., Edwards, R. L., Jones, W. L., Michaud, A.B.,
1827 Siegfried, M.R. Skidmore, M.L., Spigel, R.H., Switzer, G.W. Tulaczyk, S. & Vick-Majors, T. J. (2013) A
1828 microbiologically clean strategy for access to the Whillans Ice Stream subglacial environment. *Antarctic*
1829 *Science*, **25**(05), 637-647.

- 1830 Rignot, E., J. Mouginot, M. Morlighem, H. Seroussi & B. Scheuchl, (2014) Widespread, rapid grounding-
1831 line retreat of Pine Island, Thwaites, Smith and Kohler Glaciers, West Antarctica, from 1992-2011,
1832 *Geophys. Res. Lett.*, **41**, doi:10.1002/2014GL060140.
- 1833 Santibáñez, PA, O.J. Maselli, M.C. Greenwood, et al. Prokaryotes in the WAIS Divide ice core reflect source
1834 and transport changes between Last Glacial Maximum and the early Holocene. *Glob Change Biol.* 2018;
1835 24: 2182– 2197.
- 1836 Scambos, T.A., R.E. bell, R.B. Alley, S. Anandakrishnan, D.H. Bromwich, K. Brunt, K. Christianson, T. Creyts,
1837 S.B. Das, R. DeConto, P. Dutrieux, H.A. Fricker, D. Holland, J. MacGregor, B. Medley, J.P. Nicolas,
1838 D.Pollard, M.R. Siegfried, A.M. Smith, E.J. Steig, L.D. Trusel, D.G. Vaughan, P.L. Yager. 2017. How much,
1839 how fast?: a science review and outlook for research on the instability of Antarctica’s Thwaites Glacier
1840 in the 21st century. *Global and Planetary Change* 153, 16-34.
- 1841 Schaefer, J.M., Finkel, R., Balco, G., Alley, R.B., Caffee, M., Briner, J.P., Young, N.E., Gow, A.J., and Schwartz,
1842 R. (2016). GISP2 bedrock reveals extended periods of ice-free Greenland during the Pleistocene.
1843 *Nature*, 540, 252-255.
- 1844 Scherer, R. P., Aldahan, A., Tulaczyk, S., Possnert, G., Engelhardt, H., and Kamb, B., (1998). Pleistocene
1845 collapse of the West Antarctic ice sheet. *Science*, 281(5373), 82-85.
- 1846 Scherer, R. P., Harwood, D. M., Ishman, S. E., and Webb, P. N., 1988. Micropaleontological analysis of
1847 sediments from the Crary Ice Rise, Ross Ice Shelf. *Antarctic Journal of the United States*, 23(5), 34-36.
- 1848 Shepherd, A., D. Wingham and E. Rignot (2004) “Warm ocean is eroding West Antarctic Ice Sheet”,
1849 *Geophys. Res. Lett.* **31**, L23402, doi:10.1029/2004GL021106.
- 1850 Simon, C., Wiezer, A., Strittmatter, A. W., & Daniel, R. (2009) “Phylogenetic diversity and metabolic
1851 potential revealed in a glacier ice metagenome”, *Applied and environmental microbiology*, 75(23),
1852 7519-7526.
- 1853 Spaulding, N.E., J.A. Higgins, A.V. Kurbatov, M.L. Bender, S.A. Arcone, S.Campbell, N.W. Dunbar, L.M.
1854 Chimiak, D.S. Introne & P.A. Mayewski. (2013) “Climate Archives From 90 to 250 Ka in Horizontal and
1855 Vertical Ice Cores From the Allan Hills Blue Ice Area, Antarctica.” *Quaternary Research* 80 (3): 562–74.
- 1856 Stanton, T.P., W. J. Shaw, M. Truffer, H.F.J. Corr, L.E. Peters, K.L. Riverman, R. Bindshadler, D.M. Holland
1857 & S. Anandakrishnan, (2013) Channelized ice melting in the ocean boundary layer beneath Pine Island
1858 Glacier, Antarctica. *Science*, **341**, 1236-1239.
- 1859 Steig, E.J., K. Huybers, H.A. Singh, N.J. Steiger, D.M.W. Frierson, T. Popp, J.C.W. White (2015). Influence of
1860 West Antarctic Ice Sheet collapse on Antarctic surface climate. *Geophysical Research Letters* **42**: 4862–
1861 4868, doi: 10.1002/2015GL063861 (2015).
- 1862 Stone, J.O., G.A. Balco, D.E. Sugden, M.W. Caffee, L.C. Sass III, S.G. Cowdery and C. Siddoway (2003)
1863 Holocene deglaciation of Marie Byrd Land, West Antarctica. *Science*, **299**(5603), 99-102.
- 1864 Todd, C., J. Stone, H. Conway, B. Hall and G. Bromley (2010) “Late Quaternary evolution of Reedy Glacier,
1865 Antarctica”, *Quat. Sci. Rev.*, **29**(11-12), 1328-1341.
- 1866 Truffer, M. W. Harrison, K. Echelmeyer (1999) “Glacier motion dominated by processes deep in underlying
1867 till”, *J. Glaciol.* **46**(153), 213-221.
- 1868 Truffer, M. W. Harrison (2006) “In situ measurements of till deformation and water pressure”, *J. Glaciol.*
1869 **52**(177), 175-182.
- 1870 Venturelli, R. A., M. R. Siegfried, K. A. Roush, and others. 2020. Mid-Holocene Grounding Line Retreat
1871 and Readvance at Whillans Ice Stream, West Antarctica. *Geophys. Res. Lett.* 47.
1872 doi:10.1029/2020GL088476
- 1873 Vick-Majors TJ, A.C. Mitchell, A.M. Achberger, B.C. Christner, J.E. Dore, A.B. Michaud, et al. (2016).
1874 “Physiological ecology of microorganisms in Subglacial Lake Whillans”. *Front. Microbiol.* 7:1–16.).
- 1875 Vick-Majors, T. J., A. B. Michaud, M. L. Skidmore, and others. 2020. Biogeochemical Connectivity
1876 Between Freshwater Ecosystems beneath the West Antarctic Ice Sheet and the Sub-Ice Marine
1877 Environment. *Global Biogeochem. Cycles* 34. doi:10.1029/2019GB006446

- 1878 Vogel, A.L., Lauer, A., Fang, L., Arturi, K., Bachmeier, F., Daellenback, K.R., Kaser, T., Vlachou, A.,
1879 Pospisilova, V., Baltensperger, U., Haddad, I.E., Schwikowski, M., Bjelic, S. (2019). "A Comprehensive
1880 Nontarget Analysis for the Molecular Reconstruction of Organic Aerosol Composition from Glacier Ice
1881 Cores." *Environmental Science & Technology* 53(21): 12565-12575.
1882 Waddington, E.D., T.A. Neumann, M.R. Koutnik, H-P Marshall and D.L. Morse (2007) "Inference of
1883 accumulation-rate patterns from deep layers in glaciers and ice sheets", *J. Glaciol.* **53**(183), 694-712.
1884 Waddington, E.D., H. Conway, E.J. Steig, R.B. Alley, E.J. Brook, K.C. Taylor and J.W.C. White (2005)
1885 "Decoding the dipstick: thickness of Siple Dome, West Antarctica at the Last Glacial Maximum",
1886 *Geology* **33**(4), 281-284.
1887 Wadham JL, S. Arndt, S. Tulaczyk, M Stibal, M Tranter, J Telling , et al. (2012). Potential methane reservoirs
1888 beneath Antarctica. *Nature* 488:633–7.
1889 WAIS Divide Project Members, 2013. Onset of deglacial warming in West Antarctica driven by local orbital
1890 forcing. *Nature*. **500**, 440-444, doi:10.1038/nature12376
1891 WAIS Divide Project Members, 2015. Precise interhemispheric phasing of abrupt climate change during
1892 the last ice age. *Nature* 520(7549), 661-665, doi 10.1038/nature14401.
1893 Webb, P.N., D.M. Harwood, B.C. McKelvey, J.H. Mercer and L.D Stott (1984) "Cenozoic marine
1894 sedimentation and ice-volume variation on the East Antarctic Craton", *Geology*, **12**, 287-291.
1895 Xu, J., Grannas, A.M., Xiao, C., Du, Z., Willoughby, A., Hatcher, P.G., An, W. (2018). "Highresolution mass
1896 spectrometric characterization of dissolved organic matter from warm and cold periods in the NEEM
1897 ice core." *10*(1): 38-46.
1898 Yau, A.M., M.L. Bender, D.R. Marchant, S.L. Mackay (2015) "Geochemical analysis of air from an ancient
1899 debris-covered glacier, Antarctica", *Quaternary Geochronology* 28, 29-39.
1900

1901

1902 **Acronyms**

1903

1904 AGAP: Antarctica's Gamburtsev Province

1905 ANDRILL: Antarctic Drilling Project

1906 AO: Arctic Oscillation

1907 ARA: Askaryan Radio Array

1908 ARIANNA: Antarctic Ross Ice shelf Antenna Neutrino Array

1909 ASIG: Agile Sub-Ice Geological (drill)

1910 AUV: Autonomous Underwater Vehicle

1911 BLWG: Borehole Logging Working Group

1912 CReSIS: Center for Remote Sensing of Ice Sheets

1913 DISC: Deep Ice Sheet Coring

1914 DOSECC: Drilling, Observation, Sampling of the Earths Continental Crust (drilling service)

1915 EDC: EPICA Dome C

1916 EGRIP: East Greenland Ice core Project

1917	EHWD: Enhanced Hot Water Drill
1918	ENSO: El Niño Southern Oscillation
1919	EPICA: European Project for Ice Coring in Antarctica
1920	G-2IC: Generation-2 Ice Cube
1921	GISP2: Greenland Ice Sheet Program II
1922	GRIP: Greenland Ice Core Project
1923	GZK: Greisen-Zatsepin-Kuzmin
1924	HCFC: Hydrochlorofluorocarbon
1925	ICECAP: A project name, not an acronym
1926	ICWG: Ice Core Working Group
1927	IDP: Ice Drilling Program
1928	IPCC: Intergovernmental Panel on Climate Change
1929	IPICS: International Partnerships in Ice Core Sciences
1930	LIG: Last Interglacial
1931	LRSP: Long Range Science Plan
1932	NEEM: North Greenland Eemian Ice Drilling
1933	NEGIS: Northeast Greenland Ice Stream
1934	NGRIP: North Greenland Ice Core Project
1935	NRC: National Research Council
1936	NSF: National Science Foundation
1937	PINGU: Precision IceCube Next Generation Upgrade
1938	RAID: Rapid Access Ice Drill
1939	RAM: Rapid Air Movement (drill)
1940	ROV: Remotely Operated Vehicle
1941	SAB: Science Advisory Board
1942	SALE: Subglacial Antarctic Lake Environment
1943	SCAR: Scientific Committee on Antarctic Research
1944	SHALDRIL: Shallow Drilling on the Antarctic Continental Margin
1945	SleGE: Sub-Ice Geological Exploration
1946	SPICE: South Pole Ice
1947	WAIS: West Antarctic Ice Sheet
1948	WISSARD: Whillans Ice Sheet Subglacial Access Research Drilling